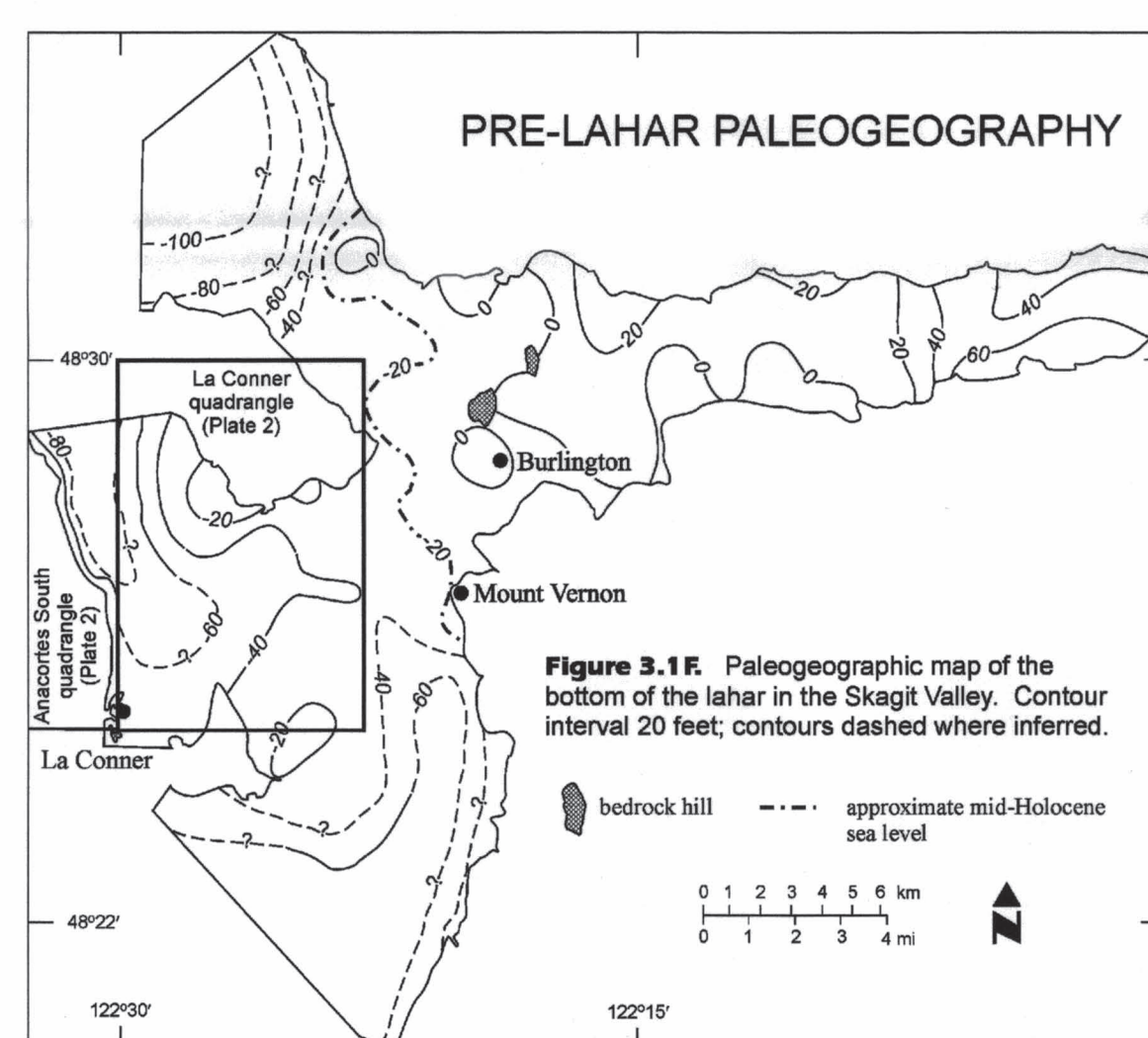
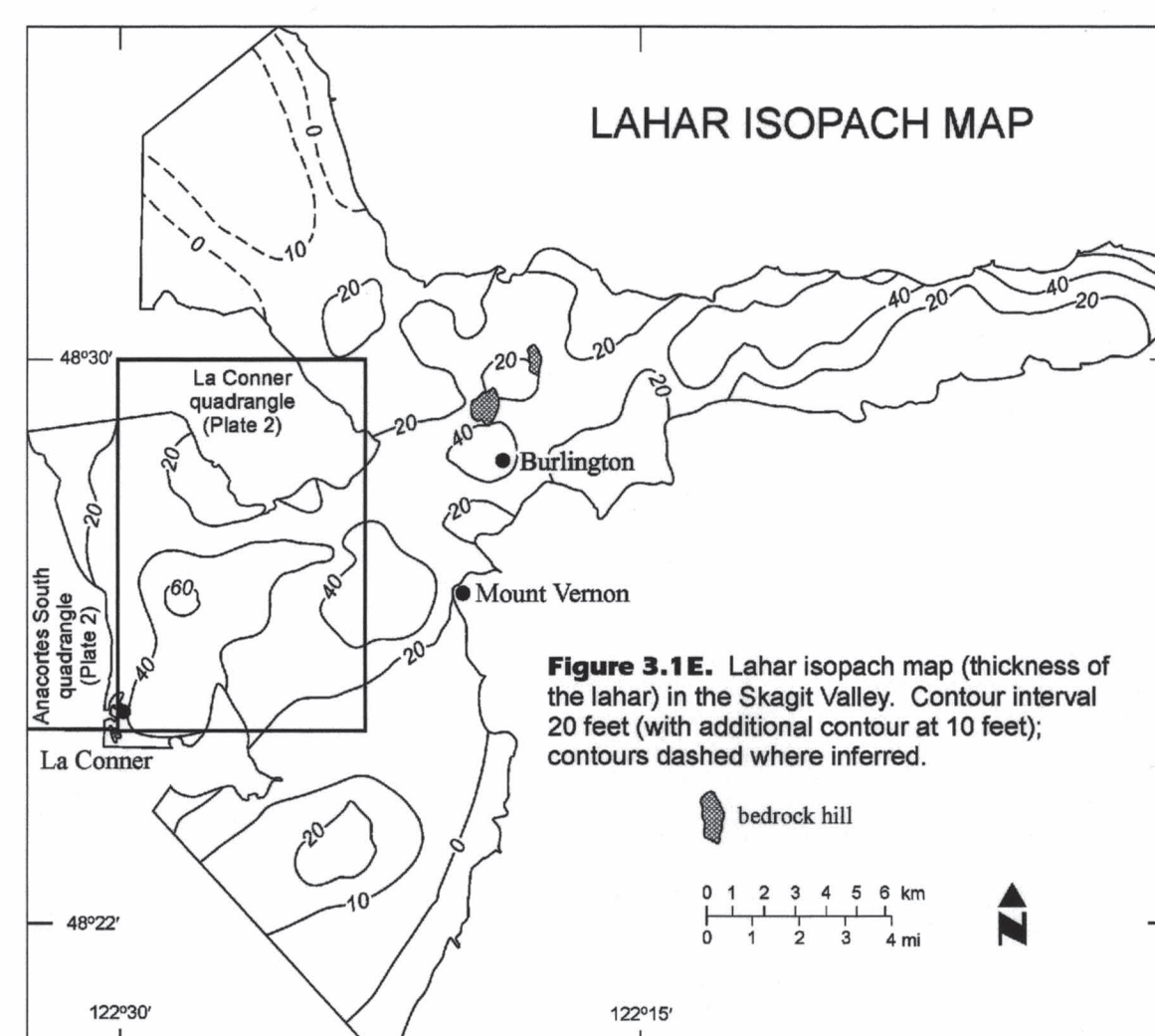
