

GEOLOGIC MAP OF WASHINGTON - NORTHEAST QUADRANT

by
**KEITH L. STOFFEL, NANCY L. JOSEPH, STEPHANIE ZURENKO WAGGONER,
CHARLES W. GULICK, MICHAEL A. KOROSSEC, and BONNIE B. BUNNING**



WASHINGTON DIVISION OF GEOLOGY AND EARTH RESOURCES
GEOLOGIC MAP GM-39

1991



WASHINGTON STATE DEPARTMENT OF
Natural Resources

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Photograph on envelope: Northwest-dipping cuesta formed on resistant layers of gneiss in the Okanogan metamorphic core complex near Riverside, Washington; Okanogan River in the foreground. Aerial view to the northeast. Photograph courtesy of K. F. Fox, Jr., U.S. Geological Survey.

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PREFACE

State geologic maps are valuable tools that can be used for exploration for mineral resources, evaluation of geologic hazards, land-use planning, and education. Previous geologic maps of the state of Washington were published by our predecessors, the Washington Division of Geology (Culver, 1936) and the Washington Division of Mines and Geology (Hunting and others, 1961). Both of these maps are now out of print. The scale of the first two state geologic maps was 1:500,000. However, 74 percent of all published and thesis geologic maps of areas in Washington postdate the 1961 state geologic map. We therefore decided to publish the new state geologic map in four quadrants at a scale of 1:250,000 to better display this accumulated information.

As in the first map of this quadrant series (Walsh and others, 1987), we have attempted to make this geologic map useful at several levels of detail and to a variety of users by (1) employing time-lithologic units on the map rather than formations; (2) portraying unconsolidated deposits in as much detail (where available) as bedrock units; (3) using a multiple-level scheme of colors, patterns, and map symbols to represent the geologic units; and (4) providing a wider range of supporting materials than is commonly found on state geologic maps. Full explanation of these characteristics is given in the text that follows.

The state geologic map program has entailed an exhaustive effort to obtain all existing published and unpublished geologic map data, and to synthesize the information into a series of 1:100,000- and 1:250,000-scale geologic maps and reports. This project has enhanced working contacts, data exchange, and mutual support among geologists at the Division of Geology and Earth Resources and at universities, in federal agencies, and in the private sector.

We are proud of this map and look forward to the continuing challenge of compiling the remaining two quadrant maps. We hope that you will realize as many benefits in the use of this map as we have in its preparation.

J. Eric Schuster, Project Manager

Bonnie B. Bunning, Assistant Project Manager

William M. Phillips, Assistant Project Manager

May 22, 1991

GEOLOGIC MAP OF WASHINGTON – NORTHEAST QUADRANT

by

Keith L. Stoffel, Nancy L. Joseph, Stephanie Zurenko Waggoner,
Charles W. Gulick, Michael A. Korosec, and Bonnie B. Bunning

INTRODUCTION

The Geologic Map of Washington – Northeast Quadrant is the second in a series of four 1:250,000-scale geologic maps, which together will make up the third geologic map of Washington published by the Washington Division of Geology and Earth Resources and its predecessors. The Geologic Map of Washington – Southwest Quadrant, the first map in the series, is available as Washington Division of Geology and Earth Resources Geologic Map GM-34 (Walsh and others, 1987). The remaining two maps in the series, the Geologic Map of Washington – Southeast Quadrant and the Geologic Map of Washington – Northwest Quadrant, are currently in preparation (May 1991).

In addition to the geologic map, the northeast quadrant publication includes a **key to geologic units, descriptions of map units, a list of named units, a correlation diagram, a 1:625,000-scale bedrock geologic and tectonic map, a summary of the geologic history of northeastern Washington, source-of-data maps, and a list of references cited.** A topographic base map is available separately.

MAP COMPILATION

Preliminary compilation of the Geologic Map of Washington – Northeast Quadrant began in 1984, when all available published and unpublished geologic information was compiled on sixteen 1:100,000-scale quadrangles. (See Sheet 1.) Between 1984 and 1988, Division of Geology and Earth Resources (DGER) geologists mapped (at reconnaissance and detailed levels) areas where previous geologic mapping was either inadequate or lacking. New geologic mapping was also acquired during that time through a DGER geologic mapping support program, which funded 23 university graduate student and faculty member mapping projects. (See Acknowledgments.) The new geologic mapping has been augmented by 23 new potassium-argon (K-Ar) ages, 2 new radiocarbon (¹⁴C) ages, and 230 new whole-rock geochemical analyses, which are provided in the reports listed in Table 1.

In 1990, following extensive reviews by geologists outside the DGER, fourteen of the sixteen 1:100,000-scale geologic maps were released as Washington Division of Geology and Earth Resources open-file reports (Table 1). Reports were not prepared for the Chelan and Wenatchee quadrangles because geologic maps of those quadrangles have been published by the U.S. Geological Survey (USGS) (Tabor and others, 1982, 1987). The sixteen 1:100,000-scale geologic

maps and accompanying source-of-data maps, descriptions of map units, isotopic and fossil age data, and whole-rock geochemistry data were then synthesized to prepare the 1:250,000-scale geologic map of Washington and the other components of this publication.

ACKNOWLEDGMENTS

The Geologic Map of Washington – Northeast Quadrant is the culmination of the efforts of hundreds of geologists over the better part of a century. The contributions of many of the geologists are acknowledged by citation. (See **References Cited.**) Regrettably, many other significant works are not cited. We gratefully acknowledge all those geologists who have made vital contributions toward deciphering the complex geology of northeastern Washington.

Numerous individuals have contributed greatly to the success of the state geologic map project and the publication of this report. We offer warm thanks to the people named on the following pages for their invaluable contributions.

Table 1. Compiler, 1:100,000-scale quadrangle, and DGER open-file report number for geologic maps in the northeast quadrant of Washington

Compilers	Quadrangle	Open File Report Number
Bunning, B. B.	Twisp	90-9
Gulick, C. W.	Moses Lake	90-1
Gulick, C. W.	Ritzville	90-2
Gulick, C. W.; Korosec, M. A.	Banks Lake	90-6
Gulick, C. W.; Korosec, M. A.	Omak	90-12
Joseph, N. L.	Colville	90-13
Joseph, N. L.	Nespelem	90-16
Joseph, N. L.	Spokane	90-17
Korosec, M. A.	Chelan	—
Stoffel, K. L.	Republic	90-10
Stoffel, K. L.	Oroville	90-11
Stoffel, K. L.; McGroder, M. F.	Robinson Mtn.	90-5
Waggoner, S. Z.	Rosalia	90-7
Waggoner, S. Z.	Chewelah	90-14
Waggoner, S. Z.	Coulee Dam	90-15
Walsh, T. J.	Wenatchee	—

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 Willard N. Hopkins (San Jose State University)
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 R. Wade Holder (Washington State University)
 Eugene P. Kiver (Eastern Washington University)
 Charles M. Knaack (Washington State University)
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 W. Michael Wade (San Jose State University)
 Laureen Wagoner (Washington State University)
 Patricia J. White (University of Washington)
 Glenn R. Wiedenhoft (Washington State University)

Unpublished Geologic Mapping

Brian F. Atwater (USGS)
 Eric R. Braun (consulting geologist)
 Eric S. Cheney (University of Washington)
 Edward D. Fields (Boise Cascade Corp.)
 Kenneth F. Fox, Jr. (USGS)
 John W. Goodge (formerly University of Montana)
 Vicki L. Hansen (formerly University of Montana)
 Larry G. Hanson (formerly University of Washington)
 Eugene P. Kiver (Eastern Washington University)
 Michael F. McGroder (Exxon Production Research Co.)
 Fred K. Miller (USGS)
 C. Dean Rinehart (USGS)
 J. Eric Schuster (DGER)
 James R. Snook (Eastern Washington University)
 Dale F. Stradling (Eastern Washington University)
 Gerald W. Thorsen (DGER, retired)
 Victoria R. Todd (USGS)
 Joseph R. Wilson (Law Environmental Services; formerly USGS)
 Robert G. Yates (USGS, retired)

Unpublished Isotopic Age Data

Brian F. Atwater (USGS)
 Robert J. Fleck (USGS)
 Eugene P. Kiver (Eastern Washington University)
 Fred K. Miller (USGS)
 Robert B. Miller (San Jose State University)
 C. Dean Rinehart (USGS)
 Robert C. Roback (University of Texas at Austin)
 Victoria R. Todd (USGS)

Fossil Identification and/or Data

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 J. Thomas Dutro, Jr. (USGS)
 Anita G. Harris (USGS)
 Philip R. Jackson (Teck Resources Inc.)
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Discussions of Regional Geology

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Review of 1:100,000-scale Open-File Reports

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Review of 1:250,000-scale Geologic Map and Report

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 Rowland W. Tabor (USGS)

MAP DESIGN

Base Map

No appropriate topographic base map of northeastern Washington existed when the state geologic map project was initiated. Therefore, the Division chose to construct a new 1:250,000-scale base map by simplifying 1:100,000-scale, metric-contour topographic maps of northeastern Washington, which were published recently by the USGS or are in preparation by that agency. The 1:250,000-scale base map was manually scribed from scale-stable reductions of 1:100,000-scale map separates. The control framework was a computer-generated latitude-longitude grid plotted on a Universal Transverse Mercator Projection by the USGS. The 1:250,000-scale base map is available as a separate publication (Washington Division of Geology and Earth Resources Topographic Map TM-2).

Geologic Map

The Geologic Map of Washington - Northeast Quadrant is chiefly an age-lithology map, not a geological formation map. Formations are shown only for Miocene volcanic rocks and for Precambrian metasedimentary and metavolcanic rocks. These exceptions have been made because without this subdivision into formations, the geologic map would be too simplistic to be of much benefit to the map user.

A multiple-level scheme of colors, patterns, and map symbols has been used to portray age and lithology of geologic units on the 1:250,000-scale geologic map. (See **Key to Geologic Units** on Sheet 1.) The first level of detail, which is expressed by broad color ranges, distinguishes six general lithologic subdivisions: unconsolidated sediments, sedimentary and volcanic rocks, metasedimentary and metavolcanic rocks, metamorphic rocks, intrusive rocks, and mixed metamorphic and intrusive rocks. The second level of detail, which is expressed by variations of color within each broad color range, indicates age. The third level of detail, represented by patterns, distinguishes lithologic units or small groups of lithologic units. The fourth level of detail, represented by the map symbols, identifies individual geologic map units.

Standard USGS age symbols are used for most map units, but because of the prevalence of Tertiary rocks in Washington, each Tertiary epoch has been assigned a separate age symbol. The Tertiary symbols were chosen with the help of the editorial staff of the USGS (Denver), but the USGS has not formally adopted these symbols. Lithologic symbols used here were devised specifically for the DGER state geologic map project, and are not standard USGS symbols. Where formations are shown on the 1:250,000-scale geologic map, the map symbol includes a subscript letter or number to distinguish it from other units of similar age and lithology — for example, the symbol for the Miocene Saddle Mountains Basalt is **Mv_s**.

Because the 1:250,000-scale map units are age-lithologic units, formations consisting of diverse rock types are separated into their component lithologies and consequently are included in more than one map unit. To determine all map units by which a named unit (formation or group) is shown, consult the **List of Named Units** (Sheet 2).

DESCRIPTIONS OF MAP UNITS

A lithologic description of each map unit on the 1:250,000-scale geologic map is given in the **Descriptions of Map Units** (Sheet 2). Quaternary sediments are described first, followed by sedimentary and volcanic rocks, metasedimentary and metavolcanic rocks, intrusive rocks, mixed metamorphic and igneous rocks, and metamorphic rocks. Within each lithologic group, map units are addressed in order of increasing age. Information concerning the age and stratigraphic relations of the map units is given in the **Correlation Diagram** (Sheet 3).

At the end of each unit description is a list of named units that make up the geologic map unit. Formally named formations that are included in lexicons published by the USGS are listed without citation and with the word "Formation" or lithologic term capitalized. Informally named units are followed by the citation to the work that defines the unit; for these units, the word "formation" or lithologic term is not capitalized. Some map units consist of a single formation or part of a formation; others are a combination of named and unnamed units; still others consist only of unnamed age-lithologic units. If the map unit consists entirely of named units, the words "consists of" are applied to the list of names given. If the map unit contains both named and unnamed units, the word "includes" is used.

Sedimentary and Volcanic Rocks

Sedimentary and volcanic rocks include sub-greenschist-facies low-grade deposits only. Classification of sandstones follows the terminology of Dickinson (1970). Assignment of volcanic rock names was made on the basis of whole-rock geochemistry and the total alkali-silica (TAS) diagram (LeMaitre, 1984; Zanettin, 1984).

Metasedimentary and Metavolcanic Rocks

Metasedimentary and metavolcanic rocks include sedimentary and volcanic rocks that have been metamorphosed to the greenschist facies. Sedimentary and volcanic names are preceded by the prefix "meta" in order to readily distinguish greenschist-facies metasedimentary and metavolcanic rocks from sub-greenschist-facies sedimentary and volcanic rocks and from amphibolite-facies high-grade metamorphic rocks derived from sedimentary and volcanic rock protoliths. For example, the term "metallimestone" is applied to recrystallized greenschist-facies limestone to distinguish it from recrystallized amphibolite-facies marble and from limestone that is not recrystallized.

Intrusive Rocks

Intrusive rock names in the Descriptions of Map Units follow the International Union of Geological Sciences (IUGS) modal classification (Streckeisen, 1973). Geologic map units that consist of a single rock type or are dominated by one rock type are given lithologic names such as "granite" or "diorite". Geologic map units that contain a mixture of rock types are assigned to one of the following generic categories: (1) *acidic intrusive rocks*, which include the IUGS alkali feldspar granite, granodiorite, tonalite, and quartz monzonite fields; (2) *intermediate intrusive rocks*, which include the IUGS monzonite, quartz monzodiorite, monzodiorite, quartz diorite, and diorite fields; (3) *basic intrusive rocks*, which include the IUGS gabbro, quartz gabbro, monzogabbro, and

quartz monzogabbro fields; (4) *syenitic intrusive rocks*, which include the IUGS syenite, quartz syenite, alkali-feldspar syenite, and alkali-feldspar quartz syenite fields; and (5) *alkalic intrusive rocks*, which include the IUGS foid-bearing alkali-feldspar syenite, foid-bearing syenite, foid-bearing monzonite, foid syenite, and foid monzosyenite fields. Alkalic intrusive rock names (shonkinite, malignite, and foyaite) were assigned using the modal classification of Miller (1972).

Mixed Metamorphic and Igneous Rocks

Mixed metamorphic and igneous rocks are heterogeneous, migmatitic rocks that formed by deep burial, ultra-metamorphism, and partial fusion (anatexis) of layered sedimentary and volcanic rocks and/or by injection of magma into layered sedimentary and volcanic rocks. This category of rock types includes migmatite and agmatite, which are migmatitic at an outcrop scale, and mixed metamorphic and igneous rocks, which are migmatitic at a 1:24,000 quadrangle scale.

Metamorphic Rocks

Metamorphic rocks include sedimentary and volcanic rocks that have been metamorphosed to the amphibolite facies or higher (layered metamorphic rocks), and moderately to strongly foliated intrusive rocks (orthogneiss) that have been thermally or dynamically metamorphosed.

LIST OF NAMED UNITS

The **List of Named Units** (Sheet 2) is a summary of named geologic units that appear in published literature. The list includes *all* formal and *selected* informal unit names that are currently in use, as well as a few names that were once widely used, but are now considered archaic.

Some of the named geologic units are represented on the 1:250,000-scale geologic map (Sheet 1) by a single symbol (age-lithologic unit), whereas other named units are represented by two or more symbols. References cited for each named unit consist of the citation(s) in which the unit was defined, redefined, or extensively reviewed. Locations given for each named unit are intended to guide the map user to the general area on the 1:250,000-scale map in which the named geologic unit occurs.

SOURCES OF MAP DATA

Figures 1 through 6 constitute a series of index maps that illustrate the areas covered by the sources of geologic mapping data used to compile the 1:250,000-scale geologic map. On each of these index maps, numbered areas correspond to abbreviated citations in the explanations. Complete citations are given in the **References Cited**. Patterned areas on Figure 1 indicate areas where DGER geologists performed original geologic mapping for the state geologic map project. References that supplied other geologic data (such as isotopic and fossil ages or geochemistry) are not indicated on these figures, but are listed in the **References Cited**. Plate or sheet numbers are specified only for reports that contain more than one of these elements.

Some unpublished maps and data, cited as written or oral communications in this report, were used to compile the 1:250,000-scale geologic map. They are available for inspection at the Division's Olympia office.

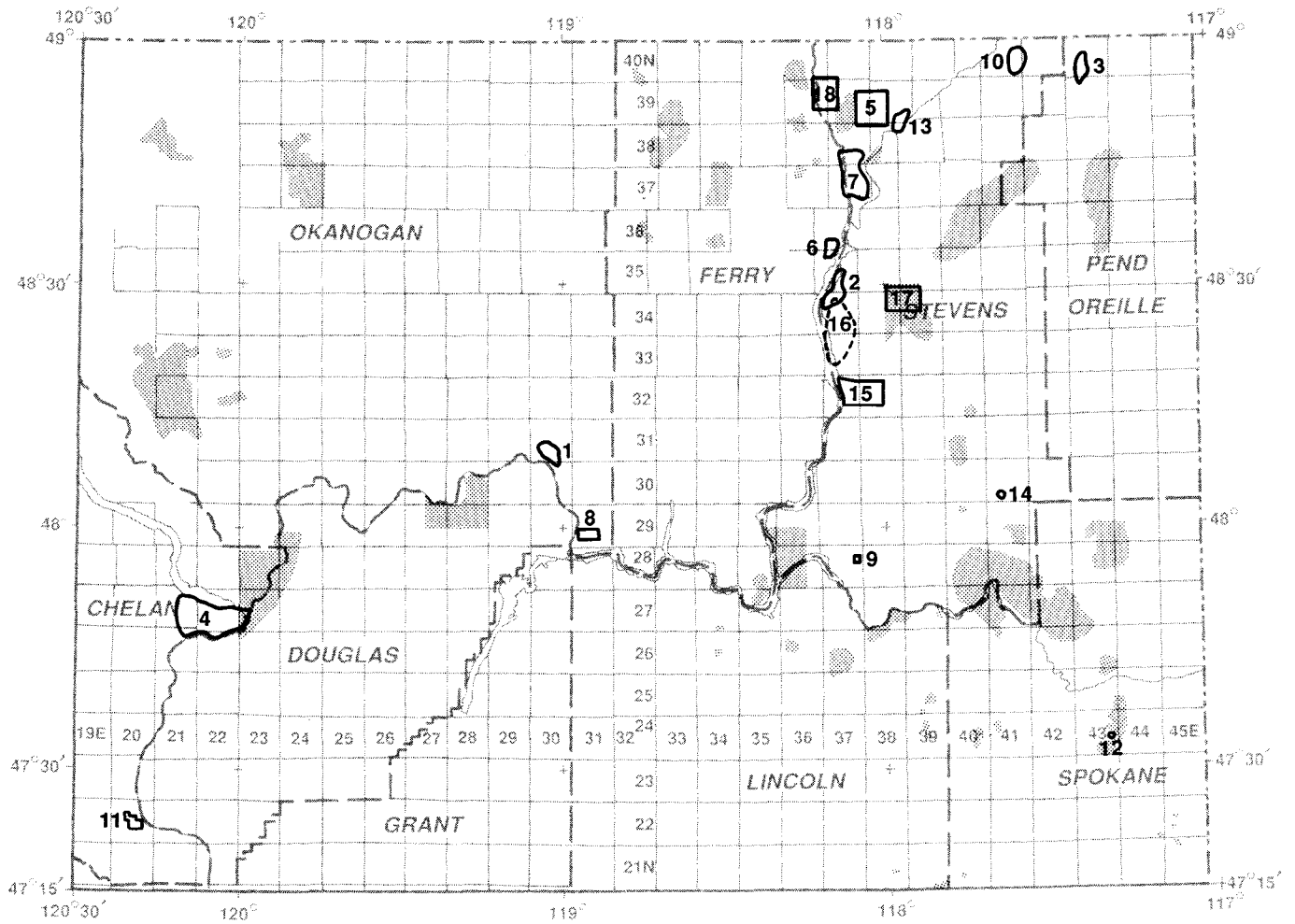


Figure 1. Index to sources of map data, scales 1:1,200 through 1:21,600. Shaded areas, original mapping by DGER geologists.