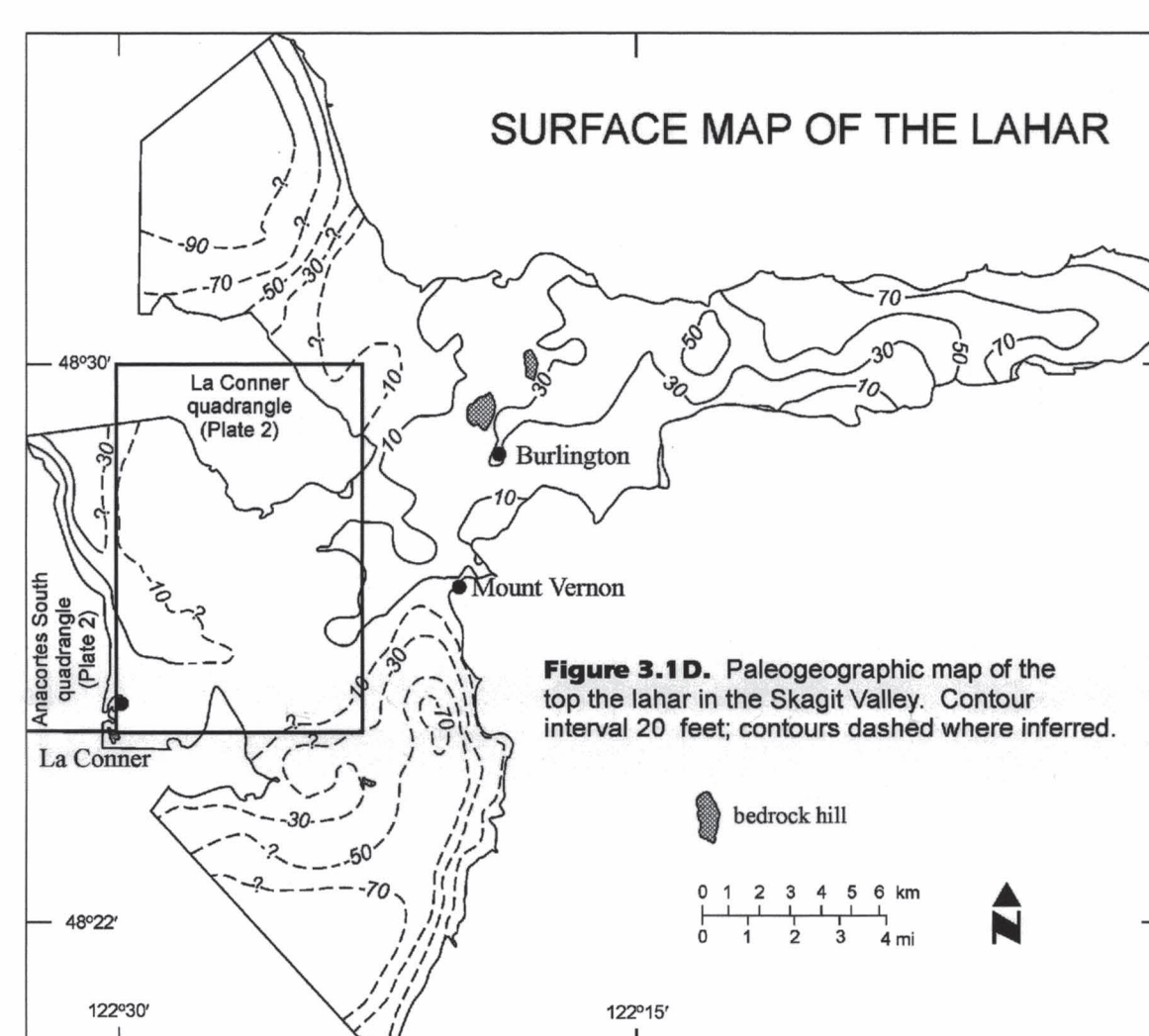
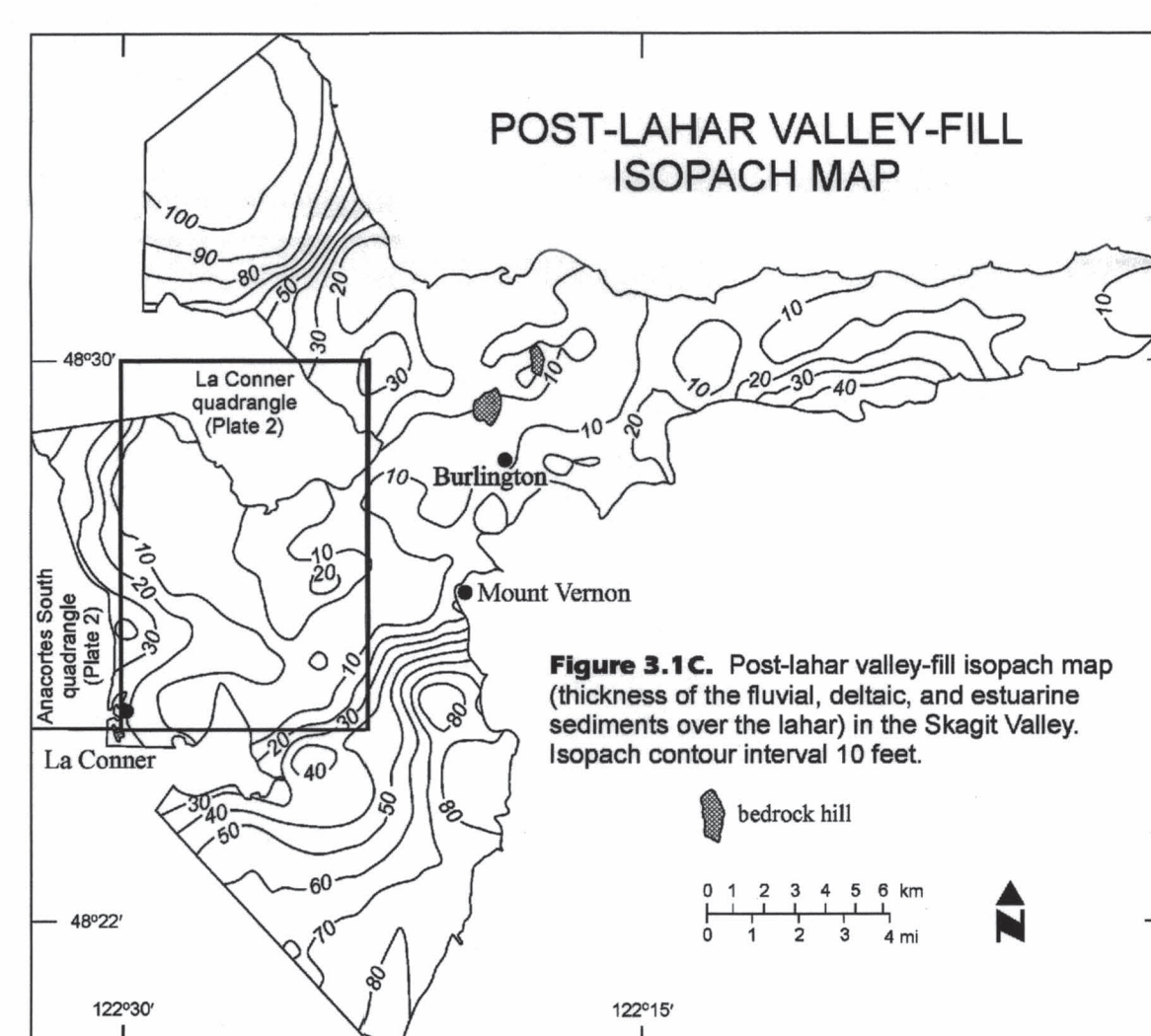
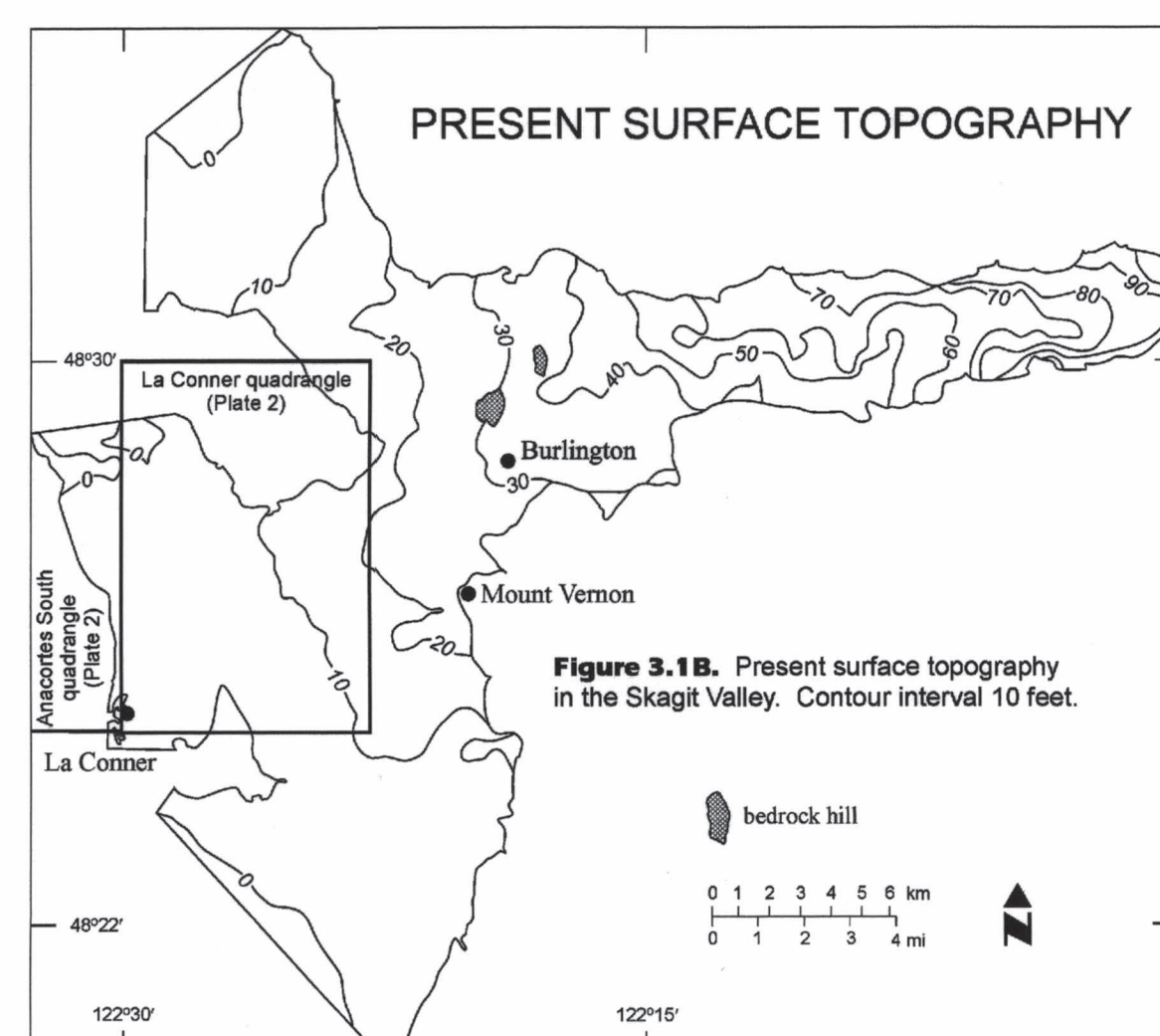