EXPLANATION

- | | |
|--|---|
| 1. Broch, 1979; plate 3.1, scale 1:12,000 | 10. O'Keefe, 1980; plate 1, scale 1:5,700 |
| 2. Cole, 1960; plate 1, scale 1:21,480 | 11. Ott, 1988; plate 1, scale 1:12,000 |
| 3. Greenman, 1976; plate 1, scale 1:12,000; plate 2, scale 1:2,570 | 12. Page, 1942; plate 32, scale 1:2,400 |
| 4. Hopson, 1955; plate 2, scale 1:19,500 | 13. Phillips, 1979; plate, 1:6,000 |
| 5. Hyde, 1985; geologic map, scale 1:12,000 | 14. Rigg, 1958; figure 224, scale 1:21,600 |
| 6. Jones and others, 1961; plate 1, scale 1:9,600 | 15. Smith, 1982; plate 1, scale 1:14,900 |
| 7. Kuenzi, 1961; plate 1, scale 1:16,800 | 16. Smith, M. T., DGER unpub. mapping, 1990; scale 1:12,000 |
| 8. Milliken, 1981; figure 4, scale 1:12,000 | 17. Snook, 1981; plate 1, scale 1:6,000 |
| 9. Nash, 1978; sheet 2, scale 1:1,200 | 18. West, J. R., 1976; figure 1, scale 1:12,000 |

6 GEOLOGIC MAP OF WASHINGTON — NORTHEAST QUADRANT

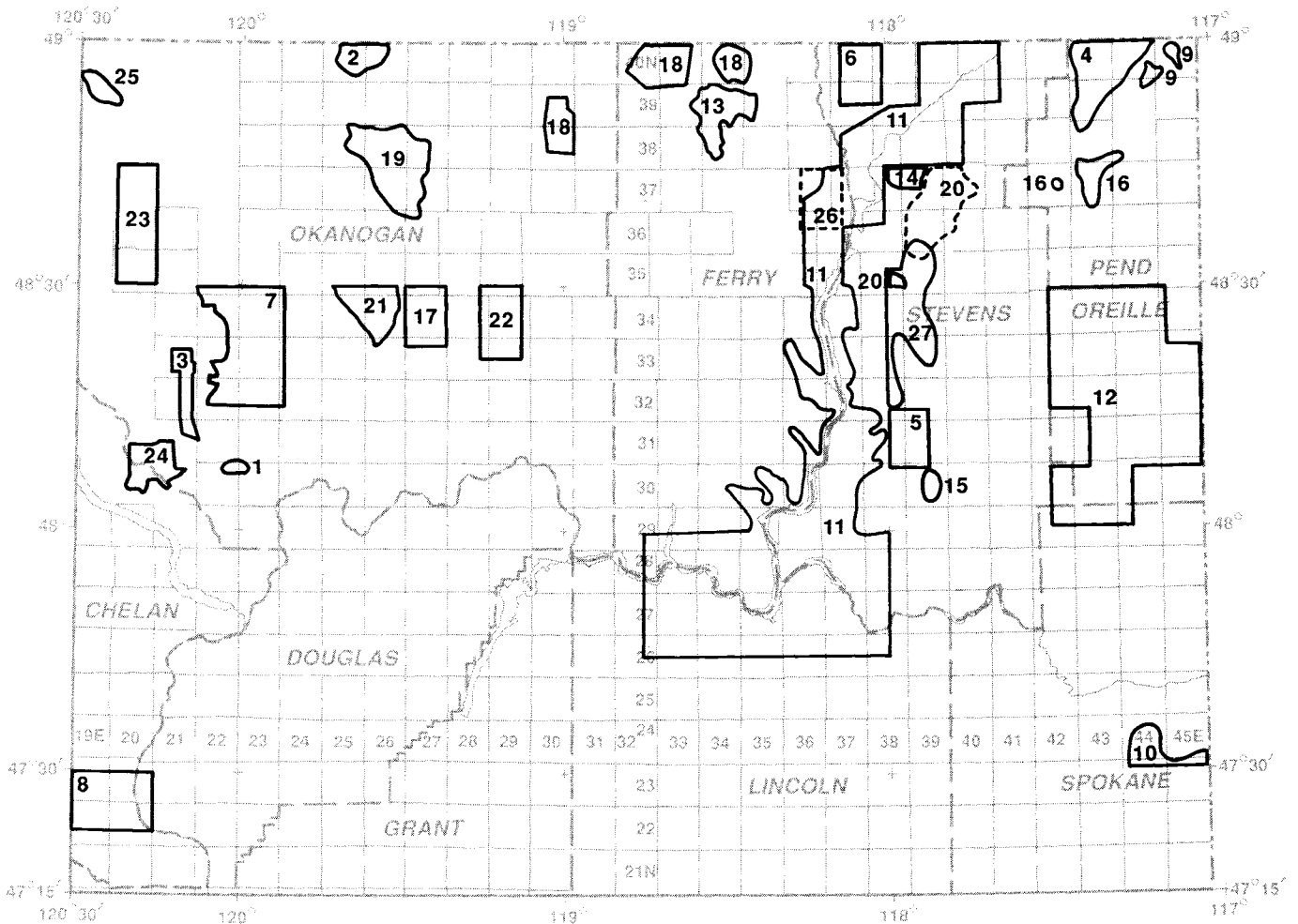


Figure 2a. Index to sources of map data, scale 1:24,000.

EXPLANATION

- | | |
|---|---|
| 1. Buddington, 1986; plate | 14. Laskowski, 1982; figure 2A |
| 2. Buddington, A. M., Western Washington Univ., written commun., 1987 | 15. Miller, F. K., USGS, written commun., 1988 |
| 3. DiLeonardo, 1987; plate 1 | 16. Miller, F. K., USGS, written commun., 1988 |
| 4. Dings and Whitebread, 1965; plate 1 | 17. Minard, 1985; sheet |
| 5. Evans, 1987; plate 1 | 18. Orr, 1985; maps A, B, and C |
| 6. Fox, 1981; sheet 1 | 19. Pine, 1985; plate 1 |
| 7. Frey, A. M., DGER unpub. mapping, 1988 | 20. Schuster, J. E., DGER, written commun., 1988 |
| 8. Gresens, 1983; plates 1 and 2 | 21. Sims, 1984; plate |
| 9. Groffman, 1986; plate | 22. Singer, 1984; plate 1 |
| 10. Hosterman, 1969; plate | 23. Todd, V. R., USGS, written commun., 1989 |
| 11. Kiver, E. P.; Stradling, D. F., Eastern Washington Univ., written commun., 1987 | 24. Wade, 1988; plate |
| 12. Kiver, E. P.; Stradling, D. F., DGER unpub. mapping, 1988 | 25. White, 1986; figure 3 |
| 13. Knaack, C. M., DGER unpub. mapping, 1987 | 26. Wilson, 1980; sheet 1 |
| | 27. Yates, R. G., USGS retired, written commun., 1986 |

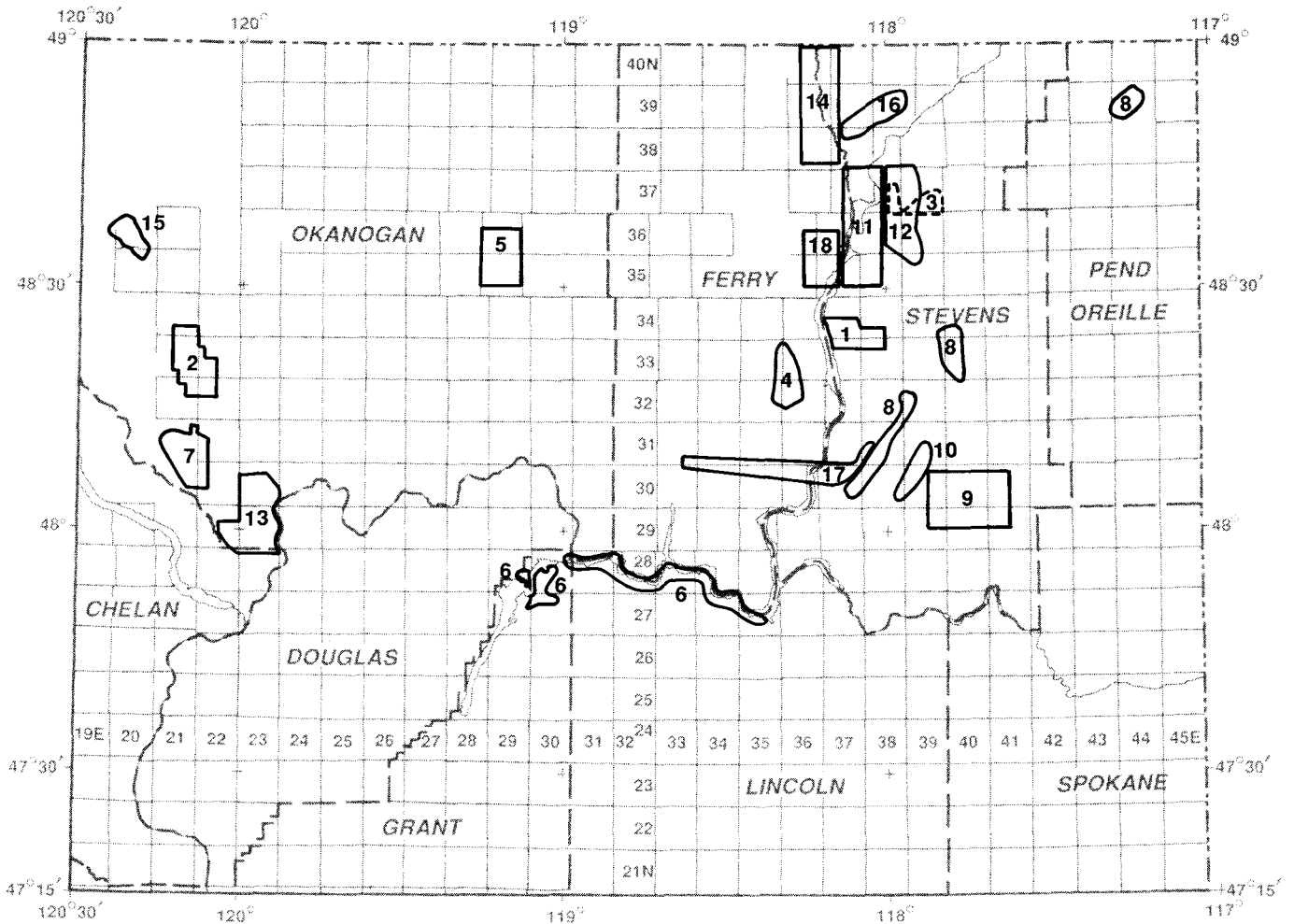


Figure 2b. Index to sources of map data, scale 1:24,000 (continued).

EXPLANATION

- | | |
|--|---|
| <ul style="list-style-type: none"> 1. Abrams, 1980; plate 1 2. Burnet, 1976; figure 21. 3. Duncan, 1982; plates 1 and 24. 4. Fullmer, 1986; plate 7. 5. Gulick, 1987; plate 1 6. Holder, R. W.; Holder, G. A. M., DGER unpub. mapping, 1987 7. Hopkins, 1987; plate 111. 8. Lindsey, 1988; plates 1, 2, 4, and 5 9. McLucas, 1980; plates 1 and 2 | <ul style="list-style-type: none"> 10. Miller, F. K., USGS, written commun., 1989 11. Mills, 1985; plates 1 and 2 12. Mills and others, 1985; sheets 1 and 2 13. Raviola, 1988; plate 14. Rhodes, 1980; plates 1 and 2 15. Riedell, 1979a; plate 1 16. Roback, R. C., Univ. of Texas at Austin, written commun., 1990 17. Smith, M. T., DGER unpub. mapping, 1989 18. Wilson, 1981a; sheet 1 |
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8 GEOLOGIC MAP OF WASHINGTON — NORTHEAST QUADRANT

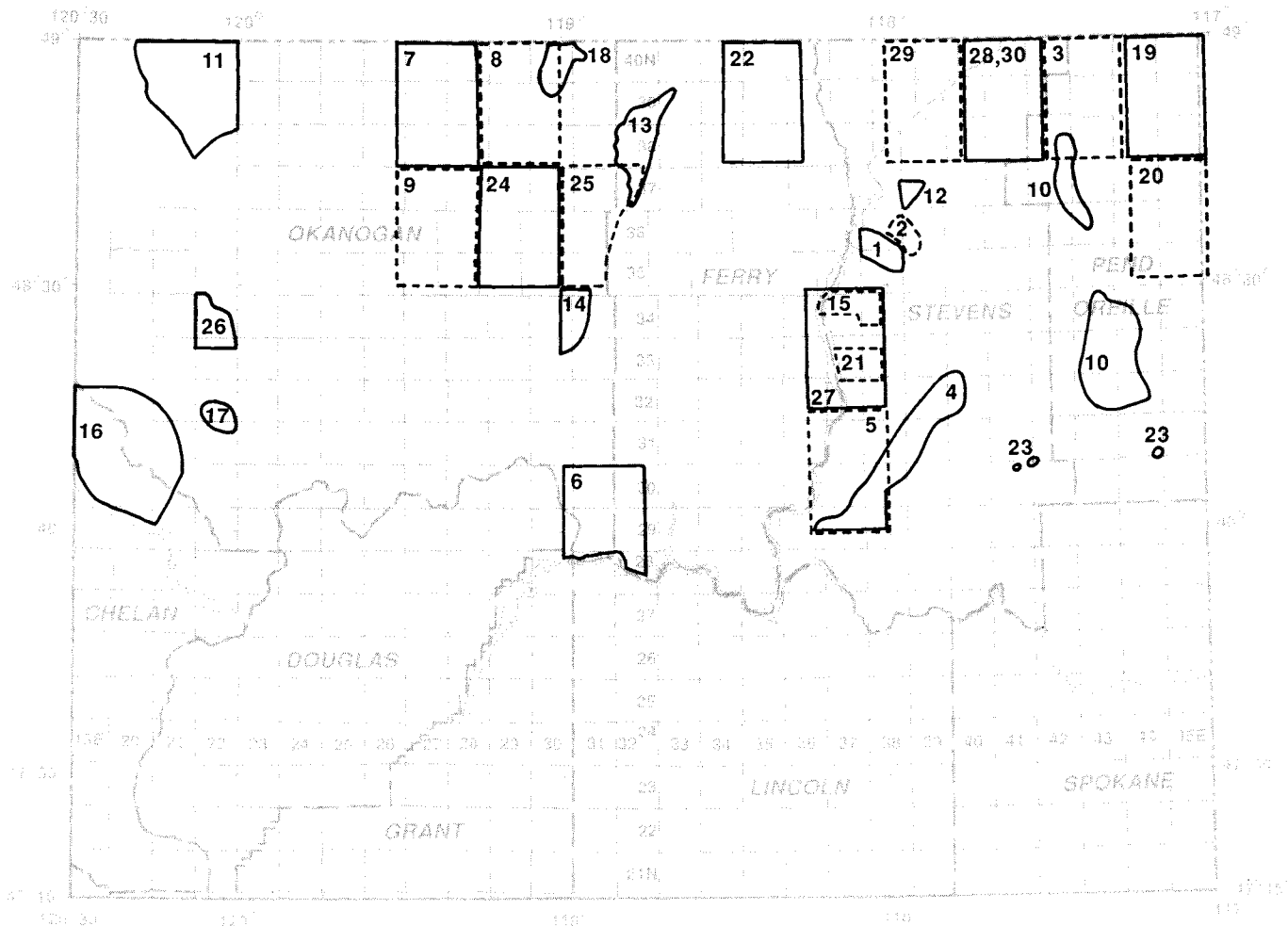


Figure 3. Index to sources of map data, scales 1:25,000 through 1:50,688.

EXPLANATION

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| <ul style="list-style-type: none"> 1. Bradshaw, 1964; plate 25, scale 1:27,500 2. Brainard, 1982; figure 2, scale 1:39,600 3. Burmester and Miller, 1983; plate, scale 1:48,000 4. Campbell and Loofbourow, 1962; plate 1, scale 1:36,000 5. Campbell and Raup, 1964; sheet, 1:48,000 6. Carlson, 1984; plate, scale 1:48,000 7. Fox, 1970; sheet, scale 1:48,000 8. Fox, 1978; sheet, scale 1:48,000 9. Fox, K. F., Jr.; Rinehart, C. D., USGS, written commun., 1988, scale 1:48,000 10. Gager, 1982; maps B and C, scale 1:33,790 11. Hawkins, 1963; plate 1, scale 1:31,680 12. Hogge, 1982; figure 2, scale 1:36,000 13. Holder, 1990; plates 1A and 1B, scale 1:50,690 14. Holder, R. W.; Holder, G. A. M., Georgia Southern College, written commun., 1989, scale 1:48,270 15. Janzen, 1981; plate, scale 1:25,000 | <ul style="list-style-type: none"> 16. Libby, 1964; plate 27, scale 1:31,680 17. McKee, Bates, Univ. of Washington, written commun., 1968, scale 1:32,000 18. McMillen, 1979; figure 7, scale 1:31,680 19. Miller, 1983; plate, scale 1:48,000 20. Miller, F. K., USGS, written commun., 1988, scale 1:48,000 21. Orlean, 1981; plate [geology], scale 1:25,000 22. Pearson, 1977; sheet, scale 1:48,000 23. Rigg, 1958; figures 125 and 221, scale 1:42,670 24. Rinehart, C. D.; Fox, K. F., Jr., USGS, written commun., 1988, scale 1:48,000 25. Rinehart and Greene, 1988; plate, scale 1:48,000 26. Ryason, 1959; fig. 11, scale 1:29,370 27. Snook and others, 1990; sheet, scale 1:48,000 28. Yates, 1964; sheet, scale 1:31,680 29. Yates, 1971; sheet, scale 1:31,680 30. Yates, 1976; plate 1, scale 1:31,680 |
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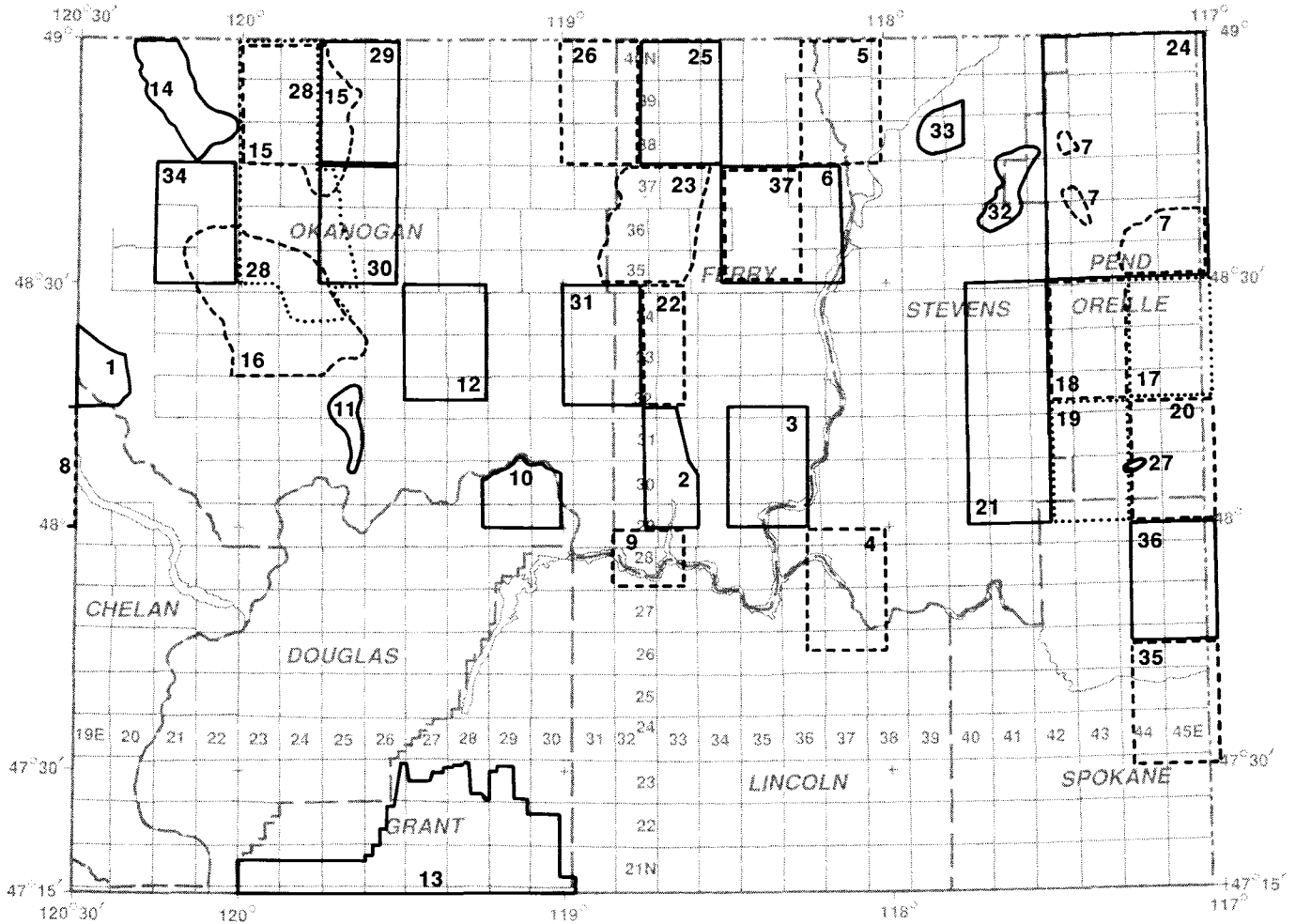


Figure 4. Index to sources of map data, scales 1:62,500 through 1:96,000.