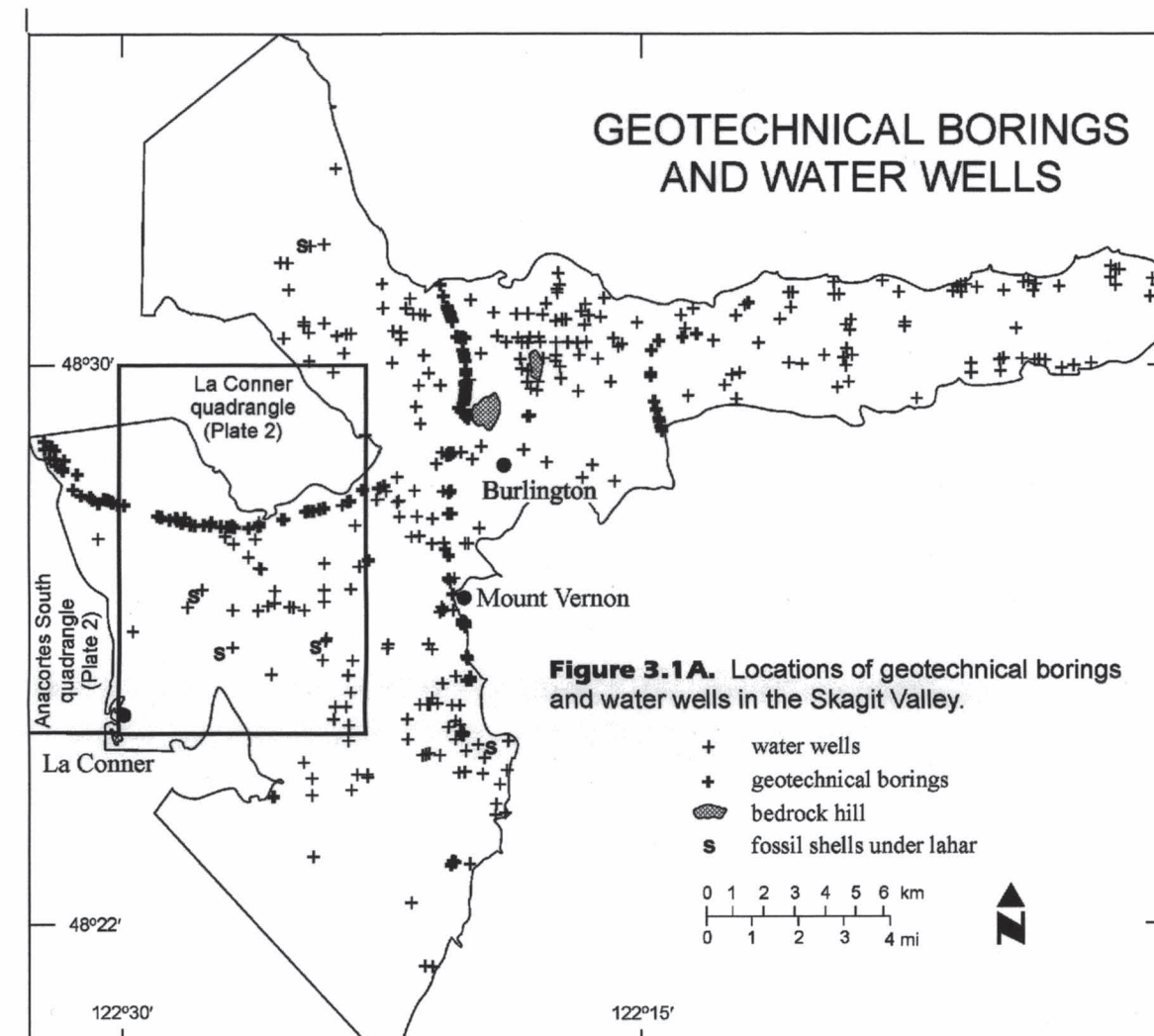


LAHAR INUNDATION OF THE LOWER SKAGIT RIVER VALLEY

MAP B

Geologic mapping compiled from Beget (1981), Tabor and Crowder (1969), Tabor and others (1988) and this study



OVERVIEW

Deposits originating from a Glacier Peak eruptive event(s) are preserved in the lower Skagit River valley at 3- to 15-m-high terraces adjacent to the Skagit River flood plain. The cities of Lyman, Sedro-Wooley, and Burlington rest on terraces composed of dacitic-rhyolite sand and sandy gravel that we interpret as nonchocolate lahars runout. New optically, $^{40}\text{Ar}/^{39}\text{Ar}$ ages¹, stratigraphic relations, dacite-rhyolite petrographic and geochemical data², and laharic sand composition³ suggest that the voluminous Kennedy Creek [volcanic] assemblage (KCA) of Beget (1981) reached Puget Sound. The KCA is mainly a lahar/proclastic surge deposit(s) that originated from a Glacier Peak vent during the Common Age (Pleistocene, 13.5 ka). The KCA is a 100-km² deposit, but distal lahars can be traced as a 10- to 18-m-thick semi-continuous stratum on the basis of well and boring logs and isolated outcrops. This distal lahar is overlain by estuarine or deltaic sediments.

The Skagit Valley sequence of the lalar appears to have originated from a single eruptive episode on the basis of an upward-fining sequence within the lalar, from crystal-dacite lithic sand/gravel (50–80% dacite clasts) to sand (41–75% dacite clasts). Hyphrethene-hornblende phryic, vesicular dacite-dacite samples collected at 21 lalar sites (from near Glacier Peak to La Conner) have nearly identical rare earth, trace, and major element geochemistry, suggesting of a single eruptive event. Charcoal in lalaric sand near Lyman yielded an age of 4,780 ± 80 yr B.P., similar to the 5,100–5,500 yr B.P. age of the KCA and thus may be part of the KCA. Along the flanks of Glacier Peak, Begert (1982) estimated a total original volume of 2–3 km³ for the KCA. Dusty Creek and Baekos Creek assemblages of Begert (1982) (The >340 and >100 µm sized fractions) yielded an original volume of >10 km³ and could correlate with the KCA. Maximum-minimum volume lalar estimates in the lower Skagit River basin suggest an additional 2 to 3 km³ for the KCA. Thus, if our correlations are correct, the preserved KCA may have a volume exceeding 15 km³.

Beget (1982) indicated that the KCA consists of noncohesive and cohesive laharc deposits, several pyroclastic flow deposits, with lesser laeunine deposits, and alluvium. The assemblage contains abundant clasts of gray hypersthene-hornblende dacite. Banded dacites, consisting of alternating layers of frothy light gray to white and darker nonvesicular dacite, are present in some of the laharc and pyroclastic flow deposits. The KCA contains hornblende with lesser hypersthene and rare olivine similar to the dacitic sands and vesicular dacite clasts of laharc-runout deposits in the lower Skagit River valley. Beget (1982) noted that olivine is a minor mineral in the KCA and is not present in many of the dacite clasts of the assemblage. We observed rare olivine in the dacite clasts from samples of the laharc deposit(s) in this study. The KCA forms terraces in the White Chuck, Sauk, and Skagit River valleys. The deposits are tens to hundreds of meters thick and are composed of dacite, andesite, and basaltic andesite. The KCA and the Dusty Creek assemblages of Beget (1982) may also have formed as a result of the 5,100 to 5,200 yr B.P. eruptive event that formed the KCA. If these other assemblages were part of the same eruptive event as the KCA, then eruptive products inundated most of the major drainages emanating from Glacier Peak. Correlation of the KCA with the laharc runouts of the lower Skagit River valley is strengthened by Beget's (1982) correlation of the KCA with a dacite-rich laharc at Minkler Lake (Map B, near Myman), "a terrace east of [Minkler] Lake is composed almost entirely of dacite-rich [laharic and alluvial deposits] and contains clasts as much as 20 cm in diameter that are similar to those of the 'KCA' assemblage." The KCA is composed of the lower, coarser grained of the hyperconcentrated flood deposits. (His site is indicated by the arrow symbol near the bottom of the terrace directly south of sites 79P and 75S on Map B.) We herein extend the KCA to near the present Puget Sound.