EXPLANATION

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| <ol style="list-style-type: none"> 1. Adams, 1961; plate 2, scale 1:63,360 2. Atwater, 1986; plate 2, scale 1:62,500 3. Becraft, 1966; sheet, scale 1:62,500 4. Becraft and Weis, 1963; plate 1, scale 1:62,500 5. Bowman, 1950; plate 1, scale 1:63,360 6. Campbell and Thorsen, 1975; sheet, scale 1:62,500 7. Castor and others, 1982; plate 22, scale 1:62,500 8. Cater and Wright, 1967; sheet, scale 1:62,500 9. Fleshman and Dodd, 1982; plate 9, scale 1:67,580 10. Fox, K. F., Jr., USGS, written commun., 1988, scale 1:62,500 11. Fritz, 1978; plate, scale 1:62,500 12. Goodge, J. W.; Hansen, V. L., formerly USGS, written commun., 1988, scale 1:62,500 | <ol style="list-style-type: none"> 13. Grolier and Bingham, 1971; sheets 1 and 2, scale 1:62,500 14. Hawkins, 1968; plate 1, scale 1:63,360 15. Hibbard, 1971; figure 3, scale 1:62,500 16. Menzer, 1982; plate 1, scale 1:63,360 17. Miller, 1974a; sheet, scale 1:62,500 18. Miller, 1974b; sheet, scale 1:62,500 19. Miller, 1974c; sheet, scale 1:62,500 20. Miller, 1974d; sheet, scale 1:62,500 21. Miller and Clark, 1975; plates 1 and 2, scale 1:62,500 22. Moye, 1984; plate 1, scale 1:62,500 23. Muessig, 1967; plate 1, scale 1:62,500 24. Park and Cannon, 1943; plate 1, scale 1:96,000 | <ol style="list-style-type: none"> 25. Parker and Calkins, 1964; plate 1, scale 1:62,500 26. Pearson, 1967; sheet, scale 1:62,500 27. Rigg, 1958; figure 124, scale 1:78,000 28. Rinehart, 1981; plate 1, scale 1:96,000 29. Rinehart and Fox, 1972; plate 1, scale 1:62,500 30. Rinehart and Fox, 1976; plate 1, scale 1:62,500 31. Staatz, 1964; plate 1, scale 1:62,500 32. Thorsen, 1966; figure 11, scale 1:63,360 33. Todd, 1973; figure 2, scale 1:84,480 34. Todd, V. R., USGS, written commun., 1988, scale 1:62,500 35. Weis, 1968; sheet, scale 1:62,500 36. Weissenborn and Weis, 1976; sheet, scale 1:62,500 37. Wilson, J. R.; formerly USGS, written commun., 1987, scale 1:62,500 |
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10 GEOLOGIC MAP OF WASHINGTON — NORTHEAST QUADRANT

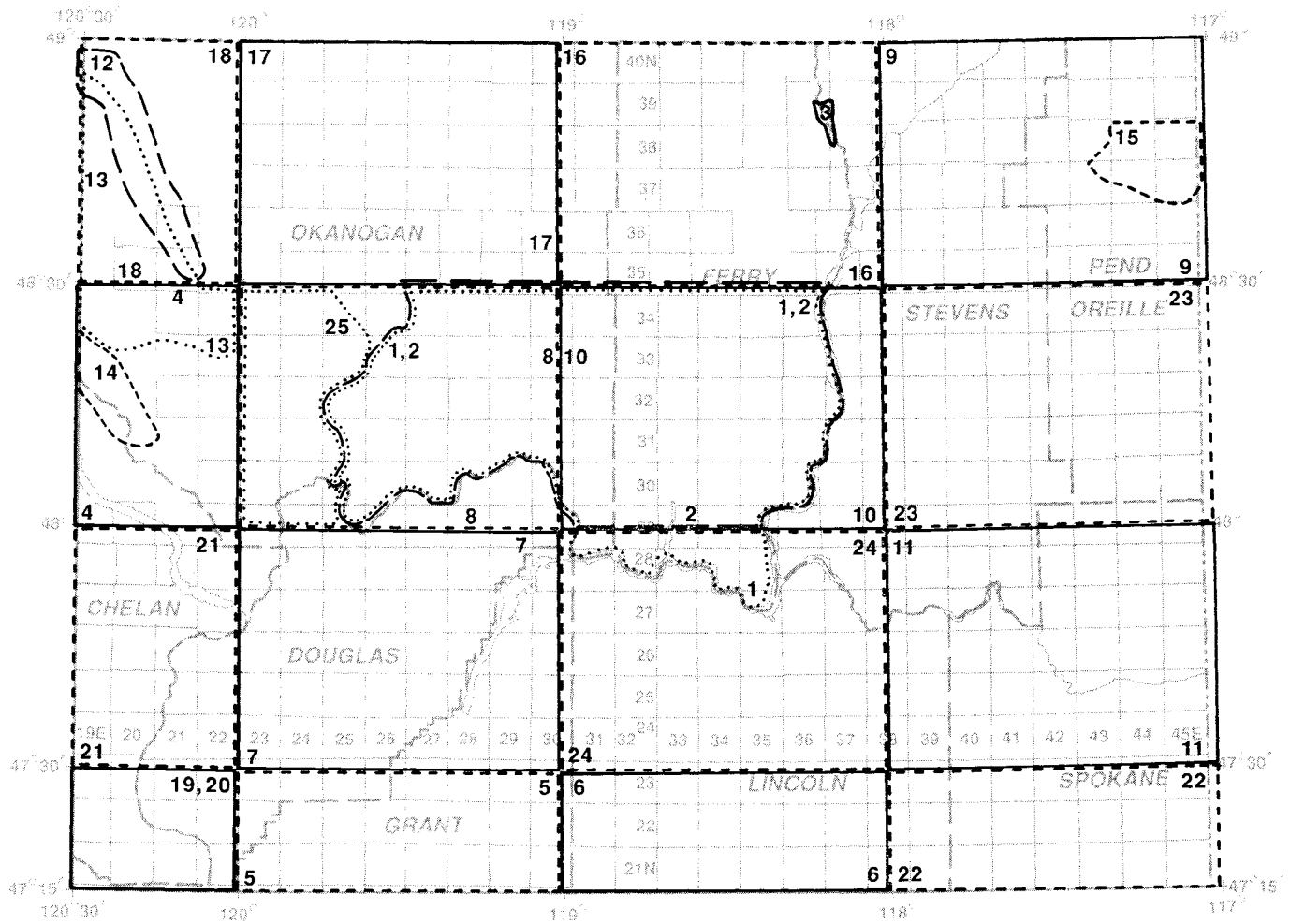


Figure 5. Index to sources of map data, scales 1:100,000 through 1:101,400.

EXPLANATION

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| 1. Atwater and Rinehart, 1984; plates 1, 2, and 3, scale 1:100,000 | 13. McGroder and others, 3 plates, 1990 |
| 2. Atwater, B. F.; Rinehart, C. D., USGS, written commun., 1987, scale 1:100,000 | 14. Miller, 1987; plate 1, scale 1:100,000 |
| 3. Braun, E. R., consultant; Fields, E. D., Boise Cascade Corp., written commun., 1989, scale 1:100,000 | 15. Miller and Frisken, 1984; figure 2, scale 1:101,380 |
| 4. Bunning, 1990; plate, scale 1:100,000 | 16. Stoffel, 1990a; plate, scale 1:100,000 |
| 5. Gulick, 1990a; plate, scale 1:100,000 | 17. Stoffel, 1990b; plate, scale 1:100,000 |
| 6. Gulick, 1990b; plate, scale 1:100,000 | 18. Stoffel and McGroder, 1990; plate, scale 1:100,000 |
| 7. Gulick and Korosec, 1990a; plate, scale 1:100,000 | 19. Tabor and others, 1977; sheet, scale 1:100,000 |
| 8. Gulick and Korosec, 1990b; plate, scale 1:100,000 | 20. Tabor and others, 1982; plate, scale 1:100,000 |
| 9. Joseph, 1990a; plate, scale 1:100,000 | 21. Tabor and others, 1987; plate, scale 1:100,000 |
| 10. Joseph, 1990b; plate, scale 1:100,000 | 22. Waggoner, 1990a; plate, scale 1:100,000 |
| 11. Joseph, 1990c; plate, scale 1:100,000 | 23. Waggoner, 1990b; plate, scale 1:100,000 |
| 12. Lawrence, 1968; plate 1, scale 1:101,380 | 24. Waggoner, 1990c; plate, scale 1:100,000 |
| | 25. Wilson, J. R., formerly USGS, written commun., 1987, scale 1:100,000 |

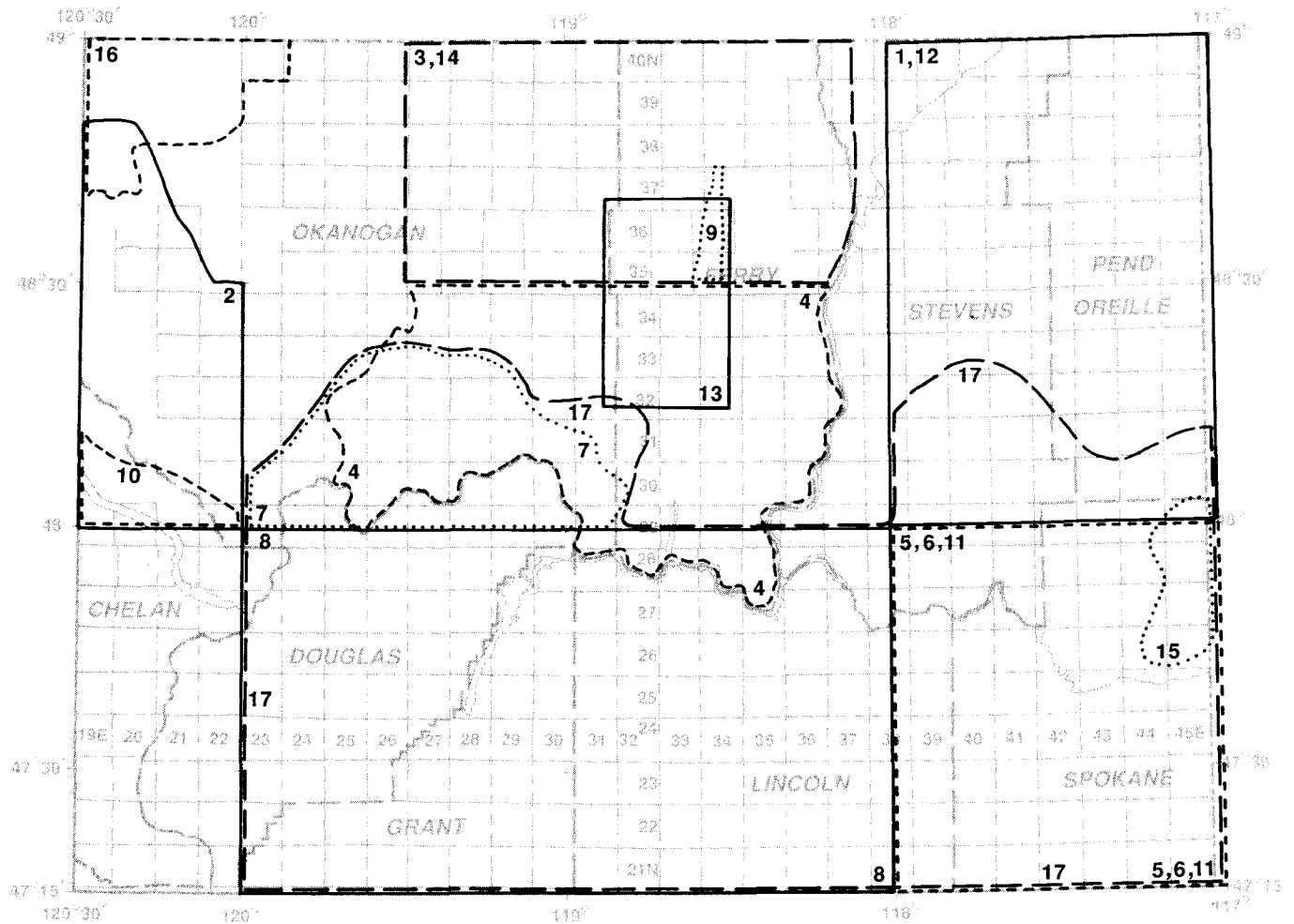


Figure 6. Index to sources of map data, scales 1:125,000 through 1:475,200.

EXPLANATION

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| <ul style="list-style-type: none"> 1. Aadland and others, 1979; sheet, scale 1:250,000 2. Barksdale, 1975; plate 1, scale 1:125,000 3. Cheney and others, 1982; figure 2, scale 1:289,650 4. Colville Confederated Tribes Geology Department, 1984; plate 1, scale 1:250,000 5. Griggs, 1966; sheet, scale 1:125,000 6. Griggs, 1973; sheet, scale 1:250,000 7. Hanson, 1979; sheet, scale 1:250,000 8. Hanson and others, 1979; sheet, scale 1:250,000 9. Holder, 1985; figure 4, scale 1:185,000 | <ul style="list-style-type: none"> 10. Hopson, 1955; plate 1, scale 1:389,900 11. Kiver and others, 1979; sheet, scale 1:250,000 12. Miller and Yates, 1976; sheet 1, scale 1:125,000 13. Moye, 1987; figure 2, scale 1:405,530 14. Orr and Cheney, 1987; figure 2, scale 1:475,200 15. Rhodes and others, 1989; figure 2, scale 1:233,950 16. Staatz and others, 1971; plate 1, scale 1:200,000 17. Swanson and others, 1979a; sheets 1, 2, 4, and 5, scale 1:250,000 |
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CORRELATION DIAGRAM

The age range for each age-lithologic unit on the 1:250,000-scale geologic map (Sheet 1) is represented by a colored box on the **Correlation Diagram** on Sheet 3. The color, pattern, and unit symbol in each box are the same as those for the unit's polygons on the map. Unit symbols on the diagram can be cross-referenced with the **Descriptions of Map Units** and the **List of Named Units**.

Solid lines on the borders of the boxes indicate that the age range of the unit is well constrained. Dashed lines along the border of the boxes indicate that the age is poorly constrained. Queries at the top and/or bottom of a box indicate that the upper or lower age limits of the unit are uncertain. Dashed lines connecting boxes indicate that the map unit includes rocks of more than one age.

Information about the age of geologic units on the diagram is summarized in the text below. More detailed information is given in the series of DGER open-file reports for this quadrant (Table 1). Formal and informal unit names are printed in bold lettering in order to facilitate cross-referencing of the **Descriptions of Map Units** with the **List of Named Units**.

We have used the geologic time scale devised for the "Correlation of Stratigraphic Units of North America (COSUNA)" project of the American Association of Petroleum Geologists (Salvador, 1985) as the basis for this Correlation Diagram. Exceptions are (1) the Eocene-Oligocene boundary is placed at 35.7 Ma (Montanari and others, 1985), (2) the Pliocene-Pleistocene boundary is assigned a 1.6 Ma age (Aguirre and Pasini, 1985), and (3) the Proterozoic time scale is from Harrison and Peterman (1982).

Quaternary Sediments

- Qs** Holocene alluvium and peat and late(?) Wisconsin¹ glacial deposits, undivided. The age is assumed from stratigraphic position.
- Qls** Holocene and late(?) Wisconsin¹ landslide deposits. The age is assumed from stratigraphic position.
- Qd** These active and stabilized dunes generally overlie late Wisconsin¹ glacial and outburst flood deposits and are assumed to be of Holocene age.
- Qp** This unit consists of Holocene organic debris deposited in depressions formed on the surfaces of late Wisconsin¹ glacial and outburst flood deposits.

¹ The ages of Quaternary sediments in the northeast quadrant of Washington are uncertain because the deposits have yielded few isotopic ages. Therefore, age assignment is based on correlation with relatively well dated deposits of the Cordilleran ice sheet in the Puget Lowland of western Washington (Blunt and others, 1987) and with deposits of the Laurentide ice sheet in the North American midcontinent (Willman and Frye, 1970). The latest glaciation, which is correlative with the Fraser glaciation in western Washington and the late Wisconsin glaciation in the midcontinent, began approximately 20 ka and ended approximately 10 ka. Older glaciations, which are correlative with pre-Fraser glaciations in western Washington and with early Wisconsin and/or older glaciations in the midcontinent, occurred before 38 ka; some are at least as old as 800-900 ka (Blunt and others, 1987).

- Qla** Holocene lake deposits. Beach(?) gravel that inter-fingers with lacustrine silt contains charcoal that has yielded a ¹⁴C age of 5,870 ± 140 years B.P. (Wagoner, 1990b).
- Qa** Holocene alluvium on modern floodplains and alluvial fans.
- Qoa** Age uncertain. Alluvial deposits in terraces along the Wenatchee River "record aggradation of the Wenatchee valley, probably caused by damming of the Columbia and Wenatchee River by flood bars, perhaps during several [glacial outburst] flood events in Columbia River valley" (Tabor and others, 1982, p. 21). Alluvial fans west of Lake Lenore are dissected and locally capped by caliche.
- Ql** The youngest loess blankets late Wisconsin¹ outburst flood deposits and locally contains Mazama tephra (7 ka) (Patton and Baker, 1978). Older loess commonly exhibits well-developed soil profiles and caliche layers and is intercalated with pre-late Wisconsin outburst flood deposits. The oldest loess is characterized by reversed magnetic polarity, indicating it was deposited prior to 790 ka (McDonald and Busacca, 1989).
- Qfs** This unit is of the same age as unit **Qfg**.
- Qfg** The youngest outburst flood deposits, which are intercalated with Mount St. Helens set S tephra (13 ka) (Mullineaux and others, 1978), were deposited between 15,300 and 12,700 years ago (Waitt, 1985). Older outburst flood deposits exhibit well-developed soil profiles and locally contain ash identified as Mount St. Helens set C tephra (36 ka) (Moody, 1987). The oldest outburst flood deposits are intercalated with loess that is characterized by reversed magnetic polarity, indicating it was deposited prior to 790 ka (McDonald and Busacca, 1989).
- Qgt** Late Wisconsin¹.
- Qgo** Late Wisconsin¹.
- Qgl** Late Wisconsin¹. South of Tiger (T36N, R43E), charcoal yielded a ¹⁴C age of 15,000 ± 170 years B.P. (Joseph, 1990a).
- Qgp** These pre-late Wisconsin¹ deposits are characterized by well-developed soil profiles, caliche horizons, and intensely weathered rock clasts.
- Qgd** In areas covered by late Wisconsin¹ ice, deposits are chiefly of late Wisconsin age. Topographically above and beyond the limit of late Wisconsin ice, this unit includes pre-late Wisconsin till and outwash as well as late Wisconsin outwash and glaciolacustrine deposits.
- Qgfl** Along the Columbia and Sanpoil River valleys, the unit is chiefly of late Wisconsin¹ age (Atwater, 1986). Along the Spokane River and its tributaries, the unit includes both late Wisconsin and pre-late Wisconsin deposits (Rigby, 1982; Kiver and Stradling, 1982a, 1982b).
- Qad** On Calispell Peak (T34N, R42E), poorly developed soils indicate a late Wisconsin¹ age (E. P. Kiver and D. F. Stradling, Eastern Washington Univ., oral commun., 1990). In the Wenatchee River valley (T23N, R19E)

and in the Chelan Mountains (T28N, R19-20E), poorly developed soils indicate that most deposits are late Wisconsin in age, but well-developed soils and weathered clasts in deposits topographically above and beyond the late Wisconsin ice limit indicate a pre-late Wisconsin age (Tabor and others, 1987). Deposition of the pre-late Wisconsin alpine drift apparently occurred between 130 and 150 ka (Tabor and others, 1987).

Sedimentary and Volcanic Rocks

- QRcg The unit contains Pleistocene mammalian fossils and Pliocene-late Pleistocene freshwater molluscs and ostracods (Grolier and Bingham, 1978).
- RMIs The unit is composed of clasts of Wanapum Basalt (14.5-15 Ma). The deeply dissected nature of the topography suggests a pre-Pleistocene age (Tabor and others, 1982).
- RMcg West of Springdale (T29-30N, R38-39E), the unit apparently overlies Grande Ronde Basalt (15.5-16.5 Ma). South of Lake Chelan (T24N, R20E and T26N, R21E), the unit includes Grande Ronde Basalt clasts. The weakly to moderately indurated nature of the conglomerates suggests a pre-Pleistocene age (Tabor and others, 1987).
- Mc This unit is intercalated with Grande Ronde and Wanapum Basalts, which have yielded numerous K-Ar ages between 14.5 and 16.5 Ma outside of the map area. The unit contains Miocene fossil floras (Knowlton, 1926).

Columbia River Basalt Group

- Mv_s This unit is correlative with the Weissenfels Ridge Member of the Saddle Mountains Basalt, which has yielded K-Ar ages between 12 and 13 Ma outside the map area (Swanson and others, 1979b; Beeson and others, 1985; Reidel and Fecht, 1987).
- Mv_w Numerous K-Ar ages between 14.5 and 15.5 Ma have been reported from this unit outside the map area (Swanson and others, 1979b; Beeson and others, 1985; Reidel and Fecht, 1987).
- Mv_g Numerous K-Ar ages between 15.5 and 16.5 Ma have been reported from this unit outside the map area (Swanson and others, 1979b; Beeson and others, 1985; Reidel and Fecht, 1987).
- Mvi_g K-Ar ages between 15.5 and 16.5 Ma have been reported from this unit outside the map area (Swanson and others, 1979b; Beeson and others, 1985; Reidel and Fecht, 1987).
- Φc Silicic tuff beds have yielded two fission-track zircon ages of approximately 34 Ma (Tabor and others, 1982). These sedimentary rocks contain Oligocene palynomorphs and fossil leaves.
- Evd₂ Rocks of this unit have yielded several K-Ar hornblende ages between 41 and 48 Ma (Pearson and Obradovich, 1977).