Dacite clasts in the lahar are composed of quartz phenocrysts (subrounded and resorbed volcanic grains), plagioclase phenocrysts (blocky rectangular grains with distinct concentric zoning), and plagioclase microlites in a glassy matrix. Hypersthene phenocrysts occur as isolated single blocky crystals and less commonly as glomerocrysts with plagioclase and are

See Plate 4 for references cited.

* Data in Washington Department of Natural Resources files for future publication.

very light green inplane-polarized light. The matrix is glassy and clear in plane-polarized light and includes microlites of plagioclase with lesser hornblende and hypersthene. Glassy magmatic flow features are conspicuous. Spherical, elliptical, and tear-drop shaped vesicles contribute to the magmatic flow appearance of these vesicular dacites.

Lahar-runout sands are distinctly dacite volcanic lithic rich*. Microlitic volcanic fragments (dacite), as well as the plagioclase, hypersthene, and some hornblende in the lahar runout sand grains are petrographically similar to the gravel-size dacite clasts in the lahar. The mineralogical similarity of the dacite lithics and relict volcanic phenocrysts in both the Skagit River alluvium and lahar-runout deposits and the abundance of these clasts in the Skagit River alluvium reflect active reworking of the laharic-volcanoanogenic sediment by the Skagit River.

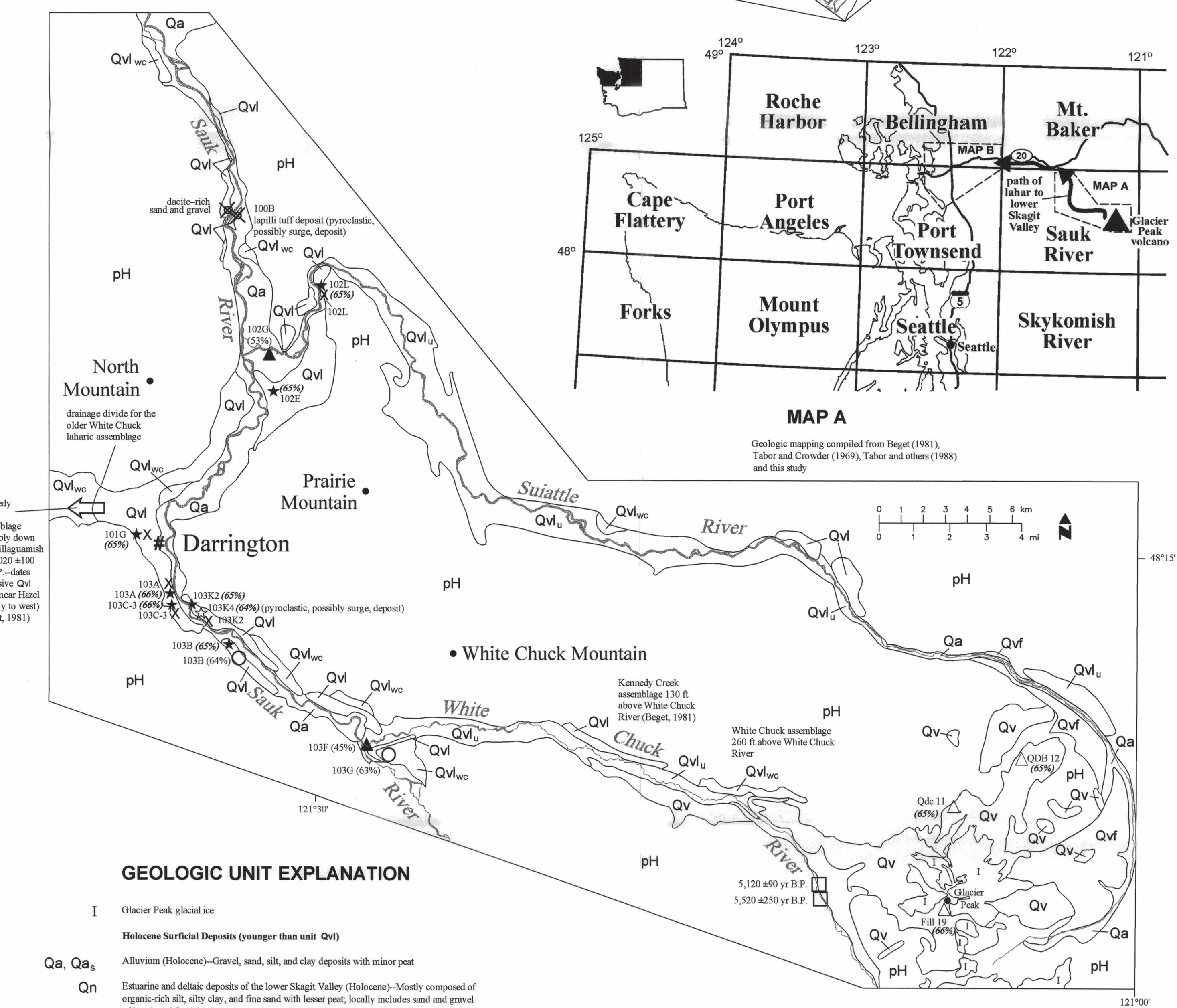
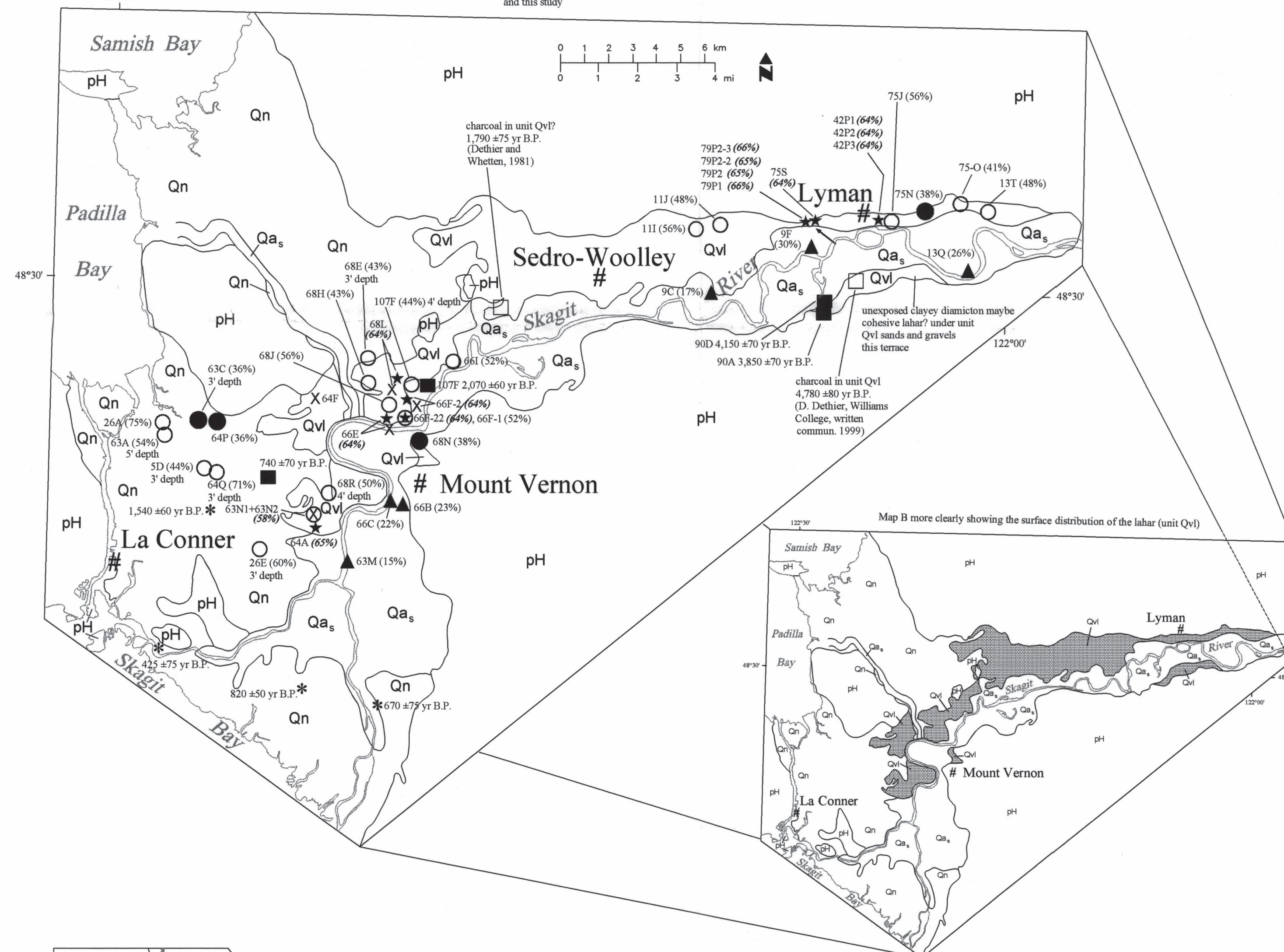
Post-lahar Skagit River channel migration and lateral erosion formed erosional terraces composed of lahar deposits. The mid-to late-Holocene Skagit River has reworked the volcaniclastic sediment of the terraces resulting in a modern Skagit alluvium that has a distinct Glacier Peak volcanogenic provenance.