- Evd₂ K-Ar biotite and whole-rock ages of approximately 48 Ma have been reported for this unit (Pearson and Obradovich, 1977; Stoffel, 1990a).
- Evt₂ Age assignment is based on stratigraphic position. The unit overlies Sanpoil Volcanics (48-52 Ma) (unit Evd₁) and underlies flows of the Klondike Mountain Formation (41-48 Ma) (unit Evd₂).
- Evc₂ Shale beds at several locations contain early middle Eocene floras (Wolfe and Wehr, 1987) and Eocene fossil fish (Pearson, 1967). The unit stratigraphically overlies dacite flows of Sanpoil Volcanics (48-52 Ma) (unit Evd₁) and locally underlies dacite and andesite flows of the Klondike Mountain Formation (41-48 Ma) (unit Evd₂).
- Ec₂ Silicic tuff beds from outside the map area yield fission-track zircon ages between 42 and 49 Ma (Gresens and others, 1981). The rocks contain Eocene palynomorphs and fossil leaves (Newman, 1971, 1975).
- Ecg₂ In the Pend Oreille River valley, the unit contains early to middle Eocene palynomorphs (Harms, 1982). Near Cusick (T33N, R44E), the unit contains clasts of Sanpoil Volcanics (48-52 Ma) (unit Evd₁). Near Wenatchee, age assignment is based on the gradational nature of the contact with unit Ec₂, which has yielded several fission-track zircon ages between 42 and 49 Ma (Gresens and others, 1981).
- Evd₁ Numerous K-Ar hornblende and biotite ages range from 43 to 52 Ma, with the majority between 48 and 52 Ma (Pearson and Obradovich, 1977; Rinehart and Fox, 1972, 1976; Yates and Engels, 1968; Atwater and Rinehart, 1984). Near Island Mountain (T39N, R19E), the unit is cut by an andesite plug (unit Eian), the age of which is 47 Ma (White, 1986).
- Evt₁ Northeast of Cusick (T33N, R44E), the unit has a K-Ar biotite age of approximately 53 Ma (Pearson and Obradovich, 1977). Southeast of Republic, the unit has yielded a K-Ar biotite age of 50 Ma (Atwater and Rinehart, 1984). North of Conconully (T37N, R24E), this unit is cut by a diorite plug (50 Ma) (unit Eid). Elsewhere, the unit's age assignment is based on its stratigraphic position between or beneath flows that have yielded ages of 48-52 Ma (unit Evd₁).
- Ev₁ Northeast of Nespelem, rocks of this unit have yielded a K-Ar biotite age of 49 Ma (B. F. Atwater, USGS, written commun., 1989). Near Island Mountain (T39N, R19E), the unit is stratigraphically overlain by unit Evd₁ and intruded by an andesite plug (47 Ma) (unit Eian). Northwest of Omak (T36N, R25E), age assignment is based on correlation with units Evd₁, Evt₁, and Eida.
- Evc₁ Near Oroville, this unit is cut by a dacite plug (52 Ma) (unit Eida). Elsewhere, age assignment based on stratigraphic position beneath dacite flows, ages of which are between 45 and 48 Ma (unit Evd₁).
- Ec₁ The **Swauk Formation** outside the map area (T21N, R19E) contains tuff that yielded ages of 49-50 Ma (Tabor and others, 1982). Unnamed rocks near Oroville (T40N, R26-27E) were intruded by dacite plugs (51-52 Ma). Sedimentary rocks in the **volcanics of Island**

- Mountain** (T39N, R19E) are stratigraphically overlain by volcanic rocks (unit Ev₁) that are cut by an andesite plug that yielded an age of 47 Ma (unit Eian). Unnamed rocks near Tonasket (T38N, R27E) are stratigraphically overlain by tuff (unit Evt₁) that underlies dacite flows that yielded radiometric ages of 48-52 Ma (unit Evd₁).
- Ec_{g1}** Near Oroville, the unit is cut by dacite plugs that have been radiometrically dated at 51-52 Ma (unit Eida). Northeast of Kettle Falls, the unit is stratigraphically overlain by volcanoclastic rocks (unit Evc₁) that are overlain by dacite flows that range in age from 48 to 52 Ma (unit Evd₁). North of Conconully (T37N, R24E), the unit is stratigraphically overlain by tuff (unit Evt₁) that is, in turn, overlain by dacite flows that range in age from 48 to 52 Ma (unit Evd₁) and is cut by a diorite plug that yielded a radiometric age of 50 Ma (unit Eid).
- TKc** The **Pipestone Canyon Formation** contains minor silicic tuff beds that have yielded a fission-track zircon age of 70 ± 5 Ma (J. A. Vance, Univ. of Washington, written commun., 1990) and a fossil flora that suggests a Paleocene age (K. R. Johnson, Yale Univ., written commun., 1988).
- Kv₂** The early Late Cretaceous (Coniacian) age of this unit is based on the gradational contact with, and similar composition to the Fawn Peak stock (88 Ma) (unit Kid). A maximum age is defined by the stratigraphic position above the late Early Cretaceous to early Late Cretaceous (middle Albian to Turonian) Winthrop Sandstone (unit Kc₂) and Virginian Ridge Formation (units Km₂ and Kcg₂).
- Kvs₂** Age assignment is based on the interfingering relation with the middle Albian to Turonian Winthrop Sandstone (units Kc₂ and Kcg₂).
- Kc₂** The **Winthrop Sandstone** contains an Albian(?) flora. Age assignment of the unit is based on the interfingering relation with the middle Albian to Turonian Virginian Ridge Formation (unit Km₂).
- Km₂** Rocks of this unit contain late Early and early Late Cretaceous pelecypods and belemnites (Albian and Cenomanian) (Barksdale, 1975; Trexler, 1984) and gastropods (Turonian) (V. R. Todd, USGS, written commun., 1989).
- Kcg₂** The **Sophie Mountain Formation** contains a Cenomanian-Campanian(?) flora (Daly, 1912; Little, 1982). The age of **Virginian Ridge Formation** conglomerate is assigned on basis of the interfingering relation with other Virginian Ridge Formation sedimentary rocks (unit Km₂).
- Kcg₁** The unit contains middle to late Albian gastropods, brachiopods, and pelecypods (Maurer, 1958; McGroder and others, 1990).
- Km₁** The **Harts Pass Formation** contains late Early Cretaceous (late Aptian to middle Albian) pelecypods, ammonites, and gastropods (Barksdale, 1975; V. R. Todd, USGS, written commun., 1989). The **Panther Creek Formation** contains early Early to late Early Cretaceous (late Hauterivian to early Albian) pelecypods (Barksdale, 1975) and Aptian-Albian gastropods, cephalopods, and pelecypods (V. R. Todd, USGS, written commun., 1989). The **Goat Creek Formation** contains no fossils in the study area, but Barremian-early Albian fossils have been found in the formation in southern British Columbia (Coates, 1974).
- KJv** Age assignment is based on probable correlation with unit KJvs.
- KJvs** Age uncertain. Northeast of Buck Mountain (T36N, R21E), the unit contains middle Early Cretaceous (Hauterivian to middle Barremian) pelecypods, ammonites, and belemnites (Maurer, 1958; Barksdale, 1975), early to middle Early Cretaceous (Berriasian to Barremian) belemnites, and Late Jurassic to Early Cretaceous (Oxfordian to Valanginian(?)) pelecypods (McGroder and others, 1990). Near Twisp (T33N, R21-22E), the unit is intruded by plutons (unit KJiq) that yield K-Ar hornblende ages of 139 Ma (Bunning, 1990; V. R. Todd, USGS, written commun., 1988), indicating that at least part of the unit is Jurassic.
- KJvt** Age assignment is based on interfingering relation with unit KJvs (Bunning, 1990).
- KJm** Age assignment is based on interfingering relation with unit KJvs (Bunning, 1990).
- Jm** In the map area, the unit contains cycad fragments and belemnite molds that are not age diagnostic. Age assignment is based on correlation with the Early to Middle Jurassic (Sinemurian to Bajocian) Ladner Group in southern British Columbia (Tennyson and Cole, 1987).
- ### Metasedimentary and Metavolcanic Rocks
- KJmv** Age uncertain. The unit is intruded by the Black Peak batholith (90 Ma) (unit Kog). Age assignment is based on possible correlation with the volcanic member of the Midnight Peak Formation (units Kv₂ and Kvc₂) or the Buck Mountain Formation and Newby Group, undivided (units KJvs, KJvt, KJm, and Jm).
- KJmt** Age uncertain. The unit is intruded by the Carlton stocks (unit KJiq), which have yielded a K-Ar hornblende age of 130 Ma (Hopkins, 1987). Age assignment is based on tentative correlation with the Newby Group, undivided (units KJvs, KJvt, KJm, and Jm).
- Jmv** The **Rossland Group** in the map area is overlain(?) by volcanic rocks that have yielded radiometric ages of 165 Ma and overlies metaconglomerate composed of granitic rocks reported to be 196 Ma (Roback and Walker, 1989). The Rossland Group in southern British Columbia contains Early Jurassic (Sinemurian-Toarcian) marine macrofossils (Høy and Andrew, 1988). The Ellemeham Formation unconformably overlies the Anarchist Group (unit Pmm) and Kobau Formation (unit Tmt), and is intruded by the Shakers Bend alkalic complex (unit Jik), which is thought to be correlative with part of the Similkameen composite pluton (170 Ma) (Rinehart and Fox, 1972).
- Jcg** The **Ellemeham Formation** metaconglomerate contains clasts of Ellemeham Formation metavolcanic rocks (unit Jmv) (Rinehart and Fox, 1972). Age assignment of

- the **Rossland Group** metaconglomerate is based on its interfingering relation with Rossland Group metavolcanic rocks (unit Jmv) (Fox, 1981).
- Tmv** **Cave Mountain Formation** metavolcanic rocks stratigraphically overlie Cave Mountain Formation metasedimentary rocks (unit Tmm) that contain Late(?) Triassic ammonites and pelecypods (Misch, 1966; Rinehart and Fox, 1976). The **Palmer Mountain Greenstone** grades into and locally intrudes the Upper Triassic(?) Kobau Formation (unit Tmt) (Rinehart and Fox, 1972). Age assignment of metavolcanic rocks at Buckhorn Mountain (T40N, R31E) is based on the compositional similarity to metavolcanic rocks in the Kobau Formation (unit Tmt).
- Tmm** **Cave Mountain Formation** metasedimentary rocks contain Late(?) Triassic ammonites and pelecypods (Misch, 1966; Rinehart and Fox, 1976). Age assignment of **metamorphic complex of Conconully** metasedimentary rocks is based on their interfingering relation with the Cave Mountain Formation (Rinehart and Fox, 1976). Age assignment of unnamed rocks near Buckhorn Mountain (T40N, R31E) is based on their tentative correlation with the Cave Mountain Formation. The upper part of the **Flagstaff Mountain sequence** (T39-40N, R38-39E) has yielded Late Triassic (Karnian to Norian) conodonts (R. C. Roback, Univ. of Texas at Austin, written commun., 1989; Joseph, 1990a).
- Tcb** **Cave Mountain Formation** metacarbonate rocks interfinger with Upper(?) Triassic Cave Mountain Formation metasedimentary rocks (unit Tmm). The upper part of the **Flagstaff Mountain sequence** (T39N, R38E) contains Late Triassic (Karnian to Norian) conodonts (R. C. Roback, Univ. of Texas at Austin, written commun., 1989; Joseph, 1990a). Unnamed rocks northeast of Curlew (T40N, R34E) have yielded Late Triassic ammonites, brachiopods, pelecypods, and crinoid columnals (Parker and Calkins, 1964).
- Tmt** The unit is intruded by the Similkameen composite pluton (170 Ma) (units Jia and Jik) and the Upper Triassic(?) Loomis pluton (unit Tgd), indicating a minimum age of Late Triassic (Rinehart and Fox, 1972). The relation with older rock units is uncertain.
- Tpmm** This unit locally contains discontinuous layers and pods of metalimestone that have yielded Early Triassic (Scythian) pelecypods, cephalopods, and gastropods and Early to Late Permian fusulinids, crinoids, bryozoans, brachiopods, and corals (Kuenzi, 1961, 1965; Skinner and Wilde, 1966).
- Tpmv** Near Curlew, age assignment is based on the interfingering relation with unit TPmt (Parker and Calkins, 1964). Northwest of Kettle Falls, age assignment is based on spatial relation with unit TPmm (Mills, 1985) and on compositional similarity to unit Tpmv near Curlew.
- TPmt** Locally, this unit contains layers and pods of metalimestone (units Tcb and Pcb) that have yielded Late Triassic ammonites, brachiopods, pelecypods, and crinoid columnals and Early Permian (Wolfcampian to Leonardian) fusulinids, bryozoans, brachiopods, and corals (Parker and Calkins, 1964; Muessig, 1967).
- Rmm** Age uncertain. Age assignment is based on possible correlation with either the Anarchist Group (units Pmm and Pcb) or the Covada Group (units Omm, Ocb, and Omv) (K. F. Fox, Jr., USGS, written commun., 1988; M. T. Smith and G. E. Gehrels, Univ. of Arizona, written commun., 1990; Joseph, 1990b; Stoffel, 1990b).
- Pmm** **Spectacle Formation** (Anarchist Group) metasedimentary rocks in the Okanogan River valley contain Permian brachiopods, pelecypods, and gastropod molds (Rinehart and Fox, 1972). **Anarchist Group** phyllite on Buckhorn Mountain (T40N, R30E) has yielded sheared and fragmented Permian(?) corals (McMillen, 1979). Unnamed metasedimentary rocks north of Kettle Falls contain Permian brachiopods, pelecypods, gastropods, scaphopods, crinoid stems, and plant fragments (Mills and Davis, 1962; B. J. West, 1976; Mills, 1985). Unnamed metasedimentary rocks east of Barstow (T38-39N, R37E) have yielded abundant late Early Permian (Leonardian) macrofossils and conodonts (R. C. Roback, Univ. of Texas at Austin, written commun., 1988; Stoffel, 1990a). Age assignment of rocks west of Northport (T40N, R37-39E) is based on correlation with the Permian Mount Roberts Formation in southern British Columbia (Little, 1982). Age assignment of the rocks northeast of Curlew and along the West Fork Sanpoil River (T35N, R30-32E) is based on tentative correlation with the Anarchist Group (Rinehart and Greene, 1988).
- Pcb** North of Kettle Falls, the unit contains Early to Late Permian (Leonardian to Guadalupian) fusulinids, tetracorals, bryozoans, ostracods, crinoids, and foraminifers (Mills and Davis, 1962; B. J. West, 1976; Mills, 1985); near Curlew Lake (T37N, R33-34E), it contains Early Permian (Wolfcampian to Leonardian) fusulinids, bryozoans, brachiopods, and corals (Muessig, 1967). Age assignment of metalimestone in the Okanogan River valley and on Buckhorn Mountain is based on its interfingering relation with unit Pmm (Rinehart and Fox, 1972; McMillen, 1979). Age assignment of metalimestone near Barstow (T38N, R37E) is based on its proximity and similarity to metalimestone bodies in Lower Permian metasedimentary rocks (unit Pmm) near Toulou Creek. Age assignment of rocks near Swan Lake (T35N, R32E) is based on their tentative correlation with the Anarchist Group (Rinehart and Greene, 1988).
- CDmm** The **Flagstaff Mountain sequence** contains late Middle to Late Devonian (Givetian to Famennian) conodonts (Beka, 1980; Adekoya, 1983). The **Pend Oreille sequence** has yielded Late Devonian conodonts (Beka, 1980). Age assignment of the **Grass Mountain sequence** is based on its correlation with the Flagstaff Mountain and Pend Oreille sequences. Correlative rocks in southern British Columbia contain brachiopods that are no older than Devonian, no younger than Permian, and may be Carboniferous (Little, 1982).
- CDcb** Unnamed rocks near Springdale (T30N, R40E) contain Late Mississippian (Meramecian to Chesterian) con-

- odonts (Waggoner, 1990b). Unnamed rocks east of Valley (T31N, R41E) have yielded Famennian conodonts and macrofossils (Miller and Clark, 1975; Waggoner, 1990b). Age assignment of metacarbonate rocks in the **Grass Mountain sequence** is based on their interfingering relation with metasedimentary rocks in the Grass Mountain sequence (unit CDmm).
- CDmv Age uncertain. The unit is interbedded with metasedimentary rocks of the Flagstaff Mountain sequence (unit CDmm) (Yates, 1971).
- CDmt Age uncertain. The unit is interbedded with metasedimentary rocks of the Flagstaff Mountain and Pend Oreille sequences (unit CDmm) (Yates, 1971).
- C0mm Age uncertain. Near the Columbia River north of Hunters (T30-35N, R36-38E), the unit conformably overlies the Covada Group (units Omm, Ocb, and Omv), suggesting a maximum age of Ordovician (M. T. Smith and G. E. Gehrels, Univ. of Arizona, written commun., 1990; Joseph, 1990b). East of Hunters (T30N, R37E), the unit contains Middle Cambrian to Early Ordovician conodonts, but their occurrence in turbidite sequences suggests that the conodonts may have been reworked from older rocks (Joseph, 1990b). East of Kettle Falls (T36-38N, R38E), these rocks contain late Famennian conodonts (Laskowski, 1982; Sandberg and others, 1988), suggesting that at least part of the unit is Late Devonian. Fossil data and lithologic compositions suggest that this unit may be correlative with unit CDmm.
- COcg The unit is interbedded with unit C0mm.
- COmv The unit is interbedded with unit C0mm.
- Dcg Metaconglomerate clasts contain reworked Lower to Middle Devonian (Lochkovian to Eifelian) conodonts; Middle Devonian (Givetian) conodonts in the conglomerate matrix apparently represent the age of deposition (Greenman, 1976; Greenman and others, 1977).
- Dcb North of Metaline Falls (T40N, R43E), the unit contains Late Devonian (Frasnian) corals and conodonts (Greenman and others, 1977). East of Hunters (T30N, R37E), the unit contains Famennian conodonts (Joseph, 1990b) and brachiopods (Dutro and Gilmour, 1989).
- Smm The unit contains Early Silurian (Llandoveryan to Ludlovian) conodonts (Greenman and others, 1977).
- Omm The unit contains middle to late Early Ordovician conodonts, trilobites, and brachiopods (Snook and others, 1981; R. J. Ross, Jr., Colorado School of Mines, written commun., 1988).
- Ocb This unit is interbedded with unit Omm.
- Omv This unit is interbedded with unit Omm.
- Oar The unit contains Early to Late Ordovician graptolites, conodonts, trilobites, and a crustacean and sponge (Park and Cannon, 1943; Schuster, 1976; Greenman and others, 1977; Snook and others, 1981; Carter, 1989a, 1989b; Schuster and others, 1989).
- OCcb The **Metaline Formation** contains Early Cambrian¹ to Middle Ordovician trilobites, brachiopods, conodonts, and fish (Bennett, 1941; Park and Cannon, 1943; Dings and Whitebread, 1965; Miller and Clark, 1975; Yates, 1976; Repetski, 1978; Repetski and others, 1989; Carter, 1989a, 1989b; Schuster and others, 1989). Age assignment of unnamed rocks north of Cusick (T35N, R43E) is based on correlation with the Metaline Formation.
- OYmm Age uncertain. Lithologic compositions suggest possible correlation with parts of the Metaline Formation (unit OCcb) and Addy Quartzite (unit CZq) (Becraft and Weis, 1963) and/or with part of the Deer Trail Group (units Ymm, Ycb₂, Yar₂, Ycb₁, and Yar₁) (Waggoner, 1990c).
- Cmm The unit is interbedded with unit Ccb.
- Ccb The unit contains Early Cambrian archeocyathids, trilobites, brachiopods, and bryozoans (Dings and Whitebread, 1965; Hampton, 1978).
- CZq The **Addy and Gypsy Quartzites** contain Early Cambrian trilobites, brachiopods, and molluscs (Park and Cannon, 1943; Okulitch, 1951; Miller and Clark, 1975; Lindsey and others, 1990). Fossils are restricted to the upper parts of these formations; the lower parts of the formations are barren of fossils. Lindsey (1987, p. 70) believes that the sudden appearance of fossils in the upper parts of the formations represents "a true 'evolutionary' first appearance, and not a facies controlled first appearance", and that the Precambrian-Cambrian boundary is in the lower parts of the formations.

Windermere Group

The minimum age of the **Windermere Group** is broadly constrained by stratigraphic position beneath the Lower Cambrian and Upper Proterozoic Addy and Gypsy Quartzites (unit CZq). The maximum age is based on the inception of rifting, which probably began shortly before extrusion of the Upper Proterozoic (meta)volcanic rocks of the Huckleberry Formation (unit Zmv) approximately 750 Ma (Devlin and others, 1988).

- Zq Age is constrained only by stratigraphic position within the Windermere Group.
- Zmm Age is constrained only by stratigraphic position within the Windermere Group.
- Zmv K-Ar ages from metavolcanic rocks of the **Huckleberry Formation** range from 238 to 928 Ma and cluster between 613 and 862 Ma (Miller and others, 1973), but because the Rb-Sr isotopic systems of the dated metavolcanic rocks have been severely disturbed, the K-Ar ages probably do not represent the age of extrusion (Devlin and others, 1985). Sm-Nd isochron ages of 674 ± 212 Ma, 762 ± 44 Ma, and 795 ± 115 Ma are also reported from the Huckleberry Formation (Devlin

¹ Early Cambrian fossils have been recovered only from rocks near Addy, which were originally assigned to the Old Dominion Limestone; elsewhere, the oldest fossils in this unit are Middle Cambrian.

and others, 1988); the 762 Ma age is apparently the most reliable of the ages and is thought to broadly represent the extrusive age of the metavolcanic rocks (Devlin and others, 1988).

Zcg Age is constrained only by stratigraphic position within the Windermere Group.

Deer Trail Group

Ymm, Ycb₂, Yar₂, Ycb₁, Yar₁

Age assignment of the Deer Trail Group is based on correlation of the entire group with units Yms₃ and Yms₄ of the Belt Supergroup (Miller and Whipple, 1989).

Priest River Group

Yar, Yq, Ycb

Age assignment of the **Priest River Group** is based on correlation with the Deer Trail Group (units Ymm, Ycb₂, Yar₂, Ycb₁, and Yar₁), which is thought to be correlative with units Yms₃ and Yms₄ of the Belt Supergroup (Miller and Whipple, 1989).

Belt Supergroup

The **Belt Supergroup** in the northern Rocky Mountains is unconformably overlain by Windermere Group metavolcanic rocks (unit Zmv) (762 Ma) (Miller and others, 1973) and unconformably overlies crystalline basement rocks that have yielded isotopic ages of approximately 1,700 Ma in Canada (Burwash and others, 1962). The age range of the Belt Supergroup shown on the correlation diagram (Sheet 3) follows the interpretative geochronology of Obradovich and others (1984) and Harrison (1984).

Yms₄ Age uncertain. Correlative rocks in western Montana have yielded Rb-Sr isochron ages of approximately 1,100 Ma (Obradovich and others, 1984).

Yms₃ Age is constrained only by stratigraphic position within the Belt Supergroup.

Yms₂ Age is constrained only by stratigraphic position within the Belt Supergroup.

Yms₁ Age uncertain. The unit is intruded by abundant meta-diorite sills, one of which (near Bonners Ferry, Idaho) has yielded a U-Pb zircon age of $1,433 \pm 13$ Ma (Zartman and others, 1982).

Intrusive Rocks

Mian The **andesite and dacite of Burch Mountain** has yielded K-Ar hornblende ages of 10-11 Ma (Engels and others, 1976; Tabor and others, 1987).

Qian The **hornblende andesite porphyry complex of Horse Lake Mountain** has yielded K-Ar hornblende ages of 25-30 Ma (Tabor and others, 1982).

Eian The **volcanics of Island Mountain** have yielded a K-Ar whole-rock age of 47 Ma (White, 1986).

Unnamed andesite near Oroville (T40N, R27E) intrudes Eocene conglomerate (unit Ecg₁).

Eir Unnamed rhyolite near Lake Chelan (T26-28N, R20-21E) intrudes the Duncan Hill pluton (48 Ma) (unit Eig) (Tabor and others, 1987).

Unnamed rhyolite dikes near Eightmile Creek (T39N, R20E) probably emanate from the Monument Peak stock (49 Ma) (Tabor and others, 1968).

The **intrusive rhyolite near West Fork** has yielded a K-Ar biotite age of 50 Ma (Atwater and Rinehart, 1984).

Eitr Unnamed trachyte east of Chesaw (T40N, R31E) intrudes the Sanpoil Volcanics (48-52 Ma) (unit Eida) (Pearson, 1967).

Eida Numerous radiometric ages (K-Ar biotite and hornblende, fission-track zircon and apatite) range from 45 to 54 Ma and cluster between 48 and 52 Ma (Joseph, 1990b; Stoffel, 1990a, 1990b; Waggoner, 1990b, 1990c).

Ei The **intrusive complex north of Grant Lake** is a heterogeneous mixture of units TKig, Eimd, Eida, and Eig. Formation of the complex probably occurred approximately 50 Ma (Gulick and Korosec, 1990b).

Eik Unnamed alkalic rocks northeast of Kettle Falls (T38N, R38E) have yielded a K-Ar biotite age of 51 Ma (Yates and Engels, 1968).

Unnamed alkalic rocks near Chesaw (T38N, R31E and T39N, R29E) intrude the O'Brien Creek Formation (unit Evc₁) (Pearson, 1967; Fox, 1973).

Unnamed alkalic rocks near Oroville intrude Eocene conglomerate (unit Ecg₁) (Fox, 1973).

Eis **Coryell intrusive rocks** have yielded K-Ar biotite ages of 51-52 Ma (Yates and Engels, 1968).

Eiat Preliminary U-Pb geochronology and Sr isotopic data (from outside the map area) suggest that the **Mount Rathdrum granite** is approximately 52 Ma (Bickford and others, 1985; Armstrong and others, 1987).

Eia The **Silver Point Quartz Monzonite** has yielded K-Ar biotite and hornblende ages of 48-52 Ma (Miller and Engels, 1975; J. C. Engels in Miller and Clark, 1975).

Eig The **Duncan Hill pluton** has yielded a K-Ar biotite age of 46 Ma (Cater and Crowder, 1967; Tabor and others, 1987).

The **Monument Peak stock** has yielded a K-Ar biotite age of 48 Ma (Tabor and others, 1968).

The **granite of Deadhorse Creek** has yielded a K-Ar biotite age of 50 Ma (Atwater and Rinehart, 1984).

The **Fire Mountain pluton** granite intrudes Fire Mountain pluton quartz monzonite (50-53 Ma) (Holder, 1985; Pearson and Obradovich, 1977).

The **Sheppard Granite** intrudes the Upper Cretaceous Sophie Mountain Formation (unit Kcg₂), is cut by lamprophyre dikes (53 Ma) (Yates and Engels, 1968), and is thought to be a phase of the Coryell intrusive rocks, which have yielded radiometric ages of 51-52 Ma (Little, 1982).

- The **Long Alec Creek pluton** has yielded a K-Ar biotite age of 53 Ma (Stoffel, 1990a).
- The **granite of upper Stepstone Creek** intrudes the Moses Mountain pluton (unit TKia) and unit Eida (Joseph, 1990b).
- Unnamed granite near Orient (T40N, R37E) intrudes the O'Brien Creek Formation (unit Evc₁) and Sanpoil Volcanics (unit Evd₁) (J. R. West, 1976).
- Eiqm The granodiorite phase of the **Cooper Mountain batholith** (unit Eigd) has yielded a radiometric age of 48 Ma (Tabor and others, 1987; V. R. Todd, USGS, written commun., 1990).
- The **Empire Lakes pluton** has yielded a K-Ar biotite age of 50 Ma (Stoffel, 1990a).
- The **Fire Mountain pluton** has yielded a K-Ar biotite age of 51 Ma (Pearson and Obradovich, 1977).
- Unnamed quartz monzonite near Loon Lake (T30N, R41E) has yielded a K-Ar biotite age of 51-52 Ma (J. C. Engels in Miller and Clark, 1975).
- The **Long Alec Creek pluton** has yielded a K-Ar biotite age of 52-53 Ma (Pearson and Obradovich, 1977; Stoffel, 1990b).
- Monzonite east of Storm King Mountain** intrudes unit Eimd (Holder, 1990).
- Eigd The **Duncan Hill pluton** has yielded a U-Pb zircon age of 48 Ma (Tabor and others, 1987) and K-Ar hornblende and biotite ages of 43-48 Ma (Cater and Crowder, 1967; Tabor and others, 1987).
- The **Cooper Mountain batholith** has yielded a K-Ar biotite age of 48 Ma (Tabor and others, 1987; V. R. Todd, USGS, written commun., 1990; Bunning, 1990).
- The **granodiorite of Joe Moses Creek** has yielded a K-Ar biotite age of 48 Ma (B. F. Atwater, USGS, written commun., 1989; Joseph, 1990b).
- The **Lost Peak stock** has yielded K-Ar biotite and fission-track zircon ages of 48-52 Ma (White, 1986).
- The **Noname stock** has yielded fission-track titanite and zircon ages of 54-56 Ma (Buddington, 1986).
- Eimd The **Kettle Crest pluton** has yielded K-Ar biotite and hornblende ages of 45-48 Ma (Atwater and Rinehart, 1984).
- The **Friedlander Meadows pluton** has yielded K-Ar biotite and hornblende ages of 47-48 Ma (B. F. Atwater, USGS, written commun., 1987; Waggoner, 1990c).
- The **Swimptkin Creek pluton** has yielded K-Ar biotite and hornblende ages of 45-50 Ma (Fox and others, 1976; Atwater and Rinehart, 1984; B. F. Atwater, USGS, written commun., 1987; Gulick and Korosec, 1990b).
- The **Devils Elbow pluton** has yielded a K-Ar biotite age of 50 Ma (Atwater and Rinehart, 1984).
- Unnamed monzodiorite near Springdale (T29N, R40E) has yielded a K-Ar biotite age of 50 Ma (J. C. Engels in Miller and Clark, 1975).
- The **Henry Creek diorite** has yielded K-Ar biotite and hornblende ages of 51-54 Ma (Pearson and Obradovich, 1977; Stoffel, 1990a).
- Eiq The **Hungry Mountain stock** has yielded a K-Ar hornblende age of 45 Ma (Hopkins, 1987).
- Eid Unnamed rocks near Sinlahekin Creek (T37N, R24E) intrude middle Eocene tuff (unit Evt₁) (Stoffel, 1990b).
- The **diorite of Little Moses Mountain** (T33N, R30E) intrudes the Moses Mountain pluton (unit TKia) and is cut by dikes of the granite of upper Stepstone Creek (unit Eiq) (Atwater and Rinehart, 1984).
- Eib Unnamed gabbro west of Toroda (T40N, R32E) intrudes middle Eocene volcanoclastic rocks (unit Evc₂) (Pearson, 1967).
- Unnamed gabbro near Wenatchee (T22N, R20E) intrudes the Chumstick Formation (unit Ec₂) and is cut by Oligocene andesite plugs (unit Qian) (Gresens, 1983).
- Unnamed gabbro and diorite near Kettle Falls intrude Triassic-Permian metasedimentary rocks (unit TPmm); because they have not been metamorphosed, the intrusions are thought to be post-Middle Jurassic in age.
- Rit The **Oval Peak batholith** has yielded U-Pb zircon and titanite ages of 61-65 Ma (Miller and Walker, 1987; Miller and others, 1989).
- TKiat The **granite of Swawilla basin** has yielded K-Ar biotite ages of 54 and 59 Ma from the border phase (unit TKia) (Atwater and Rinehart, 1984; B. F. Atwater, USGS, written commun., 1987; Waggoner, 1990c); these are minimum crystallization ages¹.
- TKiaa The **garnet-bearing granite of McGinnis Lake** (T30N, R31E) intrudes the quartz porphyry of Mount Tolman and the porphyritic granite of Keller Butte (unit TKig), which have yielded minimum crystallization ages¹ of 56-61 Ma (K-Ar biotite, muscovite, and sericite) (Atwater and Rinehart, 1984).
- Age assignment of alaskitic rocks in the **Coyote Creek, Moses Mountain, and Mount Bonaparte plutons** is based on K-Ar (biotite and muscovite) and fission-track (allanite and apatite) ages of 49-59 Ma from the main phases of the plutons (unit TKia), which are thought to represent minimum crystallization ages¹ (Fox and others, 1976; Naeser and others, 1970; Atwater and Rinehart, 1984).
- The alaskite-aplite-pegmatite phase of the **Mount Spokane granite** cuts the granitic phase (unit Kiat), the age of which is thought to be approximately 90 Ma. The alaskite-aplite-pegmatite has yielded Rb-Sr whole-rock + muscovite ages of 40 and 47 Ma and Rb-Sr whole-rock ages of 84 and 113 Ma (Armstrong and others, 1987); some of the latter ages came from samples from the

¹ Numerous K-Ar and fission-track ages between 47 and 61 Ma have been reported from Tertiary-Cretaceous intrusive rocks (units TKiat, TKiaa, TKia, TKig, and TKigd) in the Kettle, Okanogan, Lincoln, and Priest River metamorphic core complexes. These ages are thought to represent the age of uplift and cooling of the core complexes and not the magmatic crystallization ages of the intrusions, which are unknown.

granitic phase of the pluton. These data suggest a Late Cretaceous magmatic age and an Eocene metamorphic age.

Unnamed alaskite-aplite-pegmatite south of Spokane (T24N, R42-44E) intrudes Precambrian high-grade metamorphic rocks (unit pChm); age assignment is based on compositional similarity to unit TKiaa north-east of Spokane (Joseph, 1990c).

TKia The **porphyritic granodiorite of Manila Creek** has yielded K-Ar biotite and hornblende ages of 41-47 Ma (B. F. Atwater, USGS, written commun., 1989; Joseph, 1990b); these are minimum crystallization ages¹.

The **Moses Mountain pluton** has yielded K-Ar biotite ages of 48-49 Ma (Atwater and Rinehart, 1984; B. F. Atwater, USGS, written commun., 1987; Gulick and Korosec, 1990b); these are minimum crystallization ages¹.

The **plutonic complex of Johnny George** has yielded K-Ar biotite and hornblende ages of 49-50 Ma (B. F. Atwater, USGS, written commun., 1987; Waggoner, 1990c); these are minimum crystallization ages¹.

The **Mount Bonaparte pluton** has yielded K-Ar (biotite and muscovite) and fission-track (allanite and apatite) ages of 49-59 Ma (Fox and others, 1976; Naeser and others, 1970); these are minimum crystallization ages¹.

The **granite of Swawilla Basin** has yielded K-Ar biotite ages of 54-59 Ma (Atwater and Rinehart, 1984; B. F. Atwater, USGS, written commun., 1987; Waggoner, 1990c); these are minimum crystallization ages¹.

Age assignment of the **Fifteenmile Creek pluton** is based solely on compositional similarity to other unit TKia plutons.

Age of unnamed acidic intrusive rocks elsewhere in the map area is unknown, but mineralogic compositions and spatial associations suggest that they are either Late Cretaceous or early Tertiary.

TKig The **Coyote Creek pluton** has yielded K-Ar biotite ages of 49-51 Ma (Fox and others, 1976; Atwater and Rinehart, 1984); these are minimum crystallization ages¹.

The **granite of Daisy Trail** has yielded a K-Ar biotite age of 50 Ma (Atwater and Rinehart, 1984); this is a minimum crystallization age¹.

The **granite west of Armstrong Mountain** has yielded a K-Ar biotite age of 53 Ma (Atwater and Rinehart, 1984); this is a minimum crystallization age³.

The **quartz porphyry of Mount Tolman** has yielded K-Ar (sericite and muscovite) ages of 56-58 Ma (Atwater and Rinehart, 1984); these ages were determined

from secondary minerals in hydrothermally altered rocks and are therefore minimum crystallization ages¹.

Unnamed granitic rocks east of Nespelem (T31N, R29E) have yielded K-Ar biotite and muscovite and Rb-Sr isochron ages of 50-59 Ma (B. F. Atwater, USGS, written commun., 1987; R. J. Fleck, USGS, written commun., 1989; Gulick and Korosec, 1990b); these are minimum crystallization ages¹.

The **porphyritic granite of Keller Butte** has yielded K-Ar biotite ages of 53-61 Ma (Atwater and Rinehart, 1984; B. F. Atwater, USGS, written commun., 1987; Waggoner, 1990c); these are minimum crystallization ages¹.

TKigd The **porphyritic granodiorite southwest of Omak Lake** has yielded a K-Ar biotite age of 53 Ma (Atwater and Rinehart, 1984); this is a minimum crystallization age³.

TKit The **Arbuckle Mountain pluton** has yielded a K-Ar biotite age of 81.2 ± 3.0 Ma (Gulick and Korosec, 1990a), but because it intrudes rocks of the Chelan complex that yield K-Ar biotite and hornblende ages between 56 and 60 Ma, the accuracy of the 81 Ma age is questionable; a Late Cretaceous or early Tertiary crystallization age is suspected.

TKid Age assignment is based on interpretation of the unit as a border phase of the Tertiary-Cretaceous Mount Bonaparte pluton (unit TKia) (Cheney and others, 1982; Orr and Cheney, 1987).

Mkigb Age uncertain. The unit intrudes the Salmon Creek schists and gneisses (units pJhm and pJmb).

Kida Unnamed dacite and andesite dikes and sills yielded K-Ar hornblende ages of 82-93 Ma (V. R. Todd, USGS, written commun., 1988; Stoffel and McGroder, 1990).

Kiat Unnamed two-mica granite northwest of Spokane (T27N, R42E) has yielded K-Ar biotite and muscovite ages of 48-53 Ma (Miller and Engels, 1975), but these ages are probably reset; the magmatic crystallization age is interpreted as approximately 100 Ma (Joseph, 1990c).

The **granodiorite of Hall Mountain** has yielded K-Ar muscovite and biotite ages of 87-99 Ma (Miller and Engels, 1975).

The **monzogranite of Hungry Mountain** has yielded K-Ar biotite and muscovite ages of 92-97 Ma (Miller and Engels, 1975).

The **monzogranite of Granite Pass** has yielded a K-Ar muscovite age of 98 Ma (Miller and Engels, 1975).

The **Phillips Lake Granodiorite** has yielded K-Ar biotite and muscovite ages that range from 50 to 86 Ma (Miller and Engels, 1975; J. C. Engels in Miller and Clark, 1975); these ages reflect cooling during Late Cretaceous and early Tertiary uplift and/or deformation. The apparent magmatic crystallization age is approximately 100 Ma (J. C. Engels in Miller and Clark, 1975).

Unnamed two-mica granite south and west of Lone (northern extension of the **Phillips Lake Granodio-**

¹ Numerous K-Ar and fission-track ages between 47 and 61 Ma have been reported from Tertiary-Cretaceous intrusive rocks (units TKiaa, TKia, TKig, and TKigd) in the Kettle, Okanogan, Lincoln, and Priest River metamorphic core complexes. These ages are thought to represent the age of uplift and cooling of the core complexes and not the magmatic crystallization ages of the intrusions, which are unknown.

- rite**, in T35-38N, R41-43E) has yielded K-Ar biotite and muscovite ages of 49-102 Ma (Yates and Engels, 1968; Miller and Engels, 1975); these ages reflect cooling during Late Cretaceous and early Tertiary uplift and/or deformation. The apparent magmatic crystallization age is approximately 100 Ma (J. C. Engels in Miller and Clark, 1975).
- The **granodiorite of Dubious Creek** has yielded K-Ar biotite and muscovite ages of 96-102 Ma (Miller and Engels, 1975); these ages are from outside of the map area.
- The **Blickensderfer Quartz Monzonite** has yielded K-Ar biotite and muscovite ages of 101-102 Ma (Miller and Engels, 1975).
- Unnamed two-mica granite south of Sullivan Lake (T37N, R43-44E) has yielded K-Ar biotite and muscovite ages of 103 Ma (Miller and Engels, 1975).
- The **Mount Spokane granite** has yielded K-Ar biotite and muscovite ages between 47 and 49 Ma (Miller and Engels, 1975), a Rb-Sr (whole rock + biotite) age of 51 Ma (Armstrong and others, 1987), Rb-Sr whole-rock ages of 84 and 159 Ma (some of these samples coming from the pegmatite phase) (Armstrong and others, 1987), and U-Pb zircon ages between 94 and 143 Ma (Armstrong and others, 1987). These data suggest a Late Cretaceous (\approx 90 Ma) magmatic age (with older inherited zircon) and an Eocene metamorphic age (Armstrong and others, 1987).
- Kiaa Aplite east of Malott (T32N, R25E) is a phase of the **equigranular granite of Virginia Lake**, which has yielded K-Ar biotite ages of 65 and 75 Ma (B. F. Atwater, USGS, written commun., 1987; Gulick and Korosec, 1990b); these are minimum crystallization ages.
- Unnamed alaskite east of Deer Lake (T30N, R42E) has yielded a K-Ar muscovite age of 80 Ma (J. C. Engels in Miller and Clark, 1975).
- Alaskite northeast of Colville (T38N, R41E) is a phase of the **Spirit pluton**, which has yielded radiometric ages of 93-102 Ma (Yates and Engels, 1968).
- Unnamed alaskite-aplite-pegmatite east of Ruby (T35N, R44E) intrudes the Galena Point Granodiorite (101 Ma) (Miller and Engels, 1975).
- Unnamed alaskite-aplite-pegmatite north of Fort Spokane (T28-29N, R35-36E) intrudes and/or grades into granite and granodiorite plutons, the ages of which are probably 100 Ma (Waggoner, 1990c).
- Kia The **Bottle Spring pluton** has yielded a K-Ar biotite age of 89 Ma (V. R. Todd, USGS, written commun., 1988; Stoffel, 1990b).
- The **Evans Lake pluton** has yielded a K-Ar biotite age of 89 Ma (from ductilely deformed rocks along the southern border of the pluton) (Rinehart and Fox, 1976); this is a minimum crystallization age.
- Unnamed monzogranite southwest of Springdale (T28-29N, R37-39E) has yielded ages ranging from 63 to 105 Ma (U-Pb zircon and monazite, Pb-Th zircon and monazite, Pb- α , Rb-Sr isochron, and fission-track apatite and titanite) (Becraft and Weis, 1963; Ludwig and others, 1981; Asmerom and others, 1988); the age data suggest that the pluton is either a composite of intrusions (older on the east) (Asmerom and others, 1988) or an Upper Cretaceous intrusion that was locally metamorphosed (western part) during middle Eocene plutonism (Ludwig and others, 1981).
- The **Aeneas Creek pluton** has yielded K-Ar hornblende and biotite ages of 92-98 Ma (Rinehart and Fox, 1976; Stoffel, 1990b).
- The **Cathedral batholith** has yielded K-Ar biotite ages of 94-98 Ma (Hawkins, 1968; Berry and others, 1976).
- The **Fan Lake Granodiorite** has yielded K-Ar biotite and hornblende ages of 96-98 Ma (Miller and Engels, 1975).
- The **Starvation Flat Quartz Monzonite** has yielded K-Ar biotite and hornblende and Pb- α zircon ages of 99-100 Ma (Larsen and others, 1958; Miller and Engels, 1975; J. C. Engels in Miller and Clark, 1975).
- Unnamed granite east of Colville (T35N, R40-41E) has yielded a K-Ar biotite age of 100 Ma (Miller and Engels, 1975).
- The **Galena Point Granodiorite** has yielded a K-Ar biotite age of 101 Ma (Miller and Engels, 1975).
- Kig The **equigranular granite of Virginia Lake** has yielded K-Ar biotite ages of 65-75 Ma (B. F. Atwater, USGS, written commun., 1988; Gulick and Korosec, 1990b); these are minimum crystallization ages.
- The **monzogranite of Sand Creek** has yielded a K-Ar biotite age of 99 Ma (Miller and Engels, 1975).
- The **Spirit pluton** has yielded K-Ar muscovite, biotite, and hornblende ages of 93-102 Ma (Yates and Engels, 1968).
- Kiqm The **Leader Lake quartz monzonite** has yielded a Rb-Sr mineral isochron age (combined with a sample of the Conconully pluton) of 82 Ma and a fission-track (apatite) age of 94 Ma (Menzer, 1970).
- The **Horseshoe Mountain pluton** has yielded K-Ar hornblende and biotite ages of 95-103 Ma (Stoffel, 1990b).
- Kigd The **granodiorite of Soap Lake Mountain** has yielded a K-Ar biotite age of 55 Ma (B. F. Atwater, USGS, written commun., 1987); this is a minimum crystallization age.
- The **Texas Creek stock** has yielded K-Ar hornblende ages of 84-87 Ma (V. R. Todd, USGS, written commun., 1990; Bunning, 1990); these reported ages from the dioritic phase of the stock are minimum crystallization ages.
- The **Pasayten stock** has yielded K-Ar biotite and hornblende ages of 85-88 Ma (Tabor and others, 1968).
- The **Conconully pluton** has yielded Rb-Sr mineral isochron, fission-track (titanite and apatite), and Pb- α zircon ages of 82-90 Ma (Menzer, 1970). Discordant K-Ar ages between 62 and 81 Ma (Rinehart and Fox, 1976; Berry and others, 1976) "suggest that a younger