We examined all the available water well and geotechnical boring logs for the lower Skagit River valley. The above-mentioned paleogeographic topographic maps, produced using Surfer software, compare this subsurface information (433 wells and geotechnical borings; Fig. 3.1A) with our surface mapping and sample analyses. Clastic geochemistry, sand point-count data, and the common occurrence of upward-fining trends from sandy gravels to sands (as indicated in the field and the water well data) suggest a single massive laharcic event. Thus, we assume one catastrophic input of sediment into the lower Skagit River valley as the result of a Glacier Peak eruption at about 5,000 yr BP. (See Glacier Peak Kennedy Creek eruptive assemblage of Beget, 1982.) Another possible volcanic sediment inundation scenario is that the lahar deposit is a composite of the KCA and unnamed 1,750 yr BP events. Perhaps the younger vesicular-dacitic deposit locally veneers the apparently concordant terrace surface of the older event along the lower Skagit River valley. However, the terrace is not as well expressed as the high Lyman quadrangle lahar terraces and may be recorded only by subtle inset terraces farther down the Skagit Valley near Burlington. (See Dragovich and others, 2000.) However, available information (for example, dacite-dacitic geochemical homogeneity) suggests that most or all of the volcanic sediment composing the terraces in the lower Skagit Valley is correlative with the KCA (~5 ka).

Figures 3.1B-E show the paleogeography of the top and bottom of the lahar runout, as well as the thickness of the lahar and post-lahar sediments. The certainty of our correlation of the lahar diminishes away from lahar sample sites. For example, we were able to directly observe or probe to the top of the lahar north of La Conner and obtain sand samples for point-counting (Map B). Thus, we have high confidence in the subsurface correlation in the La Conner area. Confidence is also relatively high along Interstate-5 (where the lahar is exposed or at shallow

depth) as a result of the tight spacing of geotechnical borings and analytical data that support a laterally continuous larval stratum. However, the our confidence in the correlation of sand and gravel bodies below the present Sagittaw River alluvial system (Figs. 3.1C-E) with the larval uncut is only moderate to low. Confidence is significantly diminished toward Padilla Bay, as well as the present Sagittaw delta, where the larva is apparently at depth or nonexistent (Figs. 3.1C-E) and subsurface data are sparse. Some subsurface data suggest the larva is locally preserved below the modern Sagittaw channel fill; thus, the larva may have been incompletely excavated by post-lahar channel migration and incision.

Sea level during the mid-Holocene (~5,100 yr B.P.) was about 5 to 25 ft lower than at present, and the mid-Holocene Puget Sound extended several miles to the east of the present shoreline. We suggest that the mid-Holocene shoreline was approximately near the -20 ft contour on the pre-lahar paleogeographic map. (See the stylized -20 ft contour line on Fig. 3.1F.) We drew this line on the basis of marine shells reported in boring logs from stratum directly under the lahar west of Interstate-5 (Fig. 3.1A) and not reported from any borings east of the -20 ft contour line.



GEOLOGIC UNIT EXPLANATION

- I Glacier Peak glacial ice
- Holocene Surficial Deposits (younger than unit Qv)**
- Qa, Qas Alluvium (Holocene)—Gravel, sand, silt, and clay deposits with minor peat
- Qn Estuarine and deltaic deposits of the lower Skagit Valley (Holocene)—Mostly composed of organic-rich silt, clay, and fine sand with lesser peat; locally includes sand and gravel of beach and fluvial origin
- Holocene to Pleistocene Sedimentary and Volcanic Deposits**
- Qv Kennedy Creek assemblage of Bedrock (1981) (mid-Holocene) and probably correlative Backus Creek and Dundy Creek assemblages of Bedrock (1981)—Distal, nonconhesive, gravely sand and silt lacustrine and lacustrine deposits with proximal, typically thickly bedded proglacialic flow, some collapse, lar, and tepha deposits; proximal deposits locally contain interbedded lacustrine deposits. (We correlate this assemblage with the laruna-nunot deposits of the lower Skagit Valley.)
- Qv_f Volcaniclastic and volcanic fill of the White Chuck and Suiattle River valleys (Holocene and latest Pleistocene)—Suiattle fill consists of a thickly bedded assemblage of lahars, proglacialic flows, air-fall ash, alluvium, and rare lava flows that grades downvalley into lahars; White Chuck fill consists of a thickly bedded assemblage of lahars, proglacialic flows, alluvium, and reworked ash and till that partly grades downvalley into Kennedy Creek and White Chuck assemblage lahars (Tabor and others, 1988); Bedrock (1982) correlated much of the debris in these volcanic aprons with the mid-Holocene eruptive episode of unit Qv
- Qv Undivided volcanic rocks and deposits of Glacier Peak (Holocene to Pleistocene)—Includes valley-bottom, valley-side, ridge-capping, and undivided flows and flow breccias of Tabor and Crowder (1969) and Tabor and others (1988); locally includes proglacialic deposits (for example, White Chuck valley vitric tuff); available geochemical analyses suggest that some of the dacite flows on Glacier Peak are associated with the mid-Holocene unit Qv proglacialic and lacustrine events)
- Qv_{we} White Chuck assemblage of Bedrock (1982) (late Pleistocene)—Intermittent tepha eruptions and large-volume proglacialic flows, lahars, and collapse eruptions
- Qv_{lu} Undivided lahars on Glacier Peak (Holocene to Pleistocene)—Boulder-dimension to well-sorted sand and gravel; includes poorly mapped lahars correlative with units Qv_{we} and Qv as well as other volcanic assemblages of Bedrock (1981); contacts from Tabor and others (1988)
- Pre-Holocene Glacial Deposits and Bedrock**
- pH Glacial deposits and bedrock (pre-Holocene)

SYMBOL EXPLANATION FOR SAMPLE SITES*

Symbols may be combined

- Mounted thin section of lahar sand for point-counting (percent microlitic volcanics in parentheses)
- Mounted thin section of late-stage (transitional) lahar sand for point-counting (percent microlitic volcanics in parentheses) (composition and field setting suggest late-stage lahar runout or reworked lahar)
- Mounted thin section of sand from modern alluvium for point-counting (percent microlitic volcanics in parentheses)
- Geochemical sample of lahar dacite clast (*italicized and hold percent SiO₂ in parentheses*)
- Geochemical sample of angular dacite clast in lapilli tuff lake (synthetic sand deposit?) (*italicized percent SiO₂ in parentheses*) (See Map A; southeast of Derrington)
- Comparative geochemical samples of dacite flow and brecciated bomb of Taber and Crowder (1969) near Glacier Park (*italicized percent SiO₂ in parentheses*)
- Petrographic analysis of lahar dacite-sand sample (some point counted)
- Post-lahar radiocarbon dating sample site (minimum lahar ages from overlying peats or paleosols) (this study)
- Post-lahar radiocarbon dating sample site (minimum lahar ages from overlying peats or paleosols), dates were obtained from human settlements probably near the Skagit River delta front (Thompson 1978)
- Lahar radiocarbon dating sample (charcoal or logs)
- Miscellaneous notes
- Location of the most distal lahar (unit Qv) Skagit Valley stratigraphic section that is correlated by Beget (1981, 1986) with the Kennedy Creek assemblage (Map B) near the city of Lyman