- thermal event, centered east of the pluton, may have reduced the apparent mineral ages in the eastern part of the pluton" (Rinehart and Fox, 1976, p. 38).
- Unnamed granodiorite southwest of Springdale (T27N, R37-39E) has yielded Pb- α (whole rock) ages of 95-105 Ma (Becraft and Weis, 1963).
- Unnamed granodiorite southeast of Lone (T36N, R44E) has yielded K-Ar biotite and hornblende ages of 99-100 Ma (Miller and Engels, 1975).
- Unnamed granodiorite southeast of Ruby (T35N, R44E) has yielded K-Ar biotite and hornblende ages of 100-104 Ma (R. J. Fleck, USGS, written commun., 1989; Joseph, 1990a).
- Age assignment of unnamed granodiorite in the **Okanogan batholithic complex** is based on ages reported from the tonalitic phase of the complex (unit Kit).
- Kit** The **Cardinal Peak pluton** has yielded a U-Pb zircon age of 75 Ma (Miller and others, 1989).
- The **Entiat pluton** has yielded concordant U-Th-Pb zircon ages of 75-85 Ma, K-Ar biotite and hornblende ages of 56-73 Ma, and fission-track (apatite and zircon) ages of 48-58 Ma (Tabor and others, 1987); the ages indicate Late Cretaceous magmatic crystallization and early Tertiary uplift and cooling.
- Unnamed tonalite in the **Okanogan batholithic complex** has yielded a preliminary U-Pb zircon age of approximately 112 Ma (Hugh Hurlow, Univ. of Washington, oral commun., 1990) that apparently represents a magmatic crystallization age; numerous fission-track (apatite and titanite), K-Ar (biotite and muscovite), and Rb-Sr isochron ages between 76 and 108 Ma represent the ages of uplift and cooling of the complex (Stoffel, 1990b; Stoffel and McGroder, 1990; Gulick and Korosec, 1990b).
- Kiq** The **Lone Frank pluton** has yielded K-Ar hornblende and biotite ages of 93-97 Ma (Stoffel, 1990b).
- Kid** Diorite in the **Texas Creek stock** has yielded K-Ar hornblende ages of 84-87 Ma (V. R. Todd, USGS, written commun., 1990; Bunning, 1990); these are minimum crystallization ages.
- The **Fawn Peak stock** has yielded K-Ar biotite ages of 87-88 Ma (Riedell, 1979a, 1979b; V. R. Todd, USGS, written commun., 1989).
- The **Aeneas Creek pluton** has yielded K-Ar hornblende and biotite ages of 92-98 Ma (Rinehart and Fox, 1976; Stoffel, 1990b); these reported ages are from the central phase of the pluton (unit Kia).
- The **McFarland Creek stock** is intruded by the Cooper Mountain batholith (48 Ma) (units Eigd and Eiqm) and the Hungry Mountain stock (45 Ma) (unit Eiq) (Bunning, 1990); the magmatic crystallization age is uncertain.
- Kigb** Age assignment of gabbro in the **Entiat pluton** is based on concordant U-Th-Pb ages of 75-85 Ma reported from the tonalitic phase (unit Kit) of the pluton (Tabor and others, 1987).
- KJia** The **granite and granodiorite near Meteor** has yielded K-Ar biotite ages of 55 and 61 Ma and a K-Ar hornblende age of 110 Ma (Atwater and Rinehart, 1984; B. F. Atwater, USGS, written commun., 1989; Joseph, 1990b); the hornblende age is a minimum crystallization age.
- Other unit KJia intrusions cut Paleozoic and/or lower Mesozoic metasedimentary rocks and are locally intruded by lower Tertiary-Upper Cretaceous plutons (units TKia and TKig).
- KJigd** The **granodiorite of Rogers Bar** has yielded a K-Ar hornblende age of 71 Ma (Atwater and Rinehart, 1984); this is a minimum crystallization age.
- The **Anderson Creek pluton** has yielded discordant K-Ar biotite and hornblende ages of 100 and 116 Ma, respectively (Engels and others, 1976); the hornblende age is a minimum crystallization age.
- The **Whiskey Mountain pluton** has yielded a K-Ar hornblende age of 119 Ma (Stoffel, 1990b); this is a minimum crystallization age.
- Other unit KJigd intrusions cut Paleozoic and/or lower Mesozoic metasedimentary rocks; the **Wauconda pluton** (T37N, R30-31E) is intruded by the lower Tertiary-Upper Cretaceous Mount Bonaparte pluton (unit TKia).
- KJiq** Granodiorite in the **Whiskey Mountain pluton** has yielded a K-Ar hornblende age of 119 Ma (Stoffel, 1990b), which is a minimum age of crystallization.
- The **Carlton stocks** have yielded K-Ar hornblende ages of 130 Ma (Hopkins, 1987); because the stocks are locally deformed and metamorphosed, this age is a minimum age of crystallization (Bunning, 1990).
- The **Alder Creek stock** has yielded K-Ar hornblende ages of 137-139 Ma (V. R. Todd, USGS, written commun., 1990; Bunning, 1990); because the stock is locally deformed and metamorphosed, these ages are minimum ages of crystallization (Bunning, 1990).
- Quartz diorite in the **Frazer Creek complex (Wolf Canyon quartz diorite)** has yielded K-Ar biotite ages of 104-109 Ma and K-Ar hornblende ages of 130-139 Ma (V. R. Todd, USGS, written commun., 1990; Bunning, 1990); the strong discordance between biotite and hornblende ages indicates that the hornblende ages are minimum ages of crystallization.
- The **Bowers quartz diorite** (part of the **Toats Coulee pluton**) intrudes the Upper Triassic(?) Loomis pluton (unit Figd).
- KJid** Age and contact relations uncertain.
- KJigb** Age assignment of gabbro in the **Frazer Creek complex (Red Shirt gabbro)** is based on K-Ar hornblende ages of 130-139 Ma reported from quartz diorite in the complex (unit KJiq); these are minimum ages of crystallization.
- Age and contact relations of **gabbro near Stranger Creek** (T32N, R36E) are uncertain.
- Jik** Age assignment of alkalic rocks in the **Similkameen composite pluton** is based on several ages of approx-

- imately 171 Ma reported from the central phase of the pluton (unit Jigd) (Engels, 1971; Buddington and Burmester, 1990).
- Age assignment of **Shankers Bend alkalic complex** is based on its compositional similarity to the alkalic rocks in the Similkameen composite pluton, which suggests "that they were probably emplaced during the same episode of plutonism and are derived from the same magmatic source" (Rinehart and Fox, 1972, p. 59).
- Age of the **intrusive rocks of Shasket Creek** is uncertain; they intrude Triassic-Permian metasedimentary and metavolcanic rocks (unit $\bar{\text{T}}\text{Pmt}$) and are cut by Eocene dacite dikes (unit Eida) (Parker and Calkins, 1964).
- Jia The **Similkameen composite pluton** has yielded Rb-Sr isochron and K-Ar (hastingsite) ages of 171 Ma (Engels, 1971; Buddington and Burmester, 1990); a concordant U-Pb zircon age of 170 ± 2 Ma is reported from the northern border of the pluton near Keremeos, British Columbia (Parkinson, 1985).
- The **Lane Mountain pluton** has yielded K-Ar biotite and hornblende ages of 156-172 Ma (Miller and Engels, 1975; J. C. Engels in Miller and Clark, 1975; F. K. Miller, USGS, oral commun., 1988; Waggoner, 1990b).
- Jiqm The unnamed pluton west of Palmer Lake (T39N, R24E) has yielded discordant K-Ar biotite and hornblende ages of 131 and 162 Ma, respectively (Stoffel, 1990b); the hornblende age is a minimum crystallization age.
- Jigd The **Blue Goat pluton** has yielded discordant K-Ar biotite and hornblende ages of 99 and 142 Ma, respectively (Rinehart and Fox, 1976); the hornblende age is a minimum crystallization age.
- Jit The **Button Creek stock** has yielded K-Ar hornblende ages of 150-153 Ma (V. R. Todd, USGS, written commun., 1990; Stoffel and McGroder, 1990).
- Jiq Unnamed quartz diorite west of Northport (T40N, R37E) grades into basic intrusive rocks (unit Jib) that are compositionally similar to, and locally grade into metavolcanic rocks in the Jurassic Rossland Group (unit Jmv) (Rhodes, 1980).
- Jib Unnamed basic intrusive rocks west of Northport (T39-40N, R37-38E) intrude Permian metasedimentary rocks (unit Pmm) and the Jurassic Rossland Group (unit Jmv); "similarities in composition between the Rossland Group and the related mafic intrusive rocks along with the gradational nature of at least some of the contacts between the two suggests [sic] that the gabbros may be the coarse-grained equivalent of the basalts" (Rhodes, 1980, p. 33).
- Jigd The **Loomis pluton** has yielded weakly discordant K-Ar biotite and hornblende ages of 179 and 194 Ma, respectively (Rinehart and Fox, 1972); the hornblende age is a minimum crystallization age.
- The **Flowery Trail Granodiorite** has yielded K-Ar biotite ages of 66-100 Ma and K-Ar hornblende ages of 146-198 Ma (J. C. Engels in Miller and Clark, 1975); hornblende ages are minimum crystallization ages.
- $\bar{\text{T}}\text{ib}$ Because the **Chopaka intrusive complex** was metamorphosed in the Early Jurassic (K-Ar actinolitic hornblende age of the gabbro = 190 ± 15.6 Ma [Hibbard, 1971]) and was intruded by the Loomis pluton (unit $\bar{\text{T}}\text{igd}$) in the Late Triassic(?), the magmatic crystallization age of the complex must be Triassic or older.
- Other unit $\bar{\text{T}}\text{ib}$ intrusions in the map area (T36-40N, R26-30E) intrude the Anarchist Group (units Pmm and Pcb) and Kobau Formation (unit $\bar{\text{T}}\text{mt}$). The composition of the basic intrusive rocks is similar to that of basic igneous rocks in the Chopaka intrusive complex (unit $\bar{\text{T}}\text{ib}$), the Palmer Mountain Greenstone (unit $\bar{\text{T}}\text{mv}$), and the Kobau Formation (unit $\bar{\text{T}}\text{mt}$) (Hibbard, 1971; Rinehart and Fox, 1972); stratigraphic and cross-cutting relations suggest that all are roughly coeval.
- $\bar{\text{R}}\text{ib}$ Age uncertain. The unit intrudes Paleozoic metasedimentary rocks (unit $\bar{\text{R}}\text{mm}$).
- Oigb Age uncertain. The unit intrudes metasedimentary rocks of the Ordovician Covada Group (units Omm and Ocb) and may be the intrusive equivalent of metavolcanic rocks in the group (unit Omv).
- Zib This unit intrudes Proterozoic Y Deer Trail Group (units Ymm, Ycb₂, Yar₂, Ycb₁, and Yar₁) and Priest River Group (units Yar, Yq, and Ycb). The composition of unit Zib is similar to that of the Proterozoic Z metavolcanic rocks (unit Zmv), which suggests that unit Zib is the intrusive equivalent of the metavolcanic rocks (Campbell and Loofbourow, 1962; Miller and Clark, 1975) that were extruded approximately 760 Ma (Devlin and others, 1988).
- Yib This unit intrudes the Proterozoic Y Prichard Formation (unit Yms₁). No isotopic ages are known from unit Yib in the map area, but a U-Pb zircon age of 1,433 Ma is reported from correlative rocks (Crossport C sill of Bishop, 1973) near Bonners Ferry, Idaho (Zartman and others, 1982).
- u Because the "mineralogical gradation between the gabbroic [unit $\bar{\text{T}}\text{ib}$] and ultramafic portions of the [Chopaka intrusive] complex and the lack of small-scale crosscutting relationships supports [sic] the idea that they were emplaced at essentially the same time" (Hibbard, 1971, p. 3028), ultrabasic rocks in the **Chopaka intrusive complex** must be Triassic or older.
- Because ultrabasic rocks elsewhere in the map area are thought to be fault-bounded, their ages are uncertain.
- Mixed Metamorphic and Igneous Rocks
- TKmi **Skagit Gneiss** is a heterogeneous unit composed of rocks of at least three ages: Eocene (approximately 45 Ma) lineated granitic plugs and dikes; Late Cretaceous and Paleocene (60-75 Ma) directionless to strongly foliated leucocratic tonalite and tonalitic gneiss (correlative with units Rit, Rog, Kit, and Kog); and Triassic, Permian, and older(?) schist, quartzite, amphibolite, and calc-silicate rocks (correlative with units $\bar{\text{T}}\text{Pmc}$ and $\bar{\text{T}}\text{Pmb}$).

The **plutonic complex of Johnny George** mixed metamorphic and igneous rocks (T27-28N, R34-35E) consist of biotite gneiss of uncertain age and affinity, and leucocratic granite, aplite, and pegmatite that are probably equivalent to Upper Cretaceous and lower Tertiary intrusive rocks (unit TKia). Monzodiorite intrusions that are probably correlative with the Friedlander Meadows pluton (50 Ma) (unit Eimd) cut the mixed rocks. Formation of the complex probably occurred during intrusion of the leucocratic granite, aplite, and pegmatite in the Late Cretaceous and/or early Tertiary.

Kmg Migmatitic banded gneiss in the **Chelan complex** has yielded a weakly discordant U-Pb zircon age of 185 Ma (Mattinson, 1972); zircons probably represent a mixture of older inherited zircons and new magmatic zircons. The migmatitic gneiss probably formed "in the Late Cretaceous from melts that were derived by mobilization or anatexis of older rocks...The gradation of the more massive plutons into migmatitic terrane and the spectrum of U-Pb zircon ages representing mixtures of inherited and new components of this refractory mineral strongly support such an interpretation" (Tabor and others, 1987, p. 8).

KJmi, KJmg

These heterogeneous units are composed of rocks of at least two ages: weakly to strongly foliated, leucocratic tonalite and granodiorite that strongly resemble Upper Jurassic and Lower Cretaceous orthogneiss (unit KJog), and hornblende-biotite schist and gneiss, amphibolite, and minor calc-silicate rocks that are probably correlative with pre-Jurassic layered metamorphic rocks (units pJhm and pJam). The heterogeneous rocks are cut by Early Cretaceous plutons (unit Kit). They probably formed in the Late Jurassic or Early Cretaceous by anatexis of pre-Jurassic layered metamorphic rocks (units pJhm and pJam) or by injection of Upper Jurassic to Lower Cretaceous intrusive rocks (unit KJog) into the layered rocks (Hawkins, 1968; Menzer, 1983; Raviola, 1988; V. R. Todd, USGS, written commun., 1988).

Metamorphic Rocks

Rog The **Lake Juanita leucogneiss** has yielded a U-Pb zircon age of 59 Ma (Miller and others, 1989).

Orthogneiss in the **Oval Peak batholith** has yielded U-Pb zircon ages of 61-62 Ma (Miller and others, 1989); directionless rocks of the batholith (unit Rit) have yielded U-Pb (zircon and titanite) ages of 61 and 65 Ma, respectively (Miller and Walker, 1987; Miller and others, 1989).

The **War Creek gneiss** intrudes the Black Peak batholith (90 Ma) and is deformed along the Gabriel Peak tectonic zone, which was active between 45 and 65 Ma (Miller, 1987; Miller and others, 1989). The magmatic crystallization age is Late Cretaceous or early Tertiary.

Kog Leucocratic orthogneiss in the **Chelan complex** "is thought to have formed in the Late Cretaceous from melts that were derived by mobilization or anatexis of older rocks...The gradation of the more massive plutons into migmatitic terrane and the spectrum of U-Pb zircon ages representing mixtures of inherited and new components of this refractory mineral strongly support such

an interpretation...Metamorphic overprinting of the plutons must have occurred contemporaneously with or immediately after their emplacement if we accept the 70 to 86 m.y. U-Pb sphene ages...and the K-Ar hornblende and biotite ages as old as 84 m.y." (Tabor and others, 1987, p. 8).

Light-colored gneiss in the **heterogeneous schist and gneiss unit of the Mad River terrane** (reassigned to **rocks of the Napeequa River area** of the Chelan Mountains terrane by Tabor and others, 1989) has yielded concordant to moderately discordant U-Th-Pb zircon ages between 71 and 127 Ma (Tabor and others, 1987); "the light colored granodioritic gneiss appears to contain a mixture of newly crystallized Late Cretaceous and inherited and probable Paleozoic or older zircon" (Tabor and others, 1987, p. 11).

The **Black Peak batholith** has yielded U-Pb (zircon and titanite) ages of 89-91 Ma (Hoppe, 1984).

One of the **Bearcat Ridge plutons** has yielded a U-Pb zircon age of approximately 90 Ma (R. B. Miller, San Jose State Univ., oral commun., 1990).

The **Newman Lake Gneiss** has yielded discordant U-Pb zircon ages of 97 and 139 Ma and a Rb-Sr (whole rock + biotite) age of 45 Ma (Armstrong and others, 1987). Because the U-Pb ages and initial Sr isotopic ratios are similar to those of the Mount Spokane granite (unit Kiat), magmatic crystallization of the Newman Lake Gneiss is thought to have occurred during the Late Cretaceous (Armstrong and others, 1987).

Unnamed orthogneiss northeast of Spokane (T27-30N, R35E) grades into the Mount Spokane granite (unit Kiat) and the Newman Lake Gneiss (unit Kog); age assignment is based on interpretation of the unit as mylonitically deformed Mount Spokane granite (Weissenborn and Weis, 1976).

Kogm Mesocratic orthogneiss in the **Chelan complex** is intimately associated with, and is probably the same age as leucocratic orthogneiss (unit Kog) in the complex.

KJog Age uncertain. K-Ar biotite and hornblende and fission-track (titanite and apatite) ages between 76 and 106 Ma (Menzer, 1970; Berry and others, 1976; V. R. Todd, USGS, written commun., 1990; Bunning, 1990; Stoffel and McGroder, 1990) represent the age of uplift and cooling or age of latest ductile deformation of the plutons. A Rb-Sr isochron age of 129 Ma from the **Leader Mountain granodioritic gneiss** (Menzer, 1970) and a U-Pb zircon age of 120 Ma from the **Methow gneiss** (R. B. Miller, San Jose State Univ., oral commun., 1990) may represent magmatic crystallization ages. U-Pb zircon ages of 140-150 Ma are reported from compositionally and texturally similar intrusions in the Eagle plutonic complex in southern British Columbia (Greig, 1988).

Jog This unit intrudes the Upper Triassic(?) Loomis pluton (unit Fjgd) and Tillman Mountain tonalitic gneiss (unit Fjog); the minimum age is constrained by discordant K-Ar biotite and hornblende ages of 151 and 170 Ma (Rinehart and Fox, 1972).

- Fig** The **Tillman Mountain tonalitic gneiss** has yielded discordant K-Ar biotite and hornblende ages of 97 and 173 Ma (Stoffel, 1990b); because it is intruded by the Similkameen composite pluton (170 Ma) (unit Jia) and grades into the Upper Triassic(?) Loomis pluton (unit Figd), a Late Triassic magmatic crystallization age is suspected.
- The **Osoyoos batholith** (in southern British Columbia) has yielded a U-Pb zircon age of 202 Ma (Parkinson, 1985); a K-Ar biotite age of 49 Ma from the map area (Fox and others, 1976) indicates that post-crystallization deformation and thermal metamorphism occurred during the middle Eocene.
- Mog** Unnamed rocks west of Tonasket (T37N, R26E) intrude the Permian Anarchist Group and Triassic(?) metamorphic complex of Conconully (Rinehart and Fox, 1976).
- The **Reed Creek quartz dioritic orthogneiss** has yielded K-Ar biotite and hornblende ages of 57 and 75 Ma, respectively (Atwater and Rinehart, 1984) and Rb-Sr isochron and Pb- α zircon ages of 83 and 170 Ma, respectively (Menzer, 1983).
- Unnamed rocks of this unit elsewhere intrude the Leecher Metamorphics and Salmon Creek schists and gneisses (units pJhm and pJmb) and amphibolite and schist of Twentyfive Mile Creek (units TPhmc and TPmb); they are locally cut by Upper Cretaceous or Eocene plutons.
- TPhmc** Granofels in the **younger gneissic rocks of the Holden area** (Cater and Crowder, 1967), which are correlative with the amphibolite and schist of Twentyfive Mile Creek (Tabor and others, 1987), has yielded a discordant U-Pb (zircon) age from biotite-quartz-oligoclase granofels that "gives on older concordia intercept of 265 +/- 15 m.y. if the younger intercept is assumed to be 60-90 m.y., the age bracket for the Late Cretaceous metamorphism (Mattinson, 1972, p. 3773). The granofels has been interpreted to be a metamorphosed keratophyre...and thus the 265 m.y. (Permian) age is believed to represent the depositional age" (Tabor and others, 1987, p. 5).
- TPmb** This unit is interbedded with unit TPhmc.
- pJam** The age of both the **Basic complex** and **Ashnola Gabbro** is uncertain; metamorphic textures and compositional layering perpendicular to the regional structural trend suggest a pre-Late Cretaceous age (Daly, 1912). Age assignment is based on compositional similarity to amphibolite layers in pre-Jurassic heterogeneous rocks (unit pJhm).
- pJmb** This unit has the same age as unit pJhm.
- pJhm, pJmb** Age uncertain. The metamorphic rocks yield K-Ar ages (biotite, muscovite, and hornblende) of approximately 100 Ma (V. R. Todd, USGS, written commun., 1989; Bunning, 1990; Gulick and Korosec, 1990b). Most are intruded by Cretaceous plutons, but the **Salmon Creek schists and gneisses** and unnamed rocks north of Conconully (T37-40N, R24-25E) are intruded by the Upper Triassic(?) Loomis pluton (unit Figd), suggesting a pre-Jurassic age.
- pCam** This unit is interlayered with unit pChm; age assignment is based on interpretation of the unit as a metamorphosed equivalent of unit Yib (Miller, 1974c, 1974d; Miller and Clark, 1975).
- pCqz** Age and contact relations are uncertain. Age assignment is based on its compositional similarity to known Precambrian rocks in the region. The unit is probably correlative with part of the **Belt Supergroup** (units Yms4, Yms3, Yms2, and Yms1) and with pre-Beltian rocks.
- pChm** Age and contact relations of **gneiss near Chester Creek** and **gneiss near Round Mountain** (T24-25N, R43-45E) are uncertain. Age assignment is based on their compositional similarity to known Precambrian rocks in the region. Rb-Sr isotopic data suggest that the protolith age of the rocks is Proterozoic Y or older (Armstrong and others, 1987), which suggests that the rocks are correlative with the Belt Supergroup (units Yms4, Yms3, Yms2, and Yms1) and with pre-Beltian rocks.
- Metamorphic rocks near Priest Lake, Idaho, which are laterally continuous with the unnamed metamorphic rocks west of Newport (T30-33N, R42-44E), have yielded a Rb-Sr whole-rock age of 1,440 Ma (Armstrong and others, 1987). Because unit pChm locally grades into the Prichard Formation (unit Yms1), at least part of unit pChm probably represents thermally metamorphosed Proterozoic Y metasedimentary rocks (Miller and Clark, 1975). On the basis of Rb-Sr isotopic data, Armstrong and others (1987) suggest that the unit also includes rocks older than the Prichard Formation (pre-Belt Supergroup basement rocks).
- Age assignment of unnamed rocks south of Newport (T30N, R44-45E) is based on tentative correlation of this unit with the unnamed metamorphic rocks west of Newport (unit pChm).
- pCgn** Highly discordant U-Pb zircon ages suggest that zircon in the **Swakane Biotite Gneiss** originally crystallized at least 1,600 Ma (Mattinson, 1972). However, because the protolith of the unit is uncertain, interpretation of the age data is difficult. "If the original rock was volcanic, a Middle Proterozoic or earlier age should represent its time of formation. Many of the zircons appear rounded, however, suggesting a detrital origin. Hence, the protolith could be sedimentary, and the Precambrian age of the zircons would represent a maximum age" (Tabor and others, 1987, p. 9).
- pCbg** The **Hauser Lake Gneiss** has yielded a Rb-Sr isochron age of 2,053 Ma and a discordant U-Pb zircon age of 1,668 Ma (Armstrong and others, 1987). Near Priest River, Idaho, rocks correlative with the Hauser Lake Gneiss are intruded by augen gneiss that has yielded a Rb-Sr isochron age of 1,440 Ma (Clark, 1973). The unit is probably correlative with the Prichard Formation (unit Yms1) and/or pre-Beltian rocks.
- Age and contact relations of the **gneiss of Mica Peak** (T24-25N, R45E) are uncertain; age assignment is based on compositional similarity of this unit to known Precambrian rocks in the region. The unit is probably correlative with part of the Belt Supergroup and with pre-Beltian rocks.

pTog, pTam, pTsc, pTqz, pTmb, pThm, pTbg

Protolith ages of the metamorphic rocks in the Kettle and Okanogan metamorphic core complexes (gneiss domes) are unknown. "Specific correlation of the...rocks with sequences elsewhere in the Cordillera is difficult because the rocks in the domes are highly metamorphosed and deformed and cannot be traced out of the core into lower grade equivalents" (Orr and Cheney, 1987, p. 61). The quartzfeldspathic composition of the layered metamorphic rocks in the complexes suggests that they may be correlative with Precambrian metasedimentary rocks in the region (Cheney, 1980; Orr and Cheney, 1987). Upper Cretaceous(?) to middle Eocene plutons (units TKia, TKig, Eimd, Eiqm, and Eig) cut the metamorphic rocks in the complexes, establishing a minimum age for the metamorphic rocks (Holder and Holder, 1988; Holder and others, 1989). K-Ar biotite and hornblende ages reported from the metamorphic rocks in the core complexes range from 46 to 67 Ma (Fox and others, 1976; Engels and others, 1976; Atwater and Rinehart, 1984), indicating that uplift and cooling of the complexes occurred during the latest Cretaceous and early Tertiary. U-Pb zircon ages of 87-100 Ma from a single sample of the **Tonasket Gneiss** are the only U-Pb ages reported from the metamorphic rocks in the complexes (Fox and others, 1976); it is not clear whether these ages represent protolith or metamorphic ages.

pTgn Unnamed monzonitic and syenitic gneiss near Oroville (T39-40N, R28-29E) has yielded K-Ar biotite and hornblende ages between 50 and 58 Ma (Fox and others, 1976) and fission-track (epidote and titanite) ages of 63 and 66 Ma, respectively (Naeser and others, 1970). This unit grades into and interfingers with the Permian Anarchist Group (units Pmm and Pmcb), the Upper Triassic Osoyoos batholith (unit Tōg), and the Wauconda pluton (unit KJgd) and is cut by the Mount Bonaparte pluton (unit TKia). Gradational contacts and absence of magmatic features such as satellitic dikes and xenoliths led Fox (1973) to suggest that the gneiss formed by the alkali metasomatism of pre-existing rock units before intrusion of the Mount Bonaparte pluton.

Tectonic Rocks

tz Along the Twisp River-Foggy Dew fault (T31-34N, R19-21E), the unit consists of the mylonitically deformed Oval Peak batholith (61-65 Ma) (units Rait and Rōg), Battle Mountain and Tuckaway Lake gneisses (unit Rōg), and Twisp Valley schist (unit TPhmc) (Miller, 1987; Miller and others, 1989). Northwest of the map area, the unit includes Gabriel Peak orthogneiss (65 Ma metamorphic age), which is the deformed equivalent of the Black Peak batholith (unit Kōg) (90 Ma) (Miller, 1987). The unit is cut by the Cooper Mountain batholith (48 Ma) (unit Eigd); these relations indicate that ductile deformation occurred between 65 and 48 Ma (Miller, 1987; Miller and others, 1989).

Along the Newport fault (T31-35N, R43-45E), this unit is dominantly chloritic microbreccia; locally, it includes

chloritic breccia composed of clasts of mylonitically deformed Precambrian metamorphic rocks (unit pChm), Phillips Lake Granodiorite (100 Ma) (unit Kiat), and Silver Point Quartz Monzonite (50 Ma) (unit Eia). In the upper plate of the fault south of Tiger (T34-35N, R43E), the unit includes brecciated Addy Quartzite (unit CQz) and Metaline Formation (unit OCcb). The middle Eocene Tiger Formation (unit Ecg2) locally overlies the chloritic breccia and locally contains breccia clasts; these relations indicate deformation occurred during the middle Eocene.

Along the Omak Lake fault (T32-33N, R27-28E) this unit consists of breccia composed of clasts of the plutonic complex of Boot Mountain (unit KJmi) and Coyote Creek pluton (units TKig and TKiaa); the deformation probably occurred during the Late Cretaceous and early Tertiary.

Along the Jumpoff Joe fault (T34N, R41E), this unit consists of mylonitically and brittlely deformed Proterozoic Y metasedimentary rocks (unit Yms), Proterozoic Z metavolcanic rocks (unit Zmv), Precambrian metamorphic rocks (unit pChm), and Phillips Lake Granodiorite (100 Ma) (unit Kiat) (Miller and Clark, 1975). It is overlain by middle Eocene Tiger Formation (Ecg2); these relations indicate deformation occurred between 100 and 48 Ma.

Along the Eagle Mountain fault (T33N, R41E), this unit is composed of sheared Precambrian metasedimentary rocks (unit Yms) and Cretaceous intrusive rocks (100 Ma) (unit Kiat); the age of deformation is Late Cretaceous or younger.

Along the Pasayten fault (T34-36N, R21-22E), the unit consists of mylonitically deformed Upper Jurassic and Lower Cretaceous orthogneiss (unit KJog) and chloritic breccia composed of clasts of the mylonitic gneiss (Lawrence, 1968; V. R. Todd, USGS, written commun., 1988). Near Island Mountain (T39N, R19E), the fault is overlain or intruded by middle Eocene volcanic rocks of Island Mountain (units Ec1, Ev1, Evd1, Eian, and Eida) (White, 1986); therefore, deformation along the Pasayten fault must have occurred between the Late Jurassic and middle Eocene.

BEDROCK GEOLOGIC AND TECTONIC MAP AND TABLE SUMMARIZING THE GEOLOGIC HISTORY OF THE NORTHEAST QUADRANT

The 1:625,000-scale **bedrock geologic and tectonic map** on Sheet 3 is a simplified version of the 1:250,000-scale geologic map (Sheet 1). It graphically portrays the sedimentary, volcanic, intrusive, metamorphic, and deformational history of northeastern Washington, which is summarized in a table on Sheet 3. The colors used on this map differ from those on Sheet 1 because map units have been combined for simplicity.

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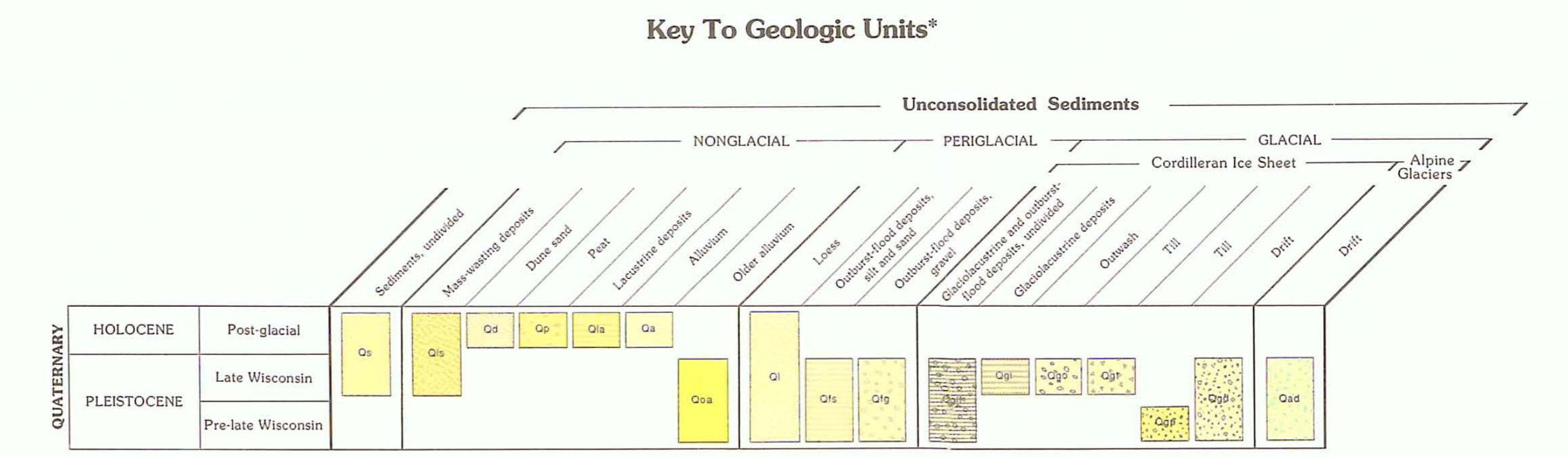
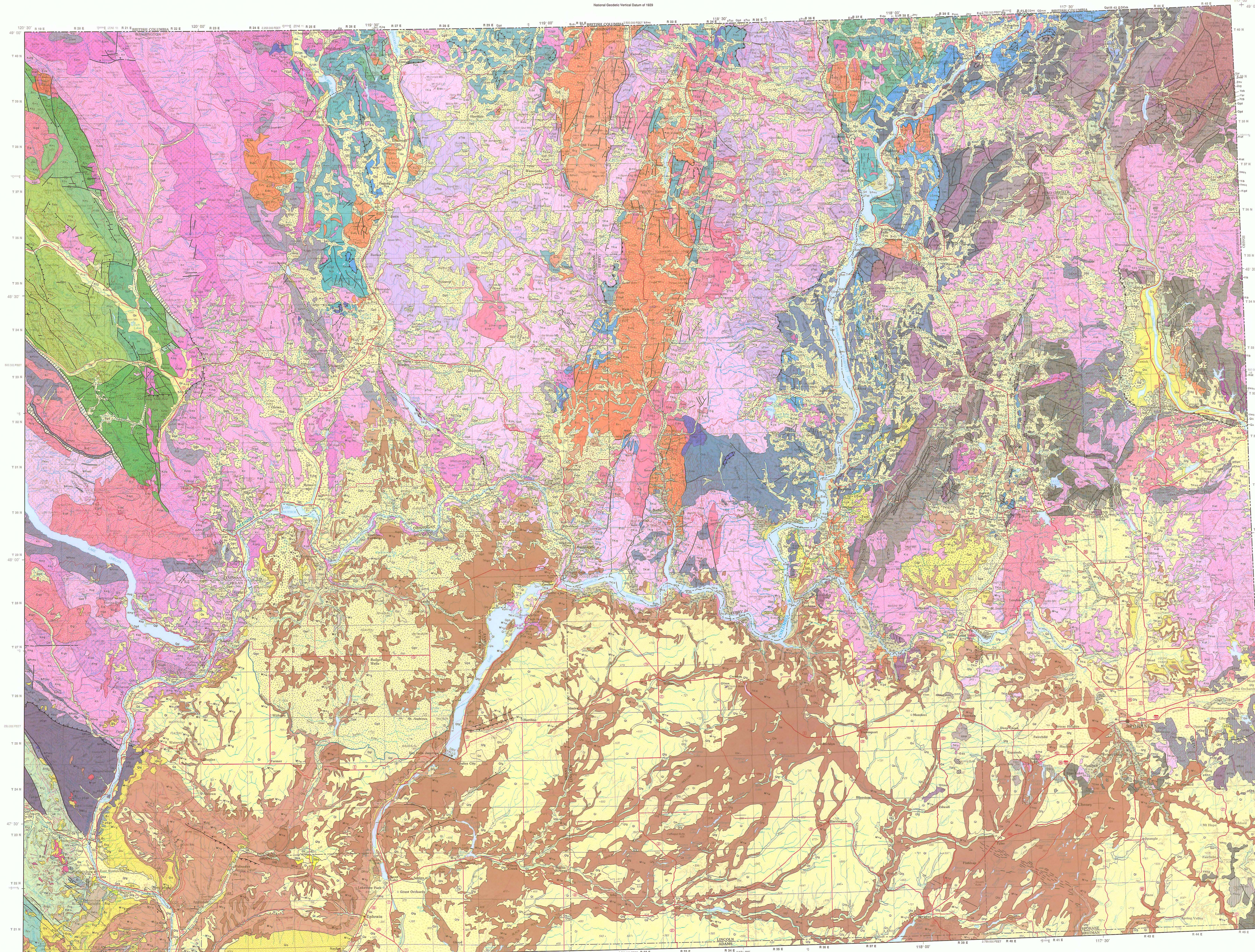
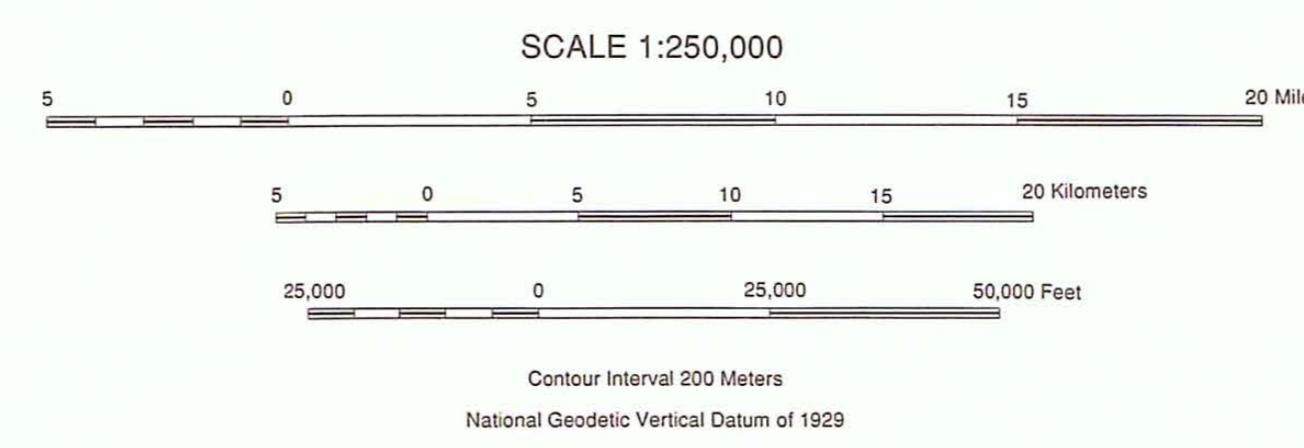
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GEOLOGIC MAP OF WASHINGTON—NORTHEAST QUADRANT

by
 Keith L. Stoffel, Nancy L. Joseph, Stephanie Zurenko Waggoner, Charles W. Gulick, Michael A. Korosec, and Bonnie B. Bunning

J. Eric Schuster, Assistant State Geologist and State Geologic Map Project Manager
 Bonnie B. Bunning and William M. Phillips, Assistant State Geologic Map Project Managers

1991



Geologic Time Period	Sedimentary Rocks		Volcanic Rocks		Metasedimentary and Metavolcanic Rocks	
	Symbol	Unit Name	Symbol	Unit Name	Symbol	Unit Name
TERTIARY	Qa	Alluvium	Ma	Mt Rainier	Ma	Mt Rainier
	Qb	Glacial drift	Ma	Mt Rainier	Ma	Mt Rainier
	Qc	Glacial drift	Ma	Mt Rainier	Ma	Mt Rainier
	Qd	Glacial drift	Ma	Mt Rainier	Ma	Mt Rainier
	Qe	Glacial drift	Ma	Mt Rainier	Ma	Mt Rainier
	Qf	Glacial drift	Ma	Mt Rainier	Ma	Mt Rainier
	Qg	Glacial drift	Ma	Mt Rainier	Ma	Mt Rainier
	Qh	Glacial drift	Ma	Mt Rainier	Ma	Mt Rainier
	Qj	Glacial drift	Ma	Mt Rainier	Ma	Mt Rainier
	Qk	Glacial drift	Ma	Mt Rainier	Ma	Mt Rainier
MESOZOIC	Tu	Triassic	Ma	Mt Rainier	Ma	Mt Rainier
	Tj	Triassic	Ma	Mt Rainier	Ma	Mt Rainier
	Tk	Triassic	Ma	Mt Rainier	Ma	Mt Rainier
	Tl	Triassic	Ma	Mt Rainier	Ma	Mt Rainier
	Tm	Triassic	Ma	Mt Rainier	Ma	Mt Rainier
	Tn	Triassic	Ma	Mt Rainier	Ma	Mt Rainier
	To	Triassic	Ma	Mt Rainier	Ma	Mt Rainier
	Tp	Triassic	Ma	Mt Rainier	Ma	Mt Rainier
	Tq	Triassic	Ma	Mt Rainier	Ma	Mt Rainier
	Tr	Triassic	Ma	Mt Rainier	Ma	Mt Rainier
PALEOZOIC	Co	Carboniferous	Ma	Mt Rainier	Ma	Mt Rainier
	Cr	Carboniferous	Ma	Mt Rainier	Ma	Mt Rainier
	Ca	Carboniferous	Ma	Mt Rainier	Ma	Mt Rainier
	Cb	Carboniferous	Ma	Mt Rainier	Ma	Mt Rainier
	Cc	Carboniferous	Ma	Mt Rainier	Ma	Mt Rainier
	Cd	Carboniferous	Ma	Mt Rainier	Ma	Mt Rainier
	Ce	Carboniferous	Ma	Mt Rainier	Ma	Mt Rainier
	Cf	Carboniferous	Ma	Mt Rainier	Ma	Mt Rainier
	Cg	Carboniferous	Ma	Mt Rainier	Ma	Mt Rainier
	Ch	Carboniferous	Ma	Mt Rainier	Ma	Mt Rainier
PRECAMBRIAN	Pr	Precambrian	Ma	Mt Rainier	Ma	Mt Rainier
	Pp	Precambrian	Ma	Mt Rainier	Ma	Mt Rainier
	Pq	Precambrian	Ma	Mt Rainier	Ma	Mt Rainier
	Pr	Precambrian	Ma	Mt Rainier	Ma	Mt Rainier
	Pp	Precambrian	Ma	Mt Rainier	Ma	Mt Rainier
	Pq	Precambrian	Ma	Mt Rainier	Ma	Mt Rainier
	Pr	Precambrian	Ma	Mt Rainier	Ma	Mt Rainier
	Pp	Precambrian	Ma	Mt Rainier	Ma	Mt Rainier
	Pq	Precambrian	Ma	Mt Rainier	Ma	Mt Rainier
	Pr	Precambrian	Ma	Mt Rainier	Ma	Mt Rainier

Geologic Time Period	Intrusive Rocks		Metamorphic and Metavolcanic Rocks	
	Symbol	Unit Name	Symbol	Unit Name
TERTIARY	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
MESOZOIC	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
PALEOZOIC	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
PRECAMBRIAN	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier
	Ma	Mt Rainier	Ma	Mt Rainier

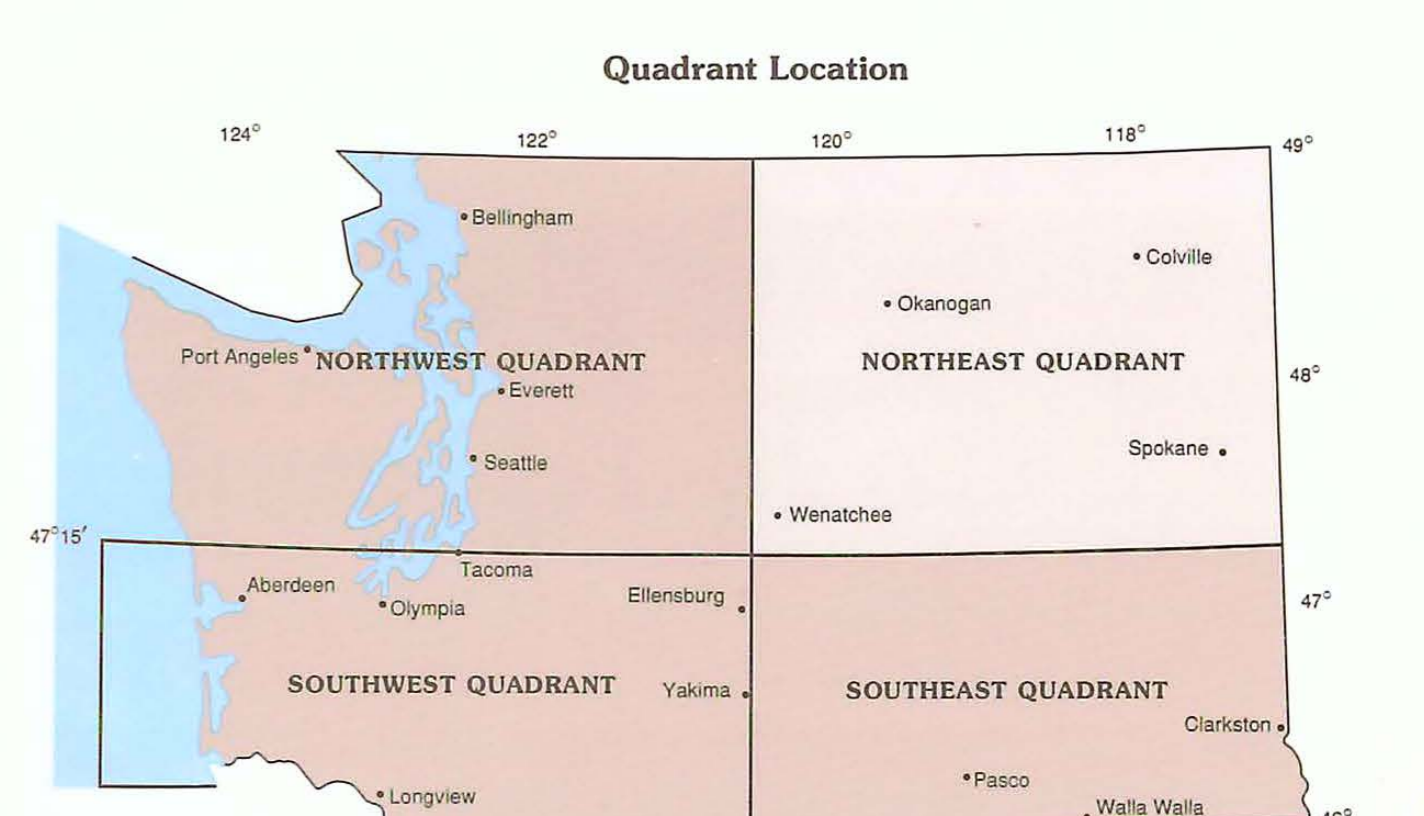
* See Sheets 2 and 3 for detailed unit descriptions and stratigraphic relations.

Geologic Symbols

- Contact
- Scratch boundary — Boundary between areas of reconnaissance and detailed mapping
- Gradational contact
- Fault — Dashed where approximately located; dotted where concealed
- Dip-slip fault — Bar and half on downthrown side; dashed where approximately located; dotted where concealed; spaced where presence or character uncertain
- Strike-slip fault — Arrows indicate direction of movement
- Low-angle normal fault — Blocks on upper plate; dotted where concealed; spaced where presence or character uncertain
- Thrust fault — Sawtooth on upper plate; dashed where approximately located; dotted where concealed; spaced where presence or character uncertain
- Anticline — Direction of plunge shown where known; dashed where approximately located; dotted where concealed
- Syncline — Direction of plunge shown where known; dashed where approximately located; dotted where concealed
- Overturned anticline
- Overturned syncline — Direction of plunge shown where known; dotted where concealed
- Monocline — Dashed where approximately located; dotted where concealed
- Approximate maximum extent of pre-late Wisconsin Coriander Ice Sheet
- Maximum extent of the late Wisconsin Coriander Ice Sheet
- Dike
- Dike swarm — Used for dike swarms of unit Ees only
- Tectonic zone
- Fracture zone

Compilation Responsibility by 1:100,000-scale Quadrangle

ROBINSON MTN Stoffel	OROVILLE Stoffel	REPUBLIC Stoffel	COLVILLE Joseph
TWISP Bunning	OMAK Gulick	NESPELEM Joseph	CHEWELAH Waggoner
CHELAN Korosec	BANKS LAKE Gulick	COULEE DAM Waggoner	SPOKANE Joseph
WENATCHEE Wells	MOSES LAKE Gulick	FITZVILLE Gulick	ROSALIA Waggoner

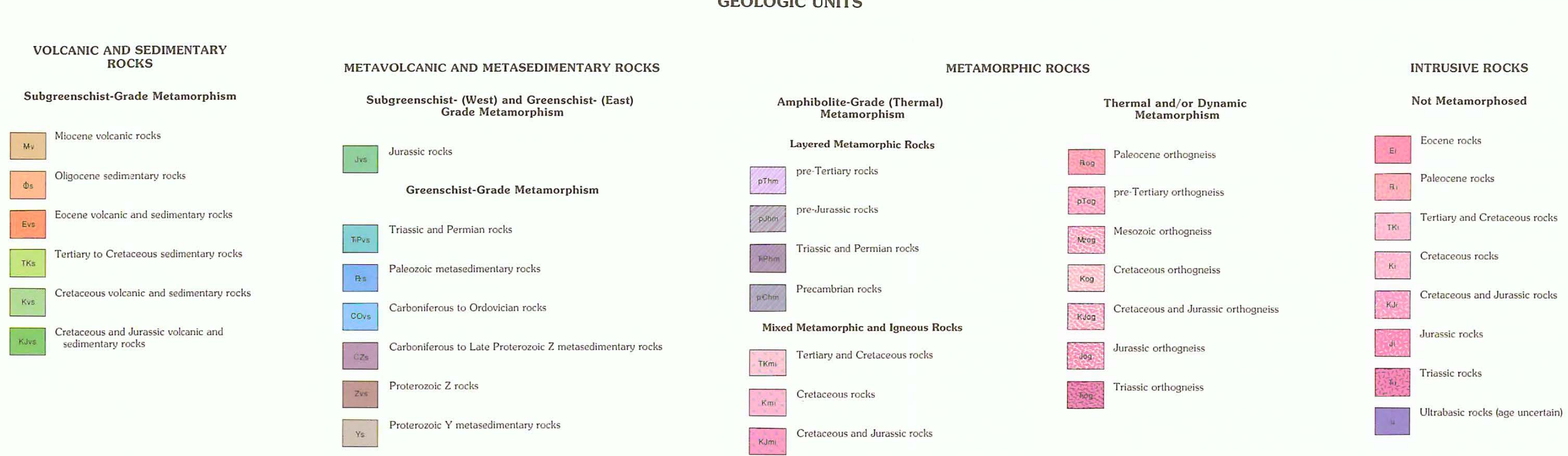
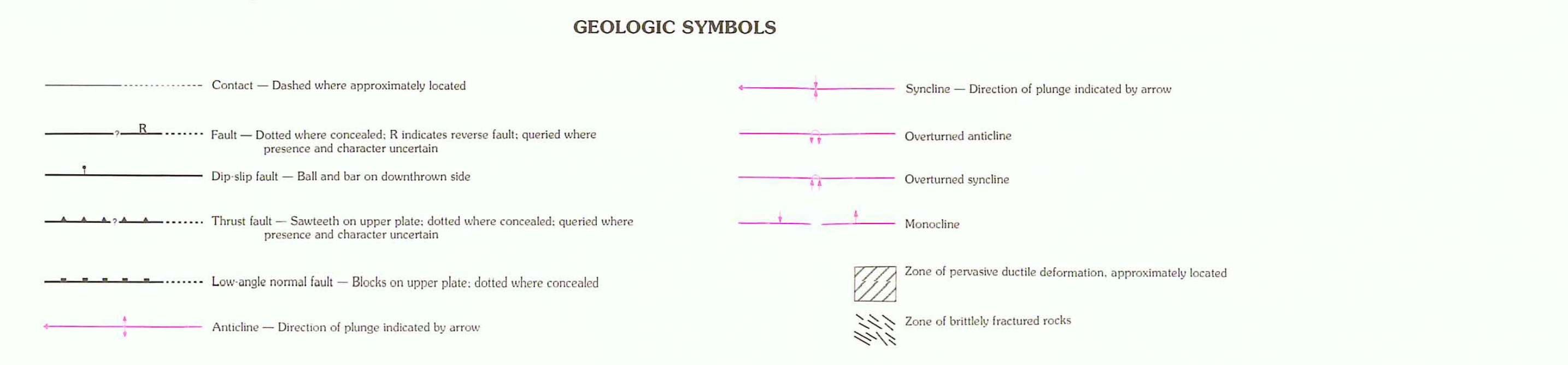
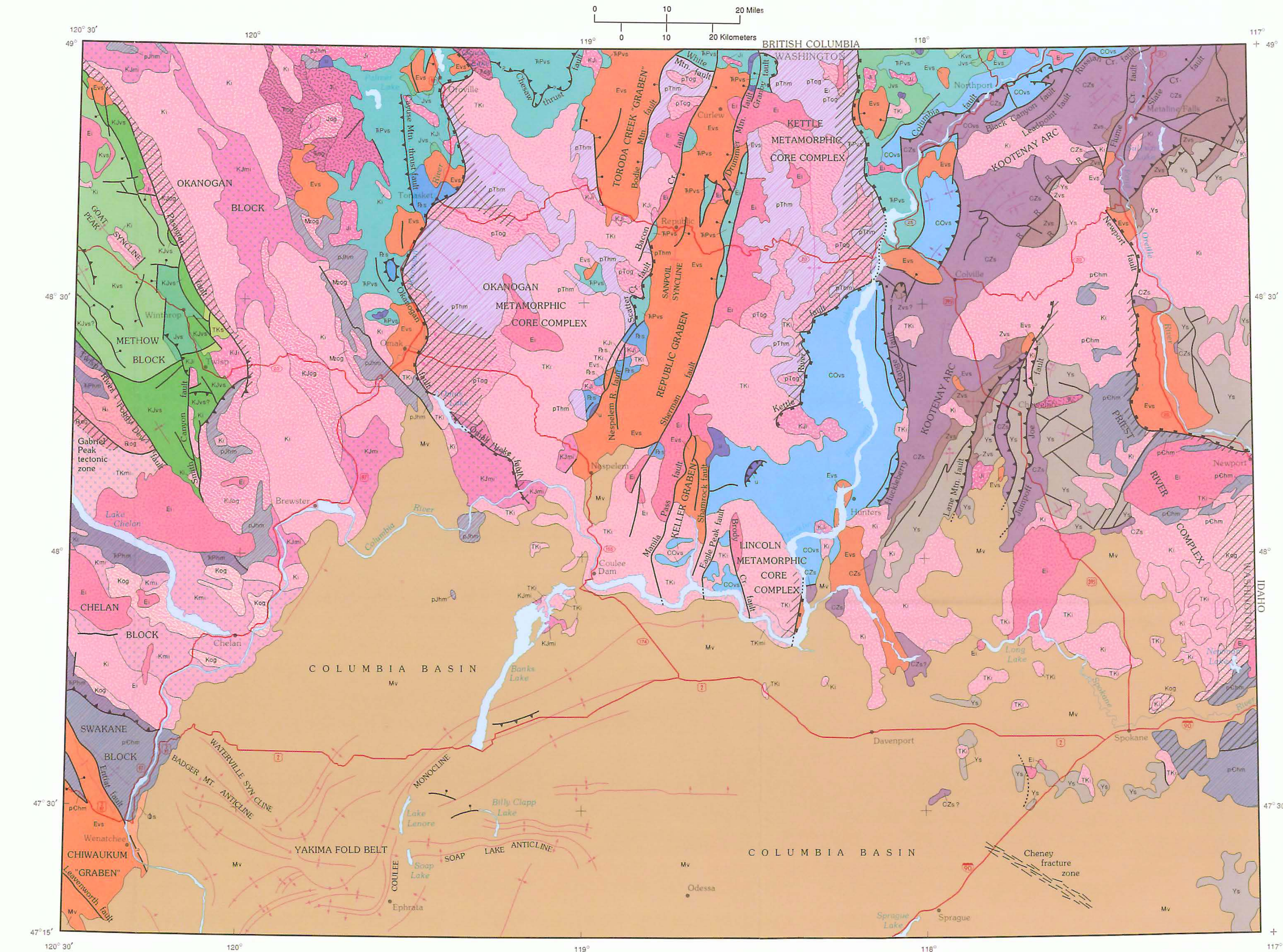


CORRELATION DIAGRAM

(To accompany geologic map, sheet 1)

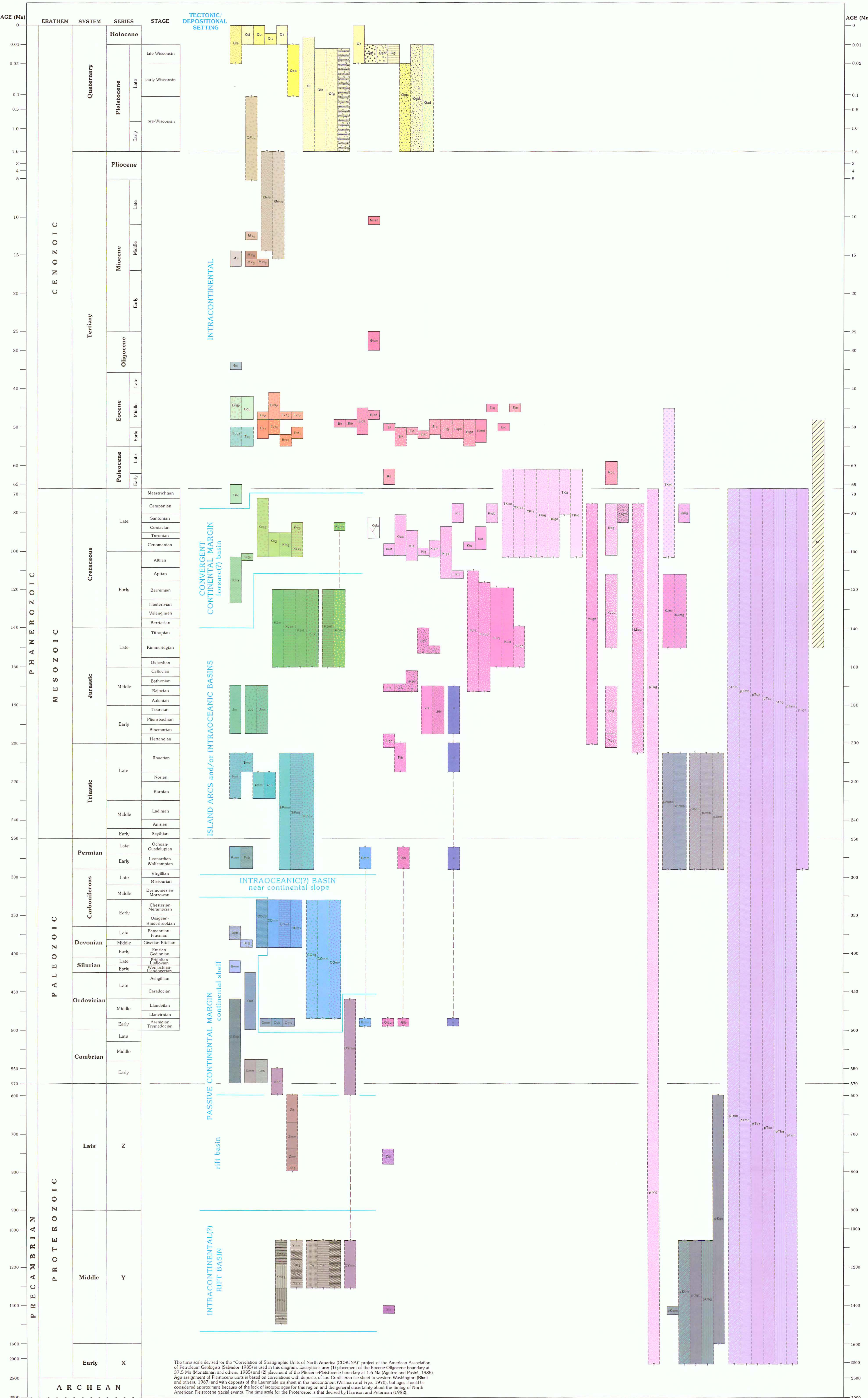
BEDROCK GEOLOGIC AND TECTONIC MAP

SCALE 1:625,000



SUMMARY OF GEOLOGIC HISTORY

AGE	VOLCANIC AND SEDIMENTARY ROCKS	INTRUSIVE ROCKS	METAMORPHISM AND DEFORMATION	TECTONIC/ DEPOSITIONAL SETTING
middle Miocene	Mv Continental flood basalt; minor continental clastic sedimentary rocks	Basaltic dikes (not shown on map)	Initial development of the Columbia Basin, Yakima fold belt, and north-northeast-trending basins (see for basalt flows) during regional subsidence and north-south(?) compression; deformation continued into the late Miocene and Pliocene	Uncertain; possible North American intracratonic back-arc or tectonic basin; possible crustal hot spot; possible incipient North American continental rift
early Oligocene	Ol Continental clastic sedimentary rocks; minor silicic volcanic rocks	Andesitic plugs and dikes (not shown on map)	?	North American intracratonic basin
early and middle Eocene	Evs Andesitic to rhyolitic volcanic and volcanoclastic rocks; continental and shallow to deep-water, marine clastic sedimentary rocks (see source)	Ei Epicentral calc-alkalic and alkalic intrusive rocks	Regional east-west extension; rapid uplift of the Swakane block and Okanogan, Methow, Lincoln, and Priest River metamorphic core complexes accompanied by low-angle, normal ("detachment") faulting (both ductile and brittle deformation) and broad folding; formation of volcano-tectonic depressions ("grabens") by high- and low-angle, normal faulting, strike-slip faulting, and broad folding(?)	Transition from compressional magmatic arc to North American intracratonic basin; back-arc extensional basin
Paleocene and Late Cretaceous	Tk, Kv Andesitic to rhyolitic volcanic and volcanoclastic rocks; continental and shallow to deep-water, marine clastic sedimentary rocks (see source)	It, Iv Mesozonal and epicentral calc-alkalic intrusive rocks	Continued development of the Methow block accompanied by strike-slip and dip-slip ductile deformation along the Trippe River-Foggy Dew fault and dip-slip and strike-slip movement on the Pasayten fault; amphibolite-grade metamorphism, migmatization, and ductile deformation of older rocks in the Chelan block (units Tpm, Tcm, and Tml) and Okanogan, Kettle, Lincoln, and Priest River metamorphic core complexes (units Kpm and Kpl) accompanied by intrusion of syntectonic plutons (Bsp, Ksp, and Tpl)	Convergent continental margin; forearc(?) basin and magmatic arc along margin of North American continent
Early Cretaceous	Kv, Kav Andesitic to rhyolitic volcanic and volcanoclastic rocks; shallow to deep-water, marine clastic sedimentary rocks (see source)	Ki K in Okanogan block, only Mesozonal calc-alkalic intrusive rocks	Initial development of the Methow block by folding, thrust faulting, and dip-slip faulting; formation of the Okanogan block during amphibolite-grade metamorphism, migmatization, and ductile deformation of older rock units (see and Kpl) and intrusion of syntectonic plutons (Ksp)	Convergent continental margin; forearc(?) basin along margin of North American continent
Jurassic	Jv, Jvs Basaltic to rhyolitic volcanic and volcanoclastic rocks; shallow to deep-water, marine clastic sedimentary rocks (see source)	Ji Gabbroic plugs and dikes (west of "Northport" and calc-alkalic intrusive rocks)	Continued(?) development of the "Kootenay arc" during Early Jurassic folding and thrust faulting that accompanied accretion of Jurassic, Triassic(?) and Permian(?) allochthonous terranes to the North American continent	Island and magmatic arcs
Triassic and Permian	Tp, Ps Basaltic to dacitic volcanic and volcanoclastic rocks; shallow to deep-water, marine clastic and igneous rocks (see source)	Ti Gabbroic plugs and dikes (not shown on map); Late Triassic calc-alkalic intrusive rocks	Initial development of the "Kootenay arc" during Late Triassic folding and thrust faulting that accompanied accretion of Triassic and Permian allochthonous terranes to the North American continent; probably regional amphibolite-grade metamorphism and migmatization of older rocks (units tpm and tpsm)	Island arc and/or intra-oceanic basin; location relative to North American continent during time of deposition uncertain
Carboniferous to Ordovician	Ord Shallow to deep-water, marine clastic sedimentary rocks (both continental and arc related) and carbonate rocks; basaltic volcanic rocks	Gabbroic plugs and dikes (not shown on map)	Regional subsidence and extension	Intra-oceanic extensional basin near continental slope; location relative to North American continent during time of deposition uncertain
Carboniferous to Late Proterozoic Z	Cz Shallow to deep-water, marine clastic sedimentary rocks (continental source) and carbonate rocks; basaltic volcanic rocks	None	Regional subsidence	North American passive continental margin, shelf
Proterozoic Z	Zs Marine (possibly) some glaciomarine sedimentary rocks (continental source) and carbonate rocks; basaltic volcanic rocks	Gabbroic sills (not shown on map)	Regional extension and subsidence; high-angle dip-slip and possible strike-slip faults; possible Late Proterozoic(?) metamorphism of Proterozoic X and/or Archaean rocks	North American continental margin rift basin
Proterozoic Y	Ys and part of ysm in the Priest River Complex; Murine and/or continental clastic sedimentary rocks (continental source); carbonate rocks; enoplistic	Gabbroic sills (not shown on map)	Regional extension and subsidence; possible strike-slip faults	North American intracratonic(?) rift basin
Proterozoic Z or Y	Zm in the Swakane block Dacitic volcanic rock(?) or clastic sedimentary rock(?) (subsequently metamorphosed to amphibolite grade)	?	?	Uncertain; location relative to North American continent during time of deposition uncertain
Proterozoic X and/or Archaean	Part of ysm in the Priest River Complex; Crystalline sedimentary rocks (subsequently metamorphosed to amphibolite grade)	?	?	?



The time scale devised for the "Correlation of Stratigraphic Units of North America (COSUNA)" project of the American Association of Petroleum Geologists (Bakker, 1982) is used in this diagram. Exceptions are: (1) placement of the Eocene-Oligocene boundary at 37.5 Ma (Montañez and others, 1982 and 1983) and (2) placement of the Pliocene-Pleistocene boundary at 1.5 Ma (Alger and Peck, 1986). Age assignment of Proterozoic units is based on correlations with deposits of the Columbia ice sheet in western Washington (Burr and others, 1987) and with deposits of the Laurentide ice sheet in the midcontinent (Wilford and Figs., 1976), but ages should be considered approximate because of the lack of isotopic ages for this region and the general uncertainty about the timing of North American Pleistocene glacial events. The time scale for the Proterozoic is that devised by Harrison and Peterman (1982).

